DOI: 10.1111/gcb.16456

RESEARCH ARTICLE

Soil organic matter molecular composition with long-term detrital alterations is controlled by site-specific forest properties

Laura Castañeda-Góme[z1](#page-0-0) | **Kate Lajtha[2](#page-0-1)** | **Richard Bowde[n3](#page-0-2)** | **Fathima Nahidha Mohammed Jauha[r1](#page-0-0)** | **Juan Jia[4](#page-0-3)** | **Xiaojuan Fen[g4](#page-0-3)** | **Myrna J. Simpson¹ ®**

¹Department of Physical and Environmental Sciences, University of Toronto, Toronto, Ontario, Canada

²Department of Crop and Soil Sciences, Oregon State University, Corvallis, Oregon, USA

3 Department of Environmental Science and Sustainability, Allegheny College, Meadville, Pennsylvania, USA

4 State Key Laboratory of Vegetation and Environmental Change, Institute of Botany, Chinese Academy of Sciences, Beijing, China

Correspondence

Myrna J. Simpson, Department of Physical and Environmental Sciences. University of Toronto, Toronto, ON, Canada. Email: myrna.simpson@utoronto.ca

Funding information Natural Sciences and Engineering Research Council of Canada

Abstract

Forest ecosystems are important global soil carbon (C) reservoirs, but their capacity to sequester C is susceptible to climate change factors that alter the quantity and quality of C inputs. To better understand forest soil C responses to altered C inputs, we integrated three molecular composition published data sets of soil organic matter (SOM) and soil microbial communities for mineral soils after 20 years of detrital input and removal treatments in two deciduous forests: Bousson Forest (BF), Harvard Forest (HF), and a coniferous forest: H.J. Andrews Forest (HJA). Soil C turnover times were estimated from radiocarbon measurements and compared with the molecularlevel data (based on nuclear magnetic resonance and specific analysis of plant- and microbial-derived compounds) to better understand how ecosystem properties control soil C biogeochemistry and dynamics. Doubled aboveground litter additions did not increase soil C for any of the forests studied likely due to long-term soil priming. The degree of SOM decomposition was higher for bacteria-dominated sites with higher nitrogen (N) availability while lower for the N-poor coniferous forest. Litter exclusions significantly decreased soil C, increased SOM decomposition state, and led to the adaptation of the microbial communities to changes in available substrates. Finally, although aboveground litter determined soil C dynamics and its molecular composition in the coniferous forest (HJA), belowground litter appeared to be more influential in broadleaf deciduous forests (BH and HF). This synthesis demonstrates that inherent ecosystem properties regulate how soil C dynamics change with litter manipulations at the molecular-level. Across the forests studied, 20 years of litter additions did not enhance soil C content, whereas litter reductions negatively impacted soil C concentrations. These results indicate that soil C biogeochemistry at these temperate forests is highly sensitive to changes in litter deposition, which are a product of environmental change drivers.

KEYWORDS

cutin, DIRT, lignin, nuclear magnetic resonance, phospholipid fatty acids, radiocarbon, soil carbon, suberin

1 | **INTRODUCTION**

Forest ecosystems store as much as 40% of the global soil carbon (C) stocks (Janzen, [2004](#page-13-0); Mayer et al., [2020](#page-14-0); Wei et al., [2014\)](#page-15-0) over 4 billion hectares of the Earth's land surface (FAO & UNEP, [2020](#page-13-1)). The future of these soil C reserves is threatened by many environmental changes that impact the dynamic balance of the soil C pool (Lal, [2005](#page-14-1); Reichstein, [2007](#page-15-1); Song et al., [2019](#page-15-2)), most of which occur via altered plant C inputs to the soil. For example, elevated atmospheric carbon dioxide $(CO₂)$ levels are often associated with increases in plant biomass that generally occur at the expense of reduced soil C stocks (Terrer et al., [2021](#page-15-3)). Elevated $CO₂$ conditions can also impact plant C and nutrient allocation between plant shoots and roots, thus changing the quantity and quality of the above- and belowground litter inputs to the soil (Gifford et al., [2000](#page-13-2); Pendall et al., [2004](#page-14-2)). Similarly, other environmental factors such as increasing temperatures (Kirschbaum, [2000](#page-13-3); Smith et al., [2008](#page-15-4); Zhao et al., [2017\)](#page-16-0), and higher nutrient availability (Matson et al., [1999](#page-14-3); Quinn Thomas et al., [2010;](#page-15-5) Yan et al., [2020](#page-16-1)) may alter plant productivity, plant litter decomposability and soil organic matter (SOM) decomposition with long-lasting impacts on the soil C pool of forests. Decreases in soil C with litter removal and increases with enhanced litter inputs seem to be a common response of most ecosystems across the globe (Feng et al., [2022](#page-13-4); Xu et al., [2013](#page-15-6), [2021\)](#page-15-7). However, temperate forests from the detrital input and removal treatment (DIRT) network exhibit distinct responses to litter additions, showing instead decreases in soil C even after 20 years of doubled litter additions (Lajtha et al., [2018\)](#page-14-4). The lack of C accumulation at these sites has been attributed to positive soil priming, referring to the microbial stimulation of SOM decomposition prompted by the fresh litter additions (Kuzyakov et al., [2000](#page-14-5); Lajtha et al., [2018;](#page-14-4) Pisani et al., [2016](#page-15-8); Wang et al., [2017](#page-15-9)). The conflicting results from DIRT temperate forests demonstrate that the relationship between ecosystem productivity, plant litter inputs, and soil C accumulation is not always linear nor easily predictable due to the many factors controlling SOM dynamics. Xu et al. [\(2021\)](#page-15-7) list the dominant tree species and their mycorrhizal symbionts as well as soil type and N availability as potential drivers of the responses of temperate forests to litter additions. A recent metanalysis by Feng et al. ([2022](#page-13-4)) suggested that the relatively higher initial soil C status of these forests before litter manipulations and the soil C saturation capacity along with comparably lower litter input rates at DIRT sites drive the lack of soil C increases with doubled litter inputs. Current global estimations by Georgiou et al. ([2022](#page-13-5)) confirm the relatively high C saturation and low C accrual rates for these DIRT temperate forests in comparison with other temperate forests (Figure [1c](#page-2-0)). Although C is not accumulating in soil with doubled litter inputs in temperate DIRT forests, the SOM molecular composition and soil microbial communities might be responding to higher litter inputs. Similarly, decreases in soil C with litter removals can have an impact on SOM molecular composition and soil microbes in unique ways that are not captured in routine soil C measurements. Accounting for the complex nature of SOM can shed light on the mechanisms driving the observed

responses to litter manipulations in these forests as distinct components of the SOM pool and soil microbial communities may respond differently to global climate change factors, ecosystem properties, and management (Feng & Simpson, [2011](#page-13-6); Kögel-Knabner, [2017](#page-14-6); Simpson & Simpson, [2012](#page-15-10)).

The DIRT network has experimental sites spanning across climatic conditions, forests, and soil types (Lajtha et al., [2018](#page-14-4); Nadelhoffer et al., [2004](#page-14-7)). Here, we focus on three forests with different vegetation composition for which SOM molecular-level data sets and soil radiocarbon measurements in mineral surface soils (0–10 cm) after 20 years of litter manipulation are available. Two temperate deciduous forests, Bousson Forest (BF) in Pennsylvania and Harvard Forest (HF) in Massachusetts, and H.J. Andrews Forest (HJA), a coniferous temperate forest in Oregon (Figure [1b\)](#page-2-0). Despite contrasting ecosystem and soil properties (Figure [1c\)](#page-2-0), soil C responses were similar with double litter additions whereby soil C did not increase in comparison with the control plots (CTs) while it decreased with litter removal treatments (above-, belowground, or both) (Bowden et al., [2014](#page-12-0); Lajtha et al., [2014](#page-14-8); Pierson et al., [2021](#page-14-9)). Across all sites, the exclusion of aboveground C inputs led to slightly higher C losses, whereas the removal of belowground C inputs did not have impacts on soil C, which points toward a greater control of aboveground litter inputs on soil C dynamics (Bowden et al., [2014;](#page-12-0) Lajtha et al., [2014](#page-14-8); Pierson et al., [2021](#page-14-9)) despite the general idea that root inputs are more persistent in soil, thus having a greater contribution to stable soil C (Angst et al., [2016](#page-12-1); Lützow et al., [2006](#page-14-10); Rasse et al., [2005](#page-15-11); Spielvogel et al., [2014](#page-15-12)). Other lines of evidence suggest that root-derived compounds degrade in soil at rates similar to leaf-derived compounds (cutin) (Angst et al., [2016](#page-12-1)) or that suberin can decompose faster than cutin (Feng et al., [2008](#page-13-7)) and that the preferential use of root- versus leaf- derived compounds can be site-specific (Crow et al., [2009](#page-13-8)).

Our objective with this synthesis is to assess changes in the response of the molecular composition of SOM and soil microbial communities to altered litter inputs across temperate forests in the Northern hemisphere. These temperate forest sites exhibit a gradient of N availability, atmospheric N deposition rates, and aboveground litter chemistry, with higher N availability at BF, followed by HF and lastly, HJA (Figure [1c](#page-2-0)). This fertility gradient can influence the responses of soil microbial community composition, the molecular biogeochemistry of the SOM and their responses to litter manipulations. Despite the similarities in the response of soil C to litter additions and removals after 20 years of manipulations across sites, we postulate that SOM components as well as soil microbial communities will have unique responses to altered litter inputs that will be partly determined by the availability of N at each site. Generally, higher N availability is expected to favor bacterial growth over fungal growth as bacteria—as *r*-strategists—can more rapidly degrade high quality (low C:N) substrates (Feng et al., [2010](#page-13-9); Hicks et al., [2021](#page-13-10); Leite et al., [2017;](#page-14-11) Rousk et al., [2011](#page-15-13); Rousk & Frey, [2015](#page-15-14); Wallenstein et al., [2006](#page-15-15)). Thus, we expect that doubled high-quality (low C:N) litter additions in sites with higher soil N availability and bacteriadominated microbial communities will lead to higher SOM decomposition, soil microbial biomass, and more significant decreases in

FIGURE 1 (a) Core experimental design of the detrital input and removal treatment network. Ecosystem (b) and soil properties (c) at the Forest sites. MAP: mean annual precipitation, MAT: mean annual temperature, N: nitrogen. In (b), percentages in parentheses represent the relative abundance of each species at each site and its litter C:N ratio, whereas *x̅*C:N represents the weighted averages of leaf C:N of the dominant tree species. Data from ^aBowden et al. [\(2014\)](#page-12-0), ^bCrow et al. ([2009](#page-13-8)), ^cFanin et al. ([2020](#page-13-11)), ^dLajtha et al. ([2014\)](#page-14-8), ^eUC Davis Soil Resource Laboratory ([2021](#page-15-17)), ^fPierson et al. [\(2021\)](#page-14-9), ^gMan et al. ([2022](#page-14-12)), ^hBF 20 year harvest unpublished data, ⁱSchwaner and Kelly ([2019](#page-15-18)), ^jAdams and Angradi (1996[\)](#page-16-2) ^kCorrigan (2008) ^IKing et al. (2001) ^mParsons et al. (2004), "Stanek et al. (2020), ^oPerakis et al. (2012), ^pYano et al. (2005),
^gEdmonds (1980), 'Edmonds and Thomas (1995), ^sAsker et al. (2003 Edmonds [\(1980\)](#page-13-13), ^rEdmonds and Thomas [\(1995](#page-13-14)), ^sAcker et al. ([2003](#page-15-20)), ^tVanderbilt et al. (2003), ^u Georgiou et al. ([2022](#page-13-5)).

concentrations of plant-derived compounds in soils in comparison with fungal-dominated sites with lower N availability and doubled low-quality (high C:N) litter inputs. Soil Δ^{14} C values reflect the time since C atoms were fixed by plants (Trumbore, [2009](#page-15-16)) and so, with doubled fresh litter inputs, Δ^{14} C values might increase as fresh litter accumulates in soil. However, congruent with the observed positive soil priming at these sites (Lajtha et al., [2018\)](#page-14-4), soil Δ^{14} C values might instead decrease with doubled litter inputs as fresh litter is rapidly decomposed, a phenomenon that might be more marked in the high N sites where litter C:N is lower and litter inputs are processed more rapidly.

Litter exclusion treatments, on the other hand, would lead to the accumulation of more processed and complex C compounds, thus further increasing SOM decomposition state accompanied by

246 | WILEY- Global Change Biology **| CASTAÑEDA-GÓMEZ ET AL.**

decreases in cutin-, suberin-, and lignin-derived compounds as all litter inputs are reduced, along with lower microbial biomass and lower Δ^{14} C values. Yet, the impact of litter removal inputs should be stronger on high N availability sites as soil microbial communities at these sites can be more dependent on the low C:N plant-derived inputs than soil microbial communities from low N availability sites, more adapted to low availability of preferred C sources and lownutrient substrates. Finally, although higher contributions to soil C accumulation from roots versus aboveground litter are generally expected, the available evidence from these DIRT sites shows that belowground litter reductions do not significantly alter soil C, whereas aboveground litter reductions decreased soil C (Lajtha et al., [2018](#page-14-4)). Therefore, this should be reflected in higher SOM decomposition state, increase degradation of lignin-, cutin-, and suberin-derived compounds and lower Δ^{14} C values for aboveground litter removal plots and not belowground litter removal plots across all sites. However, Crow et al. ([2009](#page-13-8)) concluded from the 10 year DIRT analysis that root-derived C had greater contribution to stable C pools in deciduous forests, such as BF, whereas leaf-derived C defined in greater measure the stable C pools in coniferous forests, such as HJA. With our unique SOM molecular-level multiproxy approach, we aim to ascertain what other indicators point out to a preferential persistence of root-derived C in deciduous forests, whereas leafderived C persistence in coniferous forests.

2 | **MATERIALS AND METHODS**

2.1 | **DIRT experimental design**

The core DIRT experimental design includes the variation in above and belowground inputs via: doubled aboveground litter additions (DL), no aboveground litter (NL), no roots (NR) achieved by trenching, and no C inputs (NI) (Figure [1a](#page-2-0)). CTs receive ambient detrital inputs. Replicated plots (3 $m \times 3$ m for BF and HF and 10 $m \times 15$ m for HJA. $N = 3$) are located beneath the forest canopy. Aboveground litter is excluded in NL plots with a mesh fabric, and the collected litter is added to the DL plots. Root growth in NR plots is prevented by the installation of inert barriers, while the NI treatment is a combination of NL and NR plots (Nadelhoffer et al., [2004](#page-14-7)).

2.2 | **Molecular composition of SOM determined by targeted compound analysis and nuclear magnetic resonance spectroscopy**

Targeted analysis of numerous plant- and microbial-derived compounds from BF, HF, and HJA DIRT experiments collected after 20 years of experiment was compiled into one single database. The detailed protocols and original data sets for these analyses are published and discussed in detail in Pisani et al. ([2016](#page-15-8)), Wang et al. ([2017](#page-15-9)) and Man et al. ([2022](#page-14-12)). Briefly, soil cores were collected from each experimental plot (two for BF and HF, and six for HJA)

and combined into a composite sample, freeze-dried, sieved to 2 mm, and ground for sequential chemical extractions, microbial phospholipid fatty acid (PLFA) analyses, and solid-state 13 C nuclear magnetic resonance (NMR) spectroscopy. For HF and HJA samples, sequential chemical extractions and PLFA analyses were conducted for two analytical replicates per composite sample and averaged for each plot (*n* = 3). For BF, there were no analytical replicates, and all analyses were performed on one composite sample per plot (*n* = 3). For all sites, NMR analyses were performed on a single composite sample prepared per treatment and CTs. Previous published articles and databases from all sites were consulted to obtain information about the ecosystems and soil properties (Figure [1b,c](#page-2-0)).

Sequential chemical extractions included solvent extraction followed by base hydrolysis and copper (II) oxide (CuO) oxidation (Otto et al., [2005](#page-14-15); Otto & Simpson, [2006](#page-14-16), [2007](#page-14-17)). Total solvent extractable compounds included microbial-derived short-chain aliphatic and plant-derived, long-chain aliphatic compounds, total acyclic and cyclic lipids as well as total sugars, total free lipids, and the plantderived steroid ratio calculated as degradation products over their precursors with higher values indicating increased degradation (Otto & Simpson, [2005](#page-14-18)) (Table [S1\)](#page-16-3). Variables from base hydrolysis included leaf/needle-derived cutin, root-derived suberin, and microbialderived lipids (Cai et al., [2017](#page-12-5); Jia et al., [2019\)](#page-13-15) as well as the suberin to cutin ratio, that increases with preferential degradation of cutin over suberin (Bahri et al., [2006](#page-12-6); Wang et al., [2018](#page-15-21)); the cutin degradation ratio ($\omega C_{16}/C_{16}$) as the relative change of C_{16} ω-hydroxy-alkanoic acids to all hydrolysable C_{16} fatty acids, increasing with higher cutin degradation (Goñi & Hedges, [1990](#page-13-16); Otto & Simpson, [2006](#page-14-16)); and the mid-chain-substituted acids to total suberin and cutin ratio (Σmid/ Suberin + Cutin) that decreases with higher cutin and suberin degradation (Otto & Simpson, [2006\)](#page-14-16) (Table [S2\)](#page-16-3). Lignin-derived compounds isolated from CuO oxidation include: vanillyls, syringyls, cinnamyls, total sum of vanillyls, syringyls, and cinnamyls, and ratios such as syringyls/vanillyls and cinnamyls/vanillyls that characterize the predominant botanical origin of lignin in soil. Finally, lignin oxidation ratios: acid-to-aldehyde ratio for lignin-derived vanillyls $[(Ad/A)_{\nu}]$ and syringyls $[(Ad/AI)_s]$ increase with higher lignin decomposition (Hedges & Mann, [1979\)](#page-13-17) (Table [S3\)](#page-16-3).

Microbial PLFAs were extracted using a modified Bligh–Dyer method (Bligh & Dyer, [1959](#page-12-7); Frostegård & Bååth, [1996\)](#page-13-18) and from this analysis, we compiled information of the total microbial biomass and key microbial groups including fungi, Actinobacteria, Gram-positive, and Gram-negative bacteria. Microbial community composition ratios were also compiled across all sites. These ratios included the fungi-to-bacteria ratio; and the Gram-negative to Gram-positive and Actinobacteria ratio. The Gram-negative to Gram-positive bacteria ratio is used as an indicator of C availabil-ity in soils (Fanin et al., [2019\)](#page-13-19). Here, we added Actinobacteria and Gram-positive bacteria in the denominator as these microbial groups are known to preferentially access complex forms of C in soil (Deng et al., [2015;](#page-13-20) Ghai et al., [2014;](#page-13-21) Su et al., [2020](#page-15-22)). Finally, we also included the microbial stress ratio, a measurement of microbial stress, calculated as cyclopropane PLFA to its monoenoic precursor since

higher cyclopropane PLFA is produced with substrate limitations or other stressors such as changes in temperature or moisture (Bossio & Scow, [1998\)](#page-12-8). In this study, we compare stress ratios with ambient conditions (CTs) and anticipate that a higher microbial stress ratio is indicative of preferred substrate C availability (Table [S4](#page-16-3)) or changes in community composition and structure (Frostegård et al., [2011\)](#page-13-22).

Finally, NMR spectroscopy variables correspond to four main chemical shift regions: alkyl C (0–50 ppm), *O*-alkyl C (50–110 ppm), aromatic and phenolic C (110–165 ppm), and carboxyl and carbonyl C (165–215 ppm). The ratio of alkyl/*O*-alkyl C was calculated as a measurement of the SOM decomposition state, increasing with higher SOM degradation since *O*-alkyl C components are preferentially degraded over alkyl C components (Baldock & Preston, [1995](#page-12-9)). A second ratio was calculated with alkyl C + aromatic and phenolic C over *O*-alkyl C + carboxylic and carbonyl C as a complementary measurement of SOM decomposition. The resistance to decomposition ratio also increases with higher utilization of preferred substrates such as cellulose (*O*-alkyl) and carbonyl and carboxyl C versus less preferred substrates such as plant-derived lipids (alkyl C) and lignin (aromatic and phenolic C) (Ostertag et al., [2008](#page-14-19)). For consistency, all data and ratios were revised from raw data to obtain a homogenous database for the purposes of this study. All variables, except those from NMR analysis, are expressed in μg/g dry soil.

2.3 | **Soil radiocarbon analysis and estimation of SOM turnover times**

Bulk SOM radiocarbon data were obtained and estimated C turnover time calculated to examine changes in SOM dynamics. Fraction modern $C(F^{14}C)$ was determined on composite samples of combined field replicates per treatment and forest site $(n = 15)$ using a 3MV tandem accelerator mass spectrometer by the A.E. Lalonde AMS Laboratory in Ottawa, Ontario. $F^{14}C$ was calculated based on Reimer et al. ([2013](#page-15-23)) as the ratio of the sample ${}^{14}C/{}^{12}C$ ratio to the standard $14C/12C$ ratio (Oxalic acid, Ox-II) measured concurrently with our samples. Δ^{14} C values were calculated by correcting this ratio by the radiocarbon decay of the standard up to the year of sample collection: 2011 (BF), 2010 (HF) and 2017 (HJA) (Schuur et al., [2016](#page-15-24)). The ${}^{14}C/{}^{12}C$ ratios of both samples and standards were backgroundcorrected, and the result further corrected for spectrometer and preparation fractionation using AMS measured ${}^{13}C/{}^{12}C$ ratio and normalized to δ^{13} C (PDB) following Crann et al. [\(2017](#page-13-23)). To determine carbon turnover times (in years, equal to the inverse of the turnover rate, *k*), we assumed a steady state system and used the following equation from Torn et al. ([2009](#page-15-25)):

$$
C_t \times F'_{C,t} = I \times F'_{atm,t} + C_{t-1} \times (1 - k - \lambda) \times F_{C,t-1}
$$
 (1)

where C*^t* and C*t* − 1 are organic carbon contents of the soil sampled in year *t* and 1 year before *t* − 1. *I* is plant-derived carbon input during the sampling year, $F'_{C,t}$ is the corrected measured radiocarbon of each sample by radiocarbon decay (14 C/1000+1), $F'_{\text{atm},t}$ is the atmospheric

 CASTAÑEDA-GÓMEZ et al. **[|] 247**

are governed by the applicable Creative Commons License

 F^{14} C values from 1950 to the year of sampling and λ is the radioactive decay constant of 14 C = 1.210 \times 10⁻⁴ year⁻¹. For a steady-state reservoir (C*^t* = C*t* − 1), *k* equals *I* divided by the inventory of carbon at year *t* (*I*/C*^t*) then Equation ([2](#page-4-0)) is derived from Equation [\(1\)](#page-4-1):

$$
F'_{C,t} = k \times F'_{atm,t} + (1 - k - \lambda) \times F'_{C,t-1}
$$
 (2)

Based on Equation ([2](#page-4-0)) a numerical model was constructed in MATLAB 2015 (MathWorks, Inc.) to calculate the optimal *k* to fit the measured $F^{14}C$ values of the soil samples. For this, atmospheric F^{14} C values ($F'_{atm,t}$) of the Northern Hemisphere were retrieved from the literature (Hammer & Levin, [2017;](#page-13-24) Hua et al., [2013](#page-13-25); ICOS, [2019](#page-13-26); Levin et al., [2013;](#page-14-20) Levin & Kromer, [2004](#page-14-21)) considering 1 year time lag from the soil sampling at each site, based on the assumption that most C inputs to the soil come from the decomposition of litter in the previous year. In addition, the atmospheric $F^{14}C$ values used were an average of measurements from summer months (May, June, July, August) for each year. The optimal *k* value to calculate the most likely turnover time was then selected as the value representing the smallest difference between the corrected *F*′ 14C values and the measured *F*^{'14}C in soil. Samples with Δ¹⁴C above 20‰ yielded two possible turnover times (Table [1\)](#page-5-0) and both values are reported, but for our interpretations, longer turnover times were used. Finally, error values (as standard deviation) were calculated using a Monte Carlo simulation (200 runs) of the corrected $F^{14}C$. The assumption of a steady-state system used to model the radiocarbon turnover times at these DIRT sites needs to be interpreted with caution as litter manipulations alter the assumed steady state of the system. However, we use this here as an exploratory tool to compare the relative change of the turnover times across treatments as previously done by Crow et al. ([2009](#page-13-8)) for two of these DIRT sites.

2.4 | **Statistical analyses**

All statistical analyses were performed in R version 3.6.0 (R Core Team, [2019](#page-15-26)) unless otherwise stated. Analytical replicates, when available, were averaged per plot (field replicate) within each site (*n* = 3). We tested the impact of the DIRT treatments within each site and the impact of both treatment and site on all variables. For each case, two models were built: a simple linear effect model (using the package "stats," base R) and a linear mixed effect model where plot was added as a random effect (using the package "lme4"; Bates et al., [2021\)](#page-12-10). Using the Akaike Information Criterion, the second model in which plot was included as a random effect, was deemed more appropriate in all cases to test the impact of treatment or treatment and site on the assessed variables. We performed a one-way and two-way analyses of variance (ANOVA) to test the impact of treatment within each site and the impact of treatment and site across the complete database using the ANOVA function from the "car" package (Fox et al., [2021](#page-13-27)) with Kenward–Roger degrees of freedom followed by multiple mean comparisons of all treatments using the Tukey test with the Bonferroni correction with the "glht" function from the "multcomp" package (Hothorn et al., [2021](#page-13-28)). Finally, as

TABLE 1 Summary of soil C content, radiocarbon data, and turnover time estimates from mineral soil (0–10 cm) from the Bousson Forest (BF), Harvard Forest (HF), and H.J. Andrews Forest (HJA) for the control (CT) and litter manipulation treatments: double litter (DL), no litter (NL), no roots (NR), and no inputs (NI)

| | BF | | | HF | | | HJA | | |
|------------------|---------------|---------------------|----------------------------|-----------------|--------------------|----------------------------|---------------|---------------------|----------------------------|
| Treatment | % $C†$ | Δ^{14} C (‰) | Turnover (years) | $\%C^{\dagger}$ | $\Delta^{14}C$ (‰) | Turnover (years) | $\%C$ | Δ^{14} C (‰) | Turnover (years) |
| CT | $6.0 + 0.6$ | -8.77 ± 2.58 | $378 + 41$ | $7.2 + 0.7$ | $38.53 + 2.94$ | $195 + 8$ | $5.0 + 0.8$ | 35.52 ± 2.85 | 199 \pm 8 ^b |
| DL | 5.7 ± 0.6 | $-14.69 + 2.68$ | $407 + 25$ | 6.9 ± 0.2 | $47.44 + 3.02$ | 172 ± 6^a | 4.9 ± 1.0 | $20.83 + 2.64$ | $245 \pm 9^{\circ}$ |
| NL | $5.0 + 0.7$ | $-22.67 + 2.70$ | $453 + 15$ | $5.8 + 0.8$ | $16.22 + 2.82$ | $265 + 10$ | $4.7 + 0.5$ | $28.62 + 2.75$ | 218 ± 9^d |
| NR | $5.7 + 0.8$ | $-14.66 + 2.55$ | 411 ± 25 | 6.8 ± 0.8 | 19.52 ± 2.95 | 254 ± 10 | 5.1 ± 0.5 | 8.00 ± 2.62 | 295 ± 11 |
| NI | $3.9 + 0.6$ | $-66.03 + 2.70$ | $763 + 48$ | 6.2 ± 0.5 | $27.36 + 2.98$ | $227 + 10$ | 4.1 ± 0.3 | $-5.08 + 2.61$ | $357 + 12$ |

Note: Turnover times were estimated using a limited steady-state model and are not intended to represent absolute turnover times but used to compare relative turnover shifts due to litter additions or exclusion relative to ambient conditions. Alternate turnover times were generated for samples with Δ^{14} C above 40‰ for HF: 3 3 \pm 1; and above 20‰ for HJA: $^{\rm b}$ z \pm 1, $^{\rm c}$ 4 \pm 1, $^{\rm d}$ s \pm 1. Errors for turnover years are standard deviation, for all other variables, error expressed as standard error.

† %C values taken from Wang et al. ([2017](#page-15-9)) for BF, Pisani et al. ([2016\)](#page-15-8) for HF and Man et al. ([2022](#page-14-12)).

a broader approach to capture differences between the treatments against the control within each site, we performed a Dunnett's test in paired comparisons using the "DunnettTest" function form the "DescTools" package (Signorell et al., [2021\)](#page-15-27).

3 | **RESULTS**

3.1 | **Long-term litter addition (double litter)**

The lack of net increases in soil C after 20 years of DIRT at any of the three forests (Table [1\)](#page-5-0) has been proposed as evidence for soil priming prompted by the doubled litter inputs and supported by observations of increased soil respiration in DL plots for all sites (Lajtha et al., [2018\)](#page-14-4). Here, we also detected higher SOM decomposition state and higher SOM resistance to decomposition in DL plots from the NMR analysis across all sites (Figure [2a,b](#page-6-0)). Higher values of these ratios in DL plots further supports the idea of a positive soil priming by doubled litter inputs as fresh litter is accelerating soil microbial decomposition of preferred C components while more complex C substrates are accumulated. Aligned with this, increases in microbial biomass and microbial-derived lipids alongside decreases in cutin-, suberin- and lignin-derived compounds and lower Δ^{14} C values and longer turnover times in DL plots would collectively support the use of older, previously stabilized SOM, as unequivocal evidence of positive soil priming. We were also expecting these trends to be more significant for bacterial-dominated sites with higher N availability (BF and HF). However, only BF was dominated by bacteria, whereas the microbial community at HF was dominated by fungi instead (Figure [3c](#page-7-0)) likely due to the acidic soil conditions at HF that generally favor fungal growth (Silva-Sánchez et al., [2019](#page-15-28)) and act as an environmental filter facilitating fungal growth rather than bacterial growth. Increases in microbial biomass (Figure [3a](#page-7-0)) and microbial-derived lipids (Figure [4c](#page-8-0)) with DL were only detected in fungal dominated sites (HF and HJA). Moreover, increases in lignin oxidation for DL

plots were observed at BF and HF (Figure [5a](#page-9-0)), whereas lignin oxida-tion decreased instead with doubled litter for HJA (Figure [5a](#page-9-0)). Cutin concentrations increased with DL only at HF (Table [S2](#page-16-3)) while suberin concentrations were unaltered by doubled litter additions across all sites (Table [S2\)](#page-16-3). Finally, lower Δ^{14} C values and longer turnover times in DL plots were detected at BF and HJA but not at HF, were higher Δ^{14} C values and shorter turnover times in DL plots were observed instead (Table [1](#page-5-0)).

3.2 | **Long-term above- and belowground litter exclusion**

As expected, litter removal treatments (NR, NL, and NI) led to increased SOM decomposition, particularly for NI plots, which had more than 25% higher decomposition ratio than CTs across all sites (Figure [2a](#page-6-0)). A higher degradation ratio (more than 25% increase) was also detected in NL and NR plots at HF (Figure [2a](#page-6-0)). As expected, the exclusion of fresh litter inputs (above-, belowground, or both) led to the accumulation of more processed and complex C compounds o and thus, higher SOM resistance to decomposition was also detected in litter removal treatments for all sites (Figure [2b](#page-6-0)). In parallel, lower Δ^{14} C values and longer calculated turnover times were observed for all litter exclusion treatments in comparison with CTs across all sites (Table [1\)](#page-5-0). The persistence of older SOM is traditionally associated with higher C substrate complexity, which is congruent with the higher SOM decomposition state and SOM resistance to decomposition (Figure [2a,b](#page-6-0)). A growing body of evidence also suggests that longer radiocarbon turnover times do not necessarily imply longer environmental persistence but instead relates to higher microbial processing of preferred C substrates (Gleixner, [2013](#page-13-29); Kleber et al., [2011](#page-13-30); Trumbore, [2009](#page-15-16)). Increases in alkyl C are likely associated not only with accumulation of complex C substrates like plant waxes, cutin, and suberin but also with newly synthesized C compounds that result from microbial processing of SOM (Baldock

FIGURE 2 (a) Soil organic matter (SOM) decomposition state ratio (alkyl/*O*-alkyl carbon), (b) SOM resistance to decomposition ratio (alkyl + aromatic and phenolic)/(*O*-alkyl + carboxyl and carbonyl carbon), and (c) associated data from solid-state 13C nuclear magnetic resonance spectroscopy for the mineral soil (0–10 cm) from the Bousson Forest (BF), Harvard Forest (HF), and H.J. Andrews Forest (HJA) for the control (CT) and litter manipulation treatments: double litter (DL), no litter (NL), no roots (NR), and no inputs (NI). Gray numbers in italics represent less than a 10% difference compared with the control, whereas numbers in black represent more than a 10% difference compared with the control and ← denotes more than 25% compared with the control. Arrows pointing upward represent increases in comparison with the control, whereas arrows pointing downward represent decreases versus the control. As decomposition increases, higher use of preferred C and the accumulation of more persistent forms of SOM also increases. Therefore, higher SOM decomposition state is related to a higher resistance to decomposition.

et al., [1992;](#page-12-11) Kögel-Knabner et al., [1991;](#page-14-22) Marín-Spiotta et al., [2008](#page-14-23)). Therefore, our observation of higher SOM decomposition and thus higher resistance to decomposition (Figure [2a,b\)](#page-6-0) along with longer radiocarbon turnover times at litter reduction plots may also indicate enhanced microbial processing of more available complex C sources via an adaptation to substrate limitations that facilitate sustain microbial metabolic activity. This also agrees with relatively higher microbial-derived lipids in litter reduction plots at both HF and HJA (Figure [4b](#page-8-0)), which is a proxy for microbially processed SOM residues (see section discussing these microbial-derived lipids in more detail).

The way each forest adapted to low C availability in litter exclusion plots differed according to the N status and the dominant microbial community at each site. We were expecting decreases in SOM compounds (i.e. lignin, cutin and suberin) as well as negative impacts on microbial biomass and would be stronger where N availability was higher. Indeed, we found that soil microbial biomass decreased in litter exclusion plots at the bacterial-dominated site (BF), whereas it increased at fungal-dominated sites (HF and HJA). Higher dependence of bacteria on labile C inputs (Butler et al., [2003](#page-12-12); Chen et al., [2016](#page-12-13); de Boer et al., [2005](#page-13-31)) was likely driving these

decreases in soil microbial biomass, particularly at NR and NI treatments (Figure [3a](#page-7-0)), where the exclusion of readily decomposable C in leaf litter evidenced by significant decreases in solvent extractable short-chain aliphatic lipids and simple sugars (Table [S1\)](#page-16-3). Microbial community composition shifted at this site toward microbial groups with higher capacity to process complex C such as Gram-positive and Actinobacteria (Figure [3c](#page-7-0)) (Deng et al., [2015](#page-13-20); Ghai et al., [2014](#page-13-21); Su et al., [2020](#page-15-22)). These compositional changes facilitated the degradation of more complex C substrates, such as lignin, which presented higher oxidation state in litter removal plots (Figure [5a](#page-9-0)), consistent with lower aromatic + phenolic C from NMR analysis (Figure [2c](#page-6-0)).

Similar to BF, litter exclusion treatments at the fungal-dominated HF site led to decreases in fast-cycling C components (Table [S1](#page-16-3)) and altered microbial community composition with increases in Grampositive and Actinobacteria abundance as previously detected at this forest as a microbial adaptation strategy to reduced preferred C substrates, consequence of soil warming treatments (Melillo et al., [2017\)](#page-14-24). Yet, soil microbial biomass tended to increase here rather than decrease (Figure [3a](#page-7-0)), and microbial community composition shifted to higher fungal abundance (higher fungi/bacteria ratio; Figure [3c](#page-7-0)) and

FIGURE 3 (a) Microbial biomass measured via phospholipid fatty acid (PLFA) profiling (μg PLFA/g dry soil), (b) microbial stress ratio and (c) microbial community composition as the ratio of Gram negative (Gram −)/ Gram positive (Gram +) and Actinobacteria on the *y*-axis and fungi-to-bacteria ratio. From mineral soil (0–10 cm) from the Bousson Forest (BF), Harvard Forest (HF), and H.J. Andrews Forest (HJA) for the control (CT) and litter manipulation treatments: double litter (DL), no litter (NL), no roots (NR), and no inputs (NI). For (c), significant differences between treatments and control (Dunett's test) on the *y*-axis shown in gray and significant differences on the *x*-axis shown in black. Significance levels: ≤.1 (•), ≤.05 (*), ≤.01 (**). For (b), different letters show significant differences *p*< .05 from Tukey's test, no letters mean no significant differences.

higher cutin degradation was observed instead of higher lignin decomposition (Figure [5a,b\)](#page-9-0). At the HF site, overall higher cutin avail-ability in comparison with the other sites (Table [S2](#page-16-3)) points to the preferential use of this OM source by the fungal-dominated community. Selective preservation of cutin can occur in substrates with high lignin contents (Angst et al., [2016](#page-12-1)) and lower relative abundance of extractable lignin phenols relative to cutin at HF in comparison with BF (mean \pm SE: extractable lignin/cutin_HF = 82.75 \pm 10.30; extractable lignin/cutin_BF = 36.66 ± 4.75) could have also facilitated the preferential use of cutin at HF, whereas enhancing its protection from degradation at BF.

Litter exclusion treatments at the coniferous low N-availability forest (HJA) significantly increased microbial biomass (Figure [3a](#page-7-0)) and decreased cutin and lignin oxidation (Figure [5b](#page-9-0)). Substrate limitations at this site did not lead to changes in microbial community composition. Rather, the increases in microbial biomass of this fungidominated community were likely allowing for a more extended nutrient scavenging area of the fungal communities; thus increasing fungal growth and enhancing substrate exploration (Paustian & Schnürer, [1987\)](#page-14-25) without the need to invest energy in degradation of available complex substrates such as cutin and lignin. Higher investment in hyphal length growth and fungal biomass production in litter removal plots at this site is also congruent with the lower efficiency

of fungi to decompose lignin from coniferous needles than that in broadleaf dominated litter (Osono, [2007](#page-14-26)).

3.3 | **Aboveground versus belowground controls of soil C accumulation**

At the studied deciduous forests (BF and HF), we observed higher relative use of preferred C substrates over persistent C in NR plots rather than in NL plots as evidenced by the higher SOM decomposition state and higher SOM resistance to decomposition in NL (Figure [2a](#page-6-0)). Aboveground litter exclusion (NL) in these forests did not lead to significant decreases in leaf-derived cutin (Table [S2\)](#page-16-3), even though cutin degradation significantly increased in NL plots in both forests compared with the control (Figure [4b](#page-8-0)). Root-derived suberin was, however, significantly higher in NR plots for both BF and HF (Table [S2](#page-16-3)), a sign pointing towards preferential preservation of rootderived C in these forests. Finally, lignin oxidation increased in NR plots rather than in NL plots at BF and HF, although this was not significant for HF (Figure [5a](#page-9-0)). Microbial biomass and the microbial stress ratio were not impacted in NL plots at either deciduous forest (Figure [3a](#page-7-0)), but significant decreases in microbial biomass accompanied by higher microbial stress ratio were detected in NR plots at BF

FIGURE 4 (a) Suberin and cutin ratio; (b) cutin degradation ratio (bars, $\omega C_{16}/C_{16}$), and cutin and suberin degradation ratio (circles, ΣMid/suberin + cutin) and (c) microbial hydrolysable-derived lipids (μg/g soil) from mineral soil (0–10 cm) from the Bousson Forest (BF), Harvard Forest (HF), and H.J. Andrews Forest (HJA) for the control (CT) and litter manipulation treatments: double litter (DL), no litter (NL), no roots (NR), and no inputs (NI). Different letters show significant differences $p < .05$ from Tukey's test, no letters mean no significant differences. For (b), significant differences across treatments and control for cutin degradation ratio ($\omega C_{16}/$ C_{16}) in black and for the cutin and suberin degradation ratio (ΣMid/ $suberin+cutin)$ in gray.

(Figure [3a](#page-7-0)). Combined, this evidence suggests that belowground C inputs have indeed a greater impact on SOM dynamics than aboveground C inputs at the studied deciduous forests. At the coniferous HJA forest, we detected higher SOM decomposition state in NL plots than NR plots (Figure [2a](#page-6-0)), accompanied by a significant decrease in short-chain aliphatic lipids (Table [S1\)](#page-16-3), lower leaf-derived cutin (Table [S2\)](#page-16-3), and lower microbial-derived lipids (Figure [4b,c](#page-8-0)). On the other hand, root exclusion (NR) at HJA did not alter the concentration of root-derived suberin (Table [S2](#page-16-3)) and only led to higher SOM resistance to decomposition (Figure [2b](#page-6-0)). This was congruent with the observed lower lignin decomposition (Figure [5a](#page-9-0)) and higher aromatic and phenolic C in NR plots at HJA (Figure [2c](#page-6-0)). The Δ^{14} C values and corresponding turnover times (Table [1](#page-5-0)) indicate a shift in soil C biogeochemical patterns. For BF and HF, litter exclusion resulted in longer turnover times compared with the control. HJA exhibited longer turnover times with NR and NI only but not with NL.

3.4 | **Microbial-derived lipids: a DIRT perspective on microbial inputs to SOM formation**

The quantity and quality of fresh plant litter inputs to soil is hypothesized to determine the substrate use of microbes and thus the quantity and chemical nature of the resulting microbial products (Microbial Efficiency-Matrix Stabilization— MEMS framework) (Cotrufo et al., [2013](#page-12-14)). Short-chain (C_{14} – C_{19}) branched alkanoic acids (*iso*, *anteiso*) (Cai et al., [2017](#page-12-5); Otto & Simpson, [2006](#page-14-16)) as well as short chain $(C_{10}-C_{18})$ β-hydroxyalkanoic acids (Allard, [2006](#page-12-15); Zelles, [1999](#page-16-4)) are lipids not commonly found in plants that have been useful as indicators of SOM decomposition and microbial inputs to the SOM pool (Cai et al., [2017](#page-12-5); Jia et al., [2019\)](#page-13-15). These lipids are generally extracted via an acidic or alkaline hydrolysis (Cai et al., [2017;](#page-12-5) Otto & Simpson, [2006\)](#page-14-16) and although the relevance and source of these microbial-derived lipids is still cryptic, they have been found to respond to warming (Jia et al., [2019](#page-13-15)) and to be related with the organic C to N ratio of SOM rather than with soil mineralogical properties (Cai et al., [2017\)](#page-12-5). The observed trends for these microbial-derived lipids in the studied forests agrees with recent proposed frameworks of stable SOM formation and its dependency on plant litter inputs chemistry and distinct microbial substrate use of contrasting microbial communities (Cotrufo et al., [2013\)](#page-12-14). We detected overall higher quantities of these microbial-derived lipids at the broadleaf deciduous forests with low C:N litter (BF and HF), whereas HJA, the coniferous forest with high C:N litter, had the lowest amount of microbial-derived compounds (Figure [4c](#page-8-0)). Higher SOM decomposition state was concurrently observed at BF and HF, whereas HJA had the lowest SOM decomposition state (Figure [2a](#page-6-0)). Thus, the quantity of these microbial-derived lipids might be closely related with the C:N ratio of the plant inputs and could potentially be a by-product of microbial-mediated SOM decomposition and/or microbial biomass turnover. Another trend we observed is that the response of these microbialderived lipids in DIRT varied depending on the dominant microbial community, regardless of the C:N ratio of the litter input. Although microbial-derived lipids in the bacteria-dominated site (BF) were generally unresponsive to doubled low C:N litter additions and reductions of above and belowground inputs (NL, NR, NI); fungal-dominated sites, particularly HJA did respond to DIRT treatments (Figure [4c](#page-8-0)), despite the different C:N ratios of the litter inputs. At the HF site, microbial-derived lipids tended to increase in DL and NR plots in comparison with the control, whereas at HJA, these lipids significantly increased in DL plots and NR but decreased in NL and NI plots (Figure [4c](#page-8-0)). Finally, the lack of responses of these lipids to DL in the bacterial-dominated deciduous forest (BF) could potentially be related to a generally lower substrate use efficiency of bacteria in comparison with that of fungi (Malik et al., [2016;](#page-14-27) Six et al., [2006](#page-15-29)) and the lower C:N litter inputs at this site that combined, led to a rapid and higher SOM decomposition state with minimal production of microbially derived compounds.

FIGURE 5 (a) Lignin oxidation ratios and (b) and lignin composition ratios from mineral soil (0–10 cm) from the Bousson Forest (BF), Harvard Forest (HF), and H.J. Andrews Forest (HJA) for the control (CT) and litter manipulation treatments: double litter (DL), no litter (NL), no roots (NR), and no inputs (NI). For (a), (Ad/Al)s or v: acid to aldehyde ratio for lignin-derived syringyl (s, white circles) and vanillyl (v, black) compounds. Inset (a) showing extractable lignin phenols as the sum of vanillyls, syringyls, and cinnamyls. Different letters show significant differences $p < .05$ from Tukey's test, no letters mean no significant differences. For (b), significant differences between treatments and control in the *y*-axis shown in gray and significant differences in the *x*-axis shown in black. Significance levels: ≤.1 (•), ≤.01 (**).

4 | **DISCUSSION**

With doubled litter inputs, sustained soil priming varied with forest N availability and soil microbial biomass and community composition. Although not all variables responded as expected to doubled litter inputs, the observed trends are consistent with the proposed long-term soil priming mechanism as the main reason for the lack of soil C accumulation at these sites with 20 years of annual DL. Instead, other lines of evidence from the collected molecular characterization of SOM from these sites can support the proposed soil priming although the driving mechanisms for this priming might differ in accordance with the N availability and the dominant microbial community at each site. For BF, doubled low C:N litter inputs may initially promote C mineralization of the plant material but as easily available C decreases, soil priming slows down and accumulation of processed, less preferred SOM is facilitated (Chao et al., [2019](#page-12-16); Chen et al., [2014;](#page-12-17) Tian et al., [2019](#page-15-30)). The positive soil priming occurring at BF was, thus, likely the product of a stoichiometric balance between the low C:N litter at the site and the biomass C:N of fast-growing *r*strategist microbes such as bacteria, which dominated the microbial community at BF (Figure [3c](#page-7-0)). Microbes, namely *r-*strategists mainly rely on easily available C (Butler et al., [2003](#page-12-12); Chen et al., [2016\)](#page-12-13) and the preferential use of more labile C components is supported by significant decreases in fast-cycling C compounds such as total sugars, short-chain aliphatic lipids and acyclic lipids as well as plant-derived long-chain aliphatic lipids for DL plots at BF (Table [S1](#page-16-3)). As this easily accessible C is depleted, lignin is used as an alternate energy source as indicated by the marginally higher lignin oxidation ratio [(Ad/Al)s, v; Figure [5](#page-9-0)a]. Also, Gram-positive and Actinobacteria

abundance in DL plots increased, facilitating the decomposition of more complex C substrates (Figure [3c](#page-7-0)). This stoichiometric decomposition theory may also explain the lack of increases in microbial biomass and microbial-derived lipids in DL plots at BF as pulses of microbial activity and growth after added doubled litter inputs might rapidly decrease and may have not been captured at the time of sampling (Wang et al., [2017](#page-15-9)) but were reported after 13 years of DIRT (Brant et al., [2006](#page-12-18)). Lack of more significant impacts on microbial biomass and microbial-derived lipids in the doubled litter treatment at BF could be also related with the rapid leaching of dissolved organic C from freshly dropped aboveground litter and its export out of the system or stabilization in mineral particles (Guggenberger & Kaiser, [2003](#page-13-32); Hensgens et al., [2020](#page-13-33); Neff & Asner, [2001\)](#page-14-28), a flux of easily available C that is particularly relevant in high-N systems (Smemo et al., [2006](#page-15-31)).

Similarly, soil priming at HF was likely governed by the stoichiometric balance between doubled low C:N litter inputs that generated an immediate priming effect (Fanin et al., [2020](#page-13-11)) that combined with fungal-dominated communities at HF (Figure [3c](#page-7-0)) led to efficient decomposition of high-quality litter inputs. Although fungi are generally considered *K*-strategists with slower growth rates and preferential use of more complex C substrates (Chao et al., [2019](#page-12-16); Chen et al., [2014;](#page-12-17) Fanin et al., [2020](#page-13-11)), fungal communities at HF challenge this notion as detrital fungal-dominated food webs at this site were found to associate with the use of easily available C substrates (Rousk & Frey, [2015\)](#page-15-14). Therefore, doubled litter inputs at this forest, increased the relative abundance of fungi over bacteria (higher fungi/bacteria ratio) (Figure [3c](#page-7-0)) and tended to increase the microbial-derived lipids (Figure [4c](#page-8-0)) without significantly altering lignin oxidation (Figure [5a](#page-9-0))

or suberin concentrations (Table [S2\)](#page-16-3). Instead, cutin concentration increased (Table [S2](#page-16-3)) and its degradation decreased (Figure [4b](#page-8-0)). HF had the highest litterfall rate across all sites (Figure [1b](#page-2-0)) and so, doubling litter inputs promoted microbial growth at this site and slightly increased SOM degradation state (Figure [2a](#page-6-0)) to a certain point where the nutrient and energy requirements of the soil microbial community were met followed by the accumulation of the litter surplus as evidenced by the higher Δ^{14} C values and shorter C turnover times at HF in DL plots. At HJA, soil priming is rather determined by active N mining by fungal-dominated microbial communities (Figure [3c](#page-7-0)) more adapted to process complex organic sources and with a typical *K*-selected strategy Thus, doubled low-quality (high C:N) litter inputs at HJA increased the relative abundance of fungi over bacteria (higher fungi/bacteria ratio) (Figure [3c](#page-7-0)), the microbial-derived lipids (Figure [4c](#page-8-0)) and more significantly increased the total microbial biomass (Figure [3a](#page-7-0)).

Across sites, litter exclusion led to the adaptation of microbial communities to preferred C-substrate limitations. As hypothesized, the negative impacts on microbial biomass were mostly detected

at the bacteria-dominated high N-availability broadleaf forest (BF) while increases in soil microbial biomass were instead detected in mid and low N-availability fungal-dominated sites (HF and HJA). Contrary to our hypothesis, SOM components (i.e., cutin, suberin, and lignin) did not decrease at all sites but instead, specific shifts in microbial community composition at each site drove targeted decreases of preferred C substrates. At BF, higher Gram-positive and Actinobacteria abundance facilitated higher lignin degradation, whereas at HF, higher fungal abundance led to higher cutin degradation. Finally, at the fungal-dominated low N-availability site (HJA) increases in microbial biomass might have served as a mechanism for extended resource scavenging without further degradation of available complex C substrates. Taken together, these results suggest that reduction of litter inputs may ultimately reduce soil C storage, not only via decreases in C inputs into the soil C pool but also via changes in the composition and abundance of soil microbial communities to either decompose complex C substrates or to explore new sites in soil to find their preferred nutrient and energy sources.

FIGURE 6 Summary of molecular composition of soil organic matter (SOM) and associated mechanisms defining the distinct responses of soil C across sites and treatments. SOM compositional data from Figures 2-5 are summarized in different groupings to better conceptualize relationships with mechanisms that control site-specific responses to long-term changes in litter quality and quantity.

We found that belowground litter controls soil C accumulation in deciduous forests, whereas aboveground litter is more influential in coniferous forests. Given that aboveground litter reductions (NL) led to more significant decreases in soil C across these DIRT forests than belowground litter reduction (NR) (Lajtha et al., [2018](#page-14-4)), we were expecting to detect higher SOM degradation, increased degradation of lignin, cutin and suberin and lower $\Delta^{14}C$ values for litter removal (NL) plots and not NR plots across all sites. However, these expectations were only partially met in specific cases with the SOM molecular composition of each site responding differently and more in agreement with the observations by Crow et al. ([2009](#page-13-8)) where the authors postulate a greater control of belowground litter in deciduous forests and aboveground litter being more influential for soil C accumulation at coniferous forest. Higher control of belowground litter in deciduous forests is supported here mainly by the lack of impacts in leaf-derived cutin for NL plots, whereas rootderived suberin was significantly higher in NR plots for both BF and HF. At the coniferous forest (HJA), higher control of aboveground litter on SOM dynamics is supported by the fact that leaf litter exclusions led to more significant changes in the SOM biogeochemistry than root exclusions. These trends were also observed by Crow et al. ([2009](#page-13-8)) and confirmed here, and are likely explained by the contrasting chemistry of above and belowground inputs from coniferous and deciduous forests. For coniferous forests, aboveground litter tends to have higher C:N ratio in comparison with root litter, with fine roots being a major input of N in these systems (Chen et al., [2002](#page-12-19); Yano et al., [2005](#page-16-2)) and having a faster decomposition than aboveground litter (Sulzman et al., [2005](#page-15-32)), whereas leaf litter decomposes slower with needle-derived compounds persisting in soil. For deciduous forests the opposite is true, given the higher C:N ratios of root litter in comparison with leaf litter, root litter decomposes slower than leaf litter in these forests (Kyaschenko et al., [2019](#page-14-29); Sulzman et al., [2005](#page-15-32)) with leaf-derived compounds thereby persisting in soil.

The measured Δ^{14} C values and the estimated turnover times for litter exclusion did not agree with the observed trends in SOM decomposition. For coniferous forests, where aboveground litter seems more consequential for soil C accumulation, we were expecting to detect lower Δ^{14} C values, and longer C turnover times in NL plots as the lack of aboveground C inputs would accelerate SOM decomposition. Instead, we observed higher Δ^{14} C values, and shorter C turnover times in NL plots at HJA (Table [1\)](#page-5-0). For deciduous forests, where roots were influencing soil C accumulation, we were expecting lower Δ^{14} C values, and longer C turnover times with belowground litter removal (NR) but instead, we observed higher Δ^{14} C values, and shorter C turnover times in NR plots at both BF and HF (Table [S1](#page-16-3)). These seemingly contradictory trends are potentially explained by the production rates of above- and belowground litter in coniferous and deciduous forests. Coniferous forests have higher relative root production compared with aboveground litterfall, whereas deciduous forests have higher litterfall relative to total and fine root production (Anderson-Teixeira et al., [2021](#page-12-20)). Thus, the continued C inputs from roots in NL plots at HJA and of leaf litter in

NR plots at BF and HF are driving the observed trends in the radiocarbon data and estimated turnover times.

5 | **CONCLUSIONS**

Our cross-site, comprehensive molecular-level comparison demonstrates that soil C responses to altered litter input quantity and quality as a proxy for shifts in ecosystem productivity with environmental change are contingent on site-specific ecosystem properties and respond uniquely to overall ecosystem N availability. Sustained annual doubled litter additions over 20 years of DIRT, either continuous (i.e., HJA) or seasonal (i.e., BF and HF), are unlikely to increase SOM in mature temperate coniferous and deciduous forests. The lack of increases in soil C suggests a long-term and sustained positive soil priming, yet the mechanisms for this priming might vary across sites, with a positive soil priming occurring at the N-rich, bacteria-dominated (BF), and fungal-dominated site (HF), characterized by a stoichiometric decomposition of the low C:N leaf litter. In contrast, the observed positive soil priming at the N-poor fungal-dominated site (HJA) may be determined by active N mining of available organic substrates. As expected, litter exclusion, either above- or belowground, led to even lower soil C across all sites with soil microbial communities adapting to the substrate limitation further degrading SOM in response to lower preferred C inputs. Finally, while root exclusion (NR) seemed more influential for soil C dynamics in both temperate broadleaf deciduous forests, litter exclusion (NL) altered soil C processes in the coniferous forest. Collectively, our results demonstrate that different mechanisms defined by inherent ecosystem properties are important determinants of altered soil C dynamics with litter manipulations. These mechanisms are conceptualized in Figure [6](#page-10-0) and highlight the important drivers of these molecular-level biogeochemical processes. It is important to emphasize that the observed changes in the molecular biogeochemistry of the SOM for both litter additions and reductions ultimately had a negative impact on the capacity of these temperate forests to sequester C. Thus, climate change, environmental alteration, or management decisions that increase, reduce or alter C input sources have the potential to reduce soil C and increase SOM decomposition via different mechanisms determined by the inherent ecosystem properties of specific forest ecosystems. This outcome agrees with recent metanalyses (Feng et al., [2022](#page-13-4); Xu et al., [2021](#page-15-7)) showing that the C sequestration potential of different ecosystems with increased litter inputs depends on factors such as the rate of the litter inputs as well as the initial soil C stocks. For these mature secondary temperate forests with already relatively high C stocks, increases in aboveground litter inputs will not lead to enhanced soil C stocks and instead, they might promote SOM decomposition, whereas the impacts of litter input reductions will further decrease soil C stocks. Our molecular-level synthesis also emphasizes the crucial role of microbial communities in SOM dynamics and how microbial community composition in combination with litter chemical

 CASTAÑEDA-GÓMEZ et al. **[|] 255**

properties and site-specific conditions can strongly determine the fate of soil C inputs. Consequently, forest management and climate change mitigation that enhances soil C sequestration and preservation of existing soil C stocks requires more refined measures based on dominant soil microbial groups and ecosystem properties to meet expected soil C enhancement goals or a point near soil C saturation in temperate forests.

ACKNOWLEDGMENTS

The authors acknowledge Dr. Maryam Tabatabaei Anaraki for assistance in retrieving microbial-derived lipids from archived base hydrolysis data sets. This research was supported by the Natural Sciences and Engineering Research Council (NSERC) of Canada via a Discovery Grant and a Tier 1 Canada Research Chair in Integrative Molecular Biogeochemistry to M.J.S. L.C.-G. thanks the Centre for Environmental Research in the Anthropocene for a postdoctoral fellowship. Allegheny College students are thanked for decades of maintenance on the BF DIRT experiment.

CONFLICT OF INTEREST

The authors have no conflict of interest to declare.

DATA AVAILABILITY STATEMENT

The data that support the findings of this study are openly available in Borealis at <https://doi.org/10.5683/SP3/L4MYGR>.

ORCID

Laura Castañeda-Gómez [https://orcid.](https://orcid.org/0000-0002-4167-4672) [org/0000-0002-4167-4672](https://orcid.org/0000-0002-4167-4672) *Kate Lajtha* <https://orcid.org/0000-0002-6430-4818> *Juan Jia* <https://orcid.org/0000-0002-9347-6221> *Xiaojuan Feng* <https://orcid.org/0000-0002-0443-0628> *Myrna J. Simpso[n](https://orcid.org/0000-0002-8084-411X)* <https://orcid.org/0000-0002-8084-411X>

REFERENCES

- Acker, S. A., Gregory, S. V., Lienkaemper, G., McKee, W. A., Swanson, F. J., & Miller, S. D. (2003). Composition, complexity, and tree mortality in riparian forests in the central Western Cascades of Oregon. *Forest Ecology and Management*, *173*, 293–308.
- Adams, M. B., & Angradi, T. R. (1996). Decomposition and nutrient dynamics of hardwood leaf litter in the Fernow Whole-Watershed Acidification Experiment. *Forest Ecology and Management*, *83*, 61– 69. [https://doi.org/10.1016/0378-1127\(95\)03695-4](https://doi.org/10.1016/0378-1127(95)03695-4)
- Allard, B. (2006). A comparative study on the chemical composition of humic acids from forest soil, agricultural soil and lignite deposit: Bound lipid, carbohydrate and amino acid distributions. *Geoderma*, *130*, 77–96. <https://doi.org/10.1016/j.geoderma.2005.01.010>
- Anderson-Teixeira, K. J., Herrmann, V., Banbury Morgan, R., Bond-Lamberty, B., Cook-Patton, S. C., Ferson, A. E., Muller-Landau, H. C., & Wang, M. M. H. (2021). Carbon cycling in mature and regrowth forests globally. *Environmental Research Letters*, *16*, 053009. <https://doi.org/10.1088/1748-9326/abed01>
- Angst, G., Heinrich, L., Kögel-Knabner, I., & Mueller, C. W. (2016). The fate of cutin and suberin of decaying leaves, needles and roots – Inferences from the initial decomposition of bound fatty acids. *Organic Geochemistry*, *95*, 81–92. [https://doi.org/10.1016/j.orgge](https://doi.org/10.1016/j.orggeochem.2016.02.006) [ochem.2016.02.006](https://doi.org/10.1016/j.orggeochem.2016.02.006)
- Bahri, H., Dignac, M.-F., Rumpel, C., Rasse, D. P., Chenu, C., & Mariotti, A. (2006). Lignin turnover kinetics in an agricultural soil is monomer specific. *Soil Biology and Biochemistry*, *38*, 1977–1988. [https://doi.](https://doi.org/10.1016/j.soilbio.2006.01.003) [org/10.1016/j.soilbio.2006.01.003](https://doi.org/10.1016/j.soilbio.2006.01.003)
- Baldock, J. A., Oades, J. M., Waters, A. G., Peng, X., Vassallo, A. M., & Wilson, M. A. (1992). Aspects of the chemical structure of soil organic materials as revealed by solid-state ¹³C NMR spectroscopy. *Biogeochemistry*, *16*, 1–42.<https://doi.org/10.1007/BF00024251>
- Baldock, J. A., & Preston, C. M. (1995). Chemistry of carbon decomposition processes in forests as revealed by solid-state carbon-13 nuclear magnetic resonance. In W. W. McFee & J. M. Kelly (Eds.), *Carbon forms and functions in forest soils* (pp. 89–117). Soil Science Society of America.
- Bates, D., Maechler, M., Bolker, B., Walker, S., Christense, R. H. B., Singman, H., Dai, B., Scheip, F., Grothendieck, G., Green, P., Fox, J., Baue, A., & Krivitsky, P. N. (2021). *lme4: Linear mixed-effects models using "Eigen" and S4*.
- Bligh, E. G., & Dyer, W. J. (1959). A rapid method of total lipid extraction and purification. *Canadian Journal of Biochemistry and Physiology*, *37*, 911–917.
- Bossio, D. A., & Scow, K. M. (1998). Impacts of carbon and flooding on soil microbial communities: Phospholipid fatty acid profiles and substrate utilization patterns. *Microbial Ecology*, *35*, 265–278. <https://doi.org/10.1007/s002489900082>
- Bowden, R. D., Deem, L., Plante, A. F., Peltre, C., Nadelhoffer, K., & Lajtha, K. (2014). Litter input controls on soil carbon in a temperate deciduous forest. *Soil Science Society of America Journal*, *78*, S66– S75.<https://doi.org/10.2136/sssaj2013.09.0413nafsc>
- Brant, J. B., Myrold, D. D., & Sulzman, E. W. (2006). Root controls on soil microbial community structure in forest soils. *Oecologia*, *148*, 650– 659. <https://doi.org/10.1007/s00442-006-0402-7>
- Butler, J. L., Williams, M. A., Bottomley, P. J., & Myrold, D. D. (2003). Microbial community dynamics associated with rhizosphere carbon flow. *Applied and Environmental Microbiology*, *69*, 6793–6800. <https://doi.org/10.1128/AEM.69.11.6793-6800.2003>
- Cai, Y., Tang, Z., Xiong, G., Xie, Z., Liu, Z., & Feng, X. (2017). Different composition and distribution patterns of mineral-protected versus hydrolyzable lipids in shrubland soils. *Journal of Geophysical Research: Biogeosciences*, *122*, 2206–2218. [https://doi.org/10.1002/2017J](https://doi.org/10.1002/2017JG003759) [G003759](https://doi.org/10.1002/2017JG003759)
- Chao, L., Liu, Y., Freschet, G. T., Zhang, W., Yu, X., Zheng, W., Guan, X., Yang, Q., Chen, L., Dijkstra, F. A., & Wang, S. (2019). Litter carbon and nutrient chemistry control the magnitude of soil priming effect. *Functional Ecology*, *33*, 876–888. [https://doi.](https://doi.org/10.1111/1365-2435.13278) [org/10.1111/1365-2435.13278](https://doi.org/10.1111/1365-2435.13278)
- Chen, H., Harmon, M. E., Sexton, J., & Fasth, B. (2002). Fine-root decomposition and N dynamics in coniferous forests of the Pacific Northwest, U.S.A. *Canadian Journal of Forest Research*, *32*, 320– 331. <https://doi.org/10.1139/x01-202>
- Chen, R., Senbayram, M., Blagodatsky, S., Myachina, O., Dittert, K., Lin, X., Blagodatskaya, E., & Kuzyakov, Y. (2014). Soil C and N availability determine the priming effect: Microbial N mining and stoichiometric decomposition theories. *Global Change Biology*, *20*, 2356–2367. <https://doi.org/10.1111/gcb.12475>
- Chen, Y., Chen, G., Robinson, D., Yang, Z., Guo, J., Xie, J., Fu, S., Zhou, L., & Yang, Y. (2016). Large amounts of easily decomposable carbon stored in subtropical forest subsoil are associated with r-strategydominated soil microbes. *Soil Biology and Biochemistry*, *95*, 233– 242. <https://doi.org/10.1016/j.soilbio.2016.01.004>
- Corrigan, C. K. (2008). *The decomposition of leaf litter in litter traps: Implications on forest biogeochemical cycling* (Master's Thesis, UWSpace).
- Cotrufo, M. F., Wallenstein, M. D., Boot, C. M., Denef, K., & Paul, E. (2013). The Microbial Efficiency-Matrix Stabilization (MEMS) framework integrates plant litter decomposition with soil organic matter stabilization: Do labile plant inputs form stable soil

256 | WILEY- Global Change Biology | CASTAÑEDA-GÓMEZ ET AL.

organic matter? *Global Change Biology*, *19*, 988–995. [https://doi.](https://doi.org/10.1111/gcb.12113) [org/10.1111/gcb.12113](https://doi.org/10.1111/gcb.12113)

- Crann, C. A., Murseli, S., St-Jean, G., Zhao, X., Clark, I. D., & Kieser, W. E. (2017). First status report on radiocarbon sample preparation techniques at the A.E. Lalonde AMS Laboratory (Ottawa, Canada). *Radiocarbon*, *59*, 695–704.<https://doi.org/10.1017/RDC.2016.55>
- Crow, S. E., Lajtha, K., Filley, T. R., Swanston, C., Bowden, R. D., & Caldwell, B. A. (2009). Sources of plant-derived carbon and stability of organic matter in soil: Implications for global change. *Global Change Biology*, *15*, 2003–2019. [https://doi.](https://doi.org/10.1111/j.1365-2486.2009.01850.x) [org/10.1111/j.1365-2486.2009.01850.x](https://doi.org/10.1111/j.1365-2486.2009.01850.x)
- de Boer, W., Folman, L. B., Summerbell, R. C., & Boddy, L. (2005). Living in a fungal world: Impact of fungi on soil bacterial niche development. *FEMS Microbiology Reviews*, *29*, 795–811. [https://doi.org/10.1016/j.](https://doi.org/10.1016/j.femsre.2004.11.005) [femsre.2004.11.005](https://doi.org/10.1016/j.femsre.2004.11.005)
- Deng, J., Gu, Y., Zhang, J., Xue, K., Qin, Y., Yuan, M., Yin, H., He, Z., Wu, L., Schuur, E. A. G., Tiedje, J. M., & Zhou, J. (2015). Shifts of tundra bacterial and archaeal communities along a permafrost thaw gradient in Alaska. *Molecular Ecology*, *24*, 222–234. [https://doi.org/10.1111/](https://doi.org/10.1111/mec.13015) [mec.13015](https://doi.org/10.1111/mec.13015)
- Edmonds, R. L. (1980). Litter decomposition and nutrient release in Douglas-fir, red alder, western hemlock, and Pacific silver fir ecosystems in western Washington. *Canadian Journal of Forest Research*, *10*, 327–337.<https://doi.org/10.1139/x80-056>
- Edmonds, R. L., & Thomas, T. B. (1995). Decomposition and nutrient release from green needles of western hemlock and Pacific silver fir in an old-growth temperate rain forest9 Olympic National Park, Washington. *Canadian Journal of Forest Research*, *25*, 1049–1057. <https://doi.org/10.1139/x95-115>
- Fanin, N., Alavoine, G., & Bertrand, I. (2020). Temporal dynamics of litter quality, soil properties and microbial strategies as main drivers of the priming effect. *Geoderma*, *377*, 114576. [https://doi.](https://doi.org/10.1016/j.geoderma.2020.114576) [org/10.1016/j.geoderma.2020.114576](https://doi.org/10.1016/j.geoderma.2020.114576)
- Fanin, N., Kardol, P., Farrell, M., Nilsson, M. C., Gundale, M. J., & Wardle, D. A. (2019). The ratio of Gram-positive to Gram-negative bacterial PLFA markers as an indicator of carbon availability in organic soils. *Soil Biology and Biochemistry*, *128*, 111–114. [https://doi.](https://doi.org/10.1016/j.soilbio.2018.10.010) [org/10.1016/j.soilbio.2018.10.010](https://doi.org/10.1016/j.soilbio.2018.10.010)
- FAO, & UNEP. (2020). *The state of the world's forests 2020: Forests, biodiversity and people*. FAO and UNEP.
- Feng, J., He, K., Zhang, Q., Han, M., & Zhu, B. (2022). Changes in plant inputs alter soil carbon and microbial communities in forest ecosystems. *Global Change Biology*, *28*, 3426–3440. [https://doi.](https://doi.org/10.1111/gcb.16107) [org/10.1111/gcb.16107](https://doi.org/10.1111/gcb.16107)
- Feng, X., Simpson, A. J., Schlesinger, W. H., & Simpson, M. J. (2010). Altered microbial community structure and organic matter composition under elevated $CO₂$ and N fertilization in the duke forest. *Global Change Biology*, *16*, 2104–2116. [https://doi.](https://doi.org/10.1111/j.1365-2486.2009.02080.x) [org/10.1111/j.1365-2486.2009.02080.x](https://doi.org/10.1111/j.1365-2486.2009.02080.x)
- Feng, X., Simpson, A. J., Wilson, K. P., Dudley Williams, D., & Simpson, M. J. (2008). Increased cuticular carbon sequestration and lignin oxidation in response to soil warming. *Nature Geoscience*, *1*, 836–839. <https://doi.org/10.1038/ngeo361>
- Feng, X., & Simpson, M. J. (2011). Molecular-level methods for monitoring soil organic matter responses to global climate change. *Journal of Environmental Monitoring*, *13*, 1246–1254. [https://doi.](https://doi.org/10.1039/C0EM00752H) [org/10.1039/C0EM00752H](https://doi.org/10.1039/C0EM00752H)
- Fox, J., Weisberg, S., Price, B., Adler, D., Bates, D., Baud-Bovy, G., Bolker, B., Ellison, S., Firth, D., Friendly, M., Gorjanc, G., Graves, S., Heiberger, R., Krivitsky, P., Laboissiere, R., Maechler, M., Monette, G., Murdoch, D., Nilsson, H., … R-Core. (2021). *car: Companion to applied regression*.
- Frostegård, Å., & Bååth, E. (1996). The use of phospholipid fatty acid analysis to estimate bacterial and fungal biomass in soil. *Biology and Fertility of Soils*, *22*, 59–65.
- Frostegård, Å., Tunlid, A., & Bååth, E. (2011). Use and misuse of PLFA measurements in soils. *Soil Biology and Biochemistry*, *43*, 1621– 1625. <https://doi.org/10.1016/j.soilbio.2010.11.021>
- Georgiou, K., Jackson, R. B., Vindušková, O., Abramoff, R. Z., Ahlström, A., Feng, W., Harden, J. W., Pellegrini, A. F. A., Polley, H. W., Soong, J. L., Riley, W. J., & Torn, M. S. (2022). Global stocks and capacity of mineral-associated soil organic carbon. *Nature Communications*, *13*, 3797. <https://doi.org/10.1038/s41467-022-31540-9>
- Ghai, R., Mizuno, C. M., Picazo, A., Camacho, A., & Rodriguez-Valera, F. (2014). Key roles for freshwater Actinobacteria revealed by deep metagenomic sequencing. *Molecular Ecology*, *23*, 6073–6090. <https://doi.org/10.1111/mec.12985>
- Gifford, R. M., Barrett, D. J., & Lutze, J. L. (2000). The effects of elevated [CO₂] on the C:N and C:P mass ratios of plant tissues. *Plant and Soil*, *224*, 1–14. <https://doi.org/10.1023/A:1004790612630>
- Gleixner, G. (2013). Soil organic matter dynamics: A biological perspective derived from the use of compound-specific isotopes studies. *Ecological Research*, *28*, 683–695. [https://doi.org/10.1007/s1128](https://doi.org/10.1007/s11284-012-1022-9) [4-012-1022-9](https://doi.org/10.1007/s11284-012-1022-9)
- Goñi, M. A., & Hedges, J. I. (1990). The diagenetic behavior of cutin acids in buried conifer needles and sediments from a coastal marine environment. *Geochimica et Cosmochimica Acta*, *54*, 3083–3093. [https://doi.org/10.1016/0016-7037\(90\)90124-4](https://doi.org/10.1016/0016-7037(90)90124-4)
- Guggenberger, G., & Kaiser, K. (2003). Dissolved organic matter in soil: Challenging the paradigm of sorptive preservation. *Geoderma*, *113*, 293–310. [https://doi.org/10.1016/S0016-7061\(02\)00366-X](https://doi.org/10.1016/S0016-7061(02)00366-X)
- Hammer, S., & Levin, I. (2017). Monthly mean atmospheric $D^{14}CO$, at *Jungfraujoch and Schauinsland from 1986 to 2016*. heiDATA, V2. <https://doi.org/10.11588/data/10100>
- Hedges, J. I., & Mann, D. C. (1979). The characterization of plant tissues by their lignin oxidation products. *Geochimica et Cosmochimica Acta*, *43*, 1803–1807. [https://doi.org/10.1016/0016-7037\(79\)90028-0](https://doi.org/10.1016/0016-7037(79)90028-0)
- Hensgens, G., Laudon, H., Peichl, M., Gil, I. A., Zhou, Q., & Berggren, M. (2020). The role of the understory in litter DOC and nutrient leaching in boreal forests. *Biogeochemistry*, *149*, 87–103. [https://doi.](https://doi.org/10.1007/s10533-020-00668-5) [org/10.1007/s10533-020-00668-5](https://doi.org/10.1007/s10533-020-00668-5)
- Hicks, L. C., Lajtha, K., & Rousk, J. (2021). Nutrient limitation may induce microbial mining for resources from persistent soil organic matter. *Ecology*, *102*, e03328.<https://doi.org/10.1002/ecy.3328>
- Hothorn, T., Bretz, F., Westfall, P., Heiberger, R. M., Schuetzenmeister, A., & Scheibe, S. (2021). *multcomp: Simultaneous inference in general parametric models*.
- Hua, Q., Barbetti, M., & Rakowski, A. Z. (2013). Atmospheric radiocarbon for the period 1950–2010. *Radiocarbon*, *55*, 2059–2072. [https://](https://doi.org/10.2458/azu_js_rc.v55i2.16177) doi.org/10.2458/azu_js_rc.v55i2.16177
- ICOS. (2019). *ICOS ATC/CAL 14C Release, Jungfraujoch (10.0 m)*. [https://](https://hdl.handle.net/11676/0yDrKllpdGVPtMsFIMSUEm1N) hdl.handle.net/11676/0yDrKllpdGVPtMsFIMSUEm1N
- Janzen, H. H. (2004). Carbon cycling in earth systems—a soil science perspective. *Agriculture, Ecosystems & Environment*, *104*, 399–417. <https://doi.org/10.1016/j.agee.2004.01.040>
- Jia, J., Cao, Z., Liu, C., Zhang, Z., Lin, L., Wang, Y., Haghipour, N., Wacker, L., Bao, H., Dittmar, T., Simpson, M. J., Yang, H., Crowther, T. W., Eglinton, T. I., He, J. S., & Feng, X. (2019). Climate warming alters subsoil but not topsoil carbon dynamics in alpine grassland. *Global Change Biology*, *25*, 4383–4393.<https://doi.org/10.1111/gcb.14823>
- King, J. S., Pregitzer, K. S., Zak, D. R., Kubiske, M. E., & Holmes, W. E. (2001). Correlation of foliage and litter chemistry of sugar maple, Acer saccharum, as affected by elevated $CO₂$ and varying N availability, and effects on decomposition. *Oikos*, *94*, 403–416. [https://](https://doi.org/10.1034/j.1600-0706.2001.940303.x) doi.org/10.1034/j.1600-0706.2001.940303.x
- Kirschbaum, M. U. F. (2000). Will changes in soil organic carbon act as a positive or negative feedback on global warming? *Biogeochemistry*, *48*, 21–51.
- Kleber, M., Nico, P. S., Plante, A., Filley, T., Kramer, M., Swanston, C., & Sollins, P. (2011). Old and stable soil organic matter is not

 CASTAÑEDA-GÓMEZ et al. **[|] 257**

necessarily chemically recalcitrant: Implications for modeling concepts and temperature sensitivity. *Global Change Biology*, *17*, 1097– 1107. <https://doi.org/10.1111/j.1365-2486.2010.02278.x>

- Kögel-Knabner, I. (2017). The macromolecular organic composition of plant and microbial residues as inputs to soil organic matter: Fourteen years on. *Soil Biology and Biochemistry*, *105*, A3–A8. <https://doi.org/10.1016/j.soilbio.2016.08.011>
- Kögel-Knabner, I., Zech, W., & Hatcher, P. G. (1991). Chemical structural studies of forest soil humic acids: Aromatic carbon fraction. *Soil Science Society of America Journal*, *55*, 241–247. [https://doi.](https://doi.org/10.2136/sssaj1991.03615995005500010041x) [org/10.2136/sssaj1991.03615995005500010041x](https://doi.org/10.2136/sssaj1991.03615995005500010041x)
- Kuzyakov, Y., Friedel, J. K., & Stahr, K. (2000). Review of mechanisms and quantification of priming effects. *Soil Biology and Biochemistry*, *32*, 1485–1498.
- Kyaschenko, J., Ovaskainen, O., Ekblad, A., Hagenbo, A., Karltun, E., Clemmensen, K. E., & Lindahl, B. D. (2019). Soil fertility in boreal forest relates to root-driven nitrogen retention and carbon sequestration in the mor layer. *New Phytologist*, *221*, 1492–1502. [https://](https://doi.org/10.1111/nph.15454) doi.org/10.1111/nph.15454
- Lajtha, K., Bowden, R. D., Crow, S., Fekete, I., Kotroczó, Z., Plante, A., Simpson, M. J., & Nadelhoffer, K. J. (2018). The detrital input and removal treatment (DIRT) network: Insights into soil carbon stabilization. *Science of the Total Environment*, *640–641*, 1112–1120. <https://doi.org/10.1016/j.scitotenv.2018.05.388>
- Lajtha, K., Bowden, R. D., & Nadelhoffer, K. (2014). Litter and root manipulations provide insights into soil organic matter dynamics and stability. *Soil Science Society of America Journal*, *78*, S261–S269. [https://](https://doi.org/10.2136/sssaj2013.08.0370nafsc) doi.org/10.2136/sssaj2013.08.0370nafsc
- Lal, R. (2005). Forest soils and carbon sequestration. *Forest Ecology and Management*, *220*, 242–258. [https://doi.org/10.1016/j.](https://doi.org/10.1016/j.foreco.2005.08.015) [foreco.2005.08.015](https://doi.org/10.1016/j.foreco.2005.08.015)
- Leite, M. F. A., Pan, Y., Bloem, J., Berge, H., & Kuramae, E. E. (2017). Organic nitrogen rearranges both structure and activity of the soilborne microbial seedbank. *Scientific Reports*, *7*, 42634. [https://doi.](https://doi.org/10.1038/srep42634) [org/10.1038/srep42634](https://doi.org/10.1038/srep42634)
- Levin, I., & Kromer, B. (2004). The tropospheric $^{14}CO_2$ level in midlatitudes of the northern hemisphere (1959–2003). *Radiocarbon*, *46*, 1261–1272.<https://doi.org/10.1017/S0033822200033130>
- Levin, I., Kromer, B., & Hammer, S. (2013). Atmospheric Δ^{14} CO₂ trend in Western European background air from 2000 to 2012. *Tellus B: Chemical and Physical Meteorology*, *65*, 20092. [https://doi.](https://doi.org/10.3402/tellusb.v65i0.20092) [org/10.3402/tellusb.v65i0.20092](https://doi.org/10.3402/tellusb.v65i0.20092)
- Lützow, M. V., Kögel-Knabner, I., Ekschmitt, K., Matzner, E., Guggenberger, G., Marschner, B., & Flessa, H. (2006). Stabilization of organic matter in temperate soils: Mechanisms and their relevance under different soil conditions – A review. *European Journal of Soil Science*, *57*, 426–445. [https://doi.](https://doi.org/10.1111/j.1365-2389.2006.00809.x) [org/10.1111/j.1365-2389.2006.00809.x](https://doi.org/10.1111/j.1365-2389.2006.00809.x)
- Malik, A. A., Chowdhury, S., Schlager, V., Oliver, A., Puissant, J., Vazquez, P. G., Jehmlich, N., von Bergen, M., Griffiths, R. I., & Gleixner, G. (2016). Soil fungal:bacterial ratios are linked to altered carbon cycling. *Frontiers in Microbiology*, *7*, 1247. [https://doi.org/10.3389/](https://doi.org/10.3389/fmicb.2016.01247) [fmicb.2016.01247](https://doi.org/10.3389/fmicb.2016.01247)
- Man, M., Pierson, D., Chiu, R., Tabatabaei Anaraki, M., van den Enden, L., Ye, R. X., Lajtha, K., & Simpson, M. J. (2022). Twenty years of litter manipulation reveals that above-ground litter quantity and quality controls soil organic matter molecular composition. *Biogeochemistry*, *159*, 393–411.<https://doi.org/10.1007/s10533-022-00934-8>
- Marín-Spiotta, E., Swanston, C. W., Torn, M. S., Silver, W. L., & Burton, S. D. (2008). Chemical and mineral control of soil carbon turnover in abandoned tropical pastures. *Geoderma*, *143*, 49–62. [https://doi.](https://doi.org/10.1016/j.geoderma.2007.10.001) [org/10.1016/j.geoderma.2007.10.001](https://doi.org/10.1016/j.geoderma.2007.10.001)
- Matson, P. A., McDowell, W. H., Townsend, A. R., & Vitousek, P. M. (1999). The globalization of N deposition: Ecosystem consequences in tropical environments. *Biogeochemistry*, *46*, 67–83. [https://doi.](https://doi.org/10.1023/A:1006152112852) [org/10.1023/A:1006152112852](https://doi.org/10.1023/A:1006152112852)
- Mayer, M., Prescott, C. E., Abaker, W. E. A., Augusto, L., Cécillon, L., Ferreira, G. W. D., James, J., Jandl, R., Katzensteiner, K., Laclau, J. P., Laganière, J., Nouvellon, Y., Paré, D., Stanturf, J. A., Vanguelova, E. I., & Vesterdal, L. (2020). Tamm review: Influence of forest management activities on soil organic carbon stocks: A knowledge synthesis. *Forest Ecology and Management*, *466*, 118127. [https://doi.](https://doi.org/10.1016/j.foreco.2020.118127) [org/10.1016/j.foreco.2020.118127](https://doi.org/10.1016/j.foreco.2020.118127)
- Melillo, J. M., Frey, S. D., DeAngelis, K. M., Werner, W. J., Bernard, M. J., Bowles, F. P., Pold, G., Knorr, M. A., & Grandy, A. S. (2017). Longterm pattern and magnitude of soil carbon feedback to the climate system in a warming world. *Science*, *358*, 101–105. [https://doi.](https://doi.org/10.1126/science.aan2874) [org/10.1126/science.aan2874](https://doi.org/10.1126/science.aan2874)
- Nadelhoffer, K. J., Boone, R. D., Bowden, R. D., Canary, J. K., Micks, P., Ricca, A., McDowell, W., & Aitkenhead, J. (2004). The DIRT experiment: Litter and root influences on forest soil organic matter stocks and function. In D. Foster & J. Aber (Eds.), *Forests in time: The environmental consequences of 1000 years of change in New England* (pp. 300–315). Yale University Press.
- Neff, J. C., & Asner, G. P. (2001). Dissolved organic carbon in terrestrial ecosystems: Synthesis and a model. *Ecosystems*, *4*, 29–48. [https://](https://doi.org/10.1007/s100210000058) doi.org/10.1007/s100210000058
- Osono, T. (2007). Ecology of ligninolytic fungi associated with leaf litter decomposition. *Ecological Research*, *22*, 955–974. [https://doi.](https://doi.org/10.1007/s11284-007-0390-z) [org/10.1007/s11284-007-0390-z](https://doi.org/10.1007/s11284-007-0390-z)
- Ostertag, R., Marín-Spiotta, E., Silver, W. L., & Schulten, J. (2008). Litterfall and decomposition in relation to soil carbon pools along a secondary forest chronosequence in Puerto Rico. *Ecosystems*, *11*, 701–714. <https://doi.org/10.1007/s10021-008-9152-1>
- Otto, A., Shunthirasingham, C., & Simpson, M. J. (2005). A comparison of plant and microbial biomarkers in grassland soils from the Prairie Ecozone of Canada. *Organic Geochemistry*, *36*, 425–448. [https://](https://doi.org/10.1016/j.orggeochem.2004.09.008) doi.org/10.1016/j.orggeochem.2004.09.008
- Otto, A., & Simpson, M. J. (2005). Degradation and preservation of vascular plant-derived biomarkers in grassland and forest soils from western Canada. *Biogeochemistry*, *74*, 377–409. [https://doi.](https://doi.org/10.1007/s10533-004-5834-8) [org/10.1007/s10533-004-5834-8](https://doi.org/10.1007/s10533-004-5834-8)
- Otto, A., & Simpson, M. J. (2006). Sources and composition of hydrolysable aliphatic lipids and phenols in soils from western Canada. *Organic Geochemistry*, *37*, 385–407. [https://doi.org/10.1016/j.](https://doi.org/10.1016/j.orggeochem.2005.12.011) [orggeochem.2005.12.011](https://doi.org/10.1016/j.orggeochem.2005.12.011)
- Otto, A., & Simpson, M. J. (2007). Analysis of soil organic matter biomarkers by sequential chemical degradation and gas chromatography – Mass spectrometry. *Journal of Separation Science*, *30*(2), 272– 282. <https://doi.org/10.1002/jssc.200600243>
- Parsons, W. F. J., Lindroth, R. L., & Bockheim, J. G. (2004). Decomposition of *Betula papyrifera* leaf litter under the independent and interactive effects of elevated CO₂ and O₃. Global Change Biology, 10, 1666–1677. <https://doi.org/10.1111/j.1365-2486.2004.00851.x>
- Paustian, K., & Schnürer, J. (1987). Fungal growth response to carbon and nitrogen limitation: A theoretical model. *Soil Biology and Biochemistry*, *19*, 613–620. [https://doi.org/10.1016/0038-0717\(87\)90107-6](https://doi.org/10.1016/0038-0717(87)90107-6)
- Pendall, E., Bridgham, S., Hanson, P. J., Hungate, B., Kicklighter, D. W., Johnson, D. W., Law, B. E., Luo, Y., Megonigal, J. P., Olsrud, M., Ryan, M. G., & Wan, S. (2004). Below-ground process responses to elevated $CO₂$ and temperature: A discussion of observations, measurement methods, and models. *New Phytologist*, *162*, 311–322. <https://doi.org/10.1111/j.1469-8137.2004.01053.x>
- Perakis, S. S., Matkins, J. J., & Hibbs, D. E. (2012). Interactions of tissue and fertilizer nitrogen on decomposition dynamics of ligninrich conifer litter. *Ecosphere*, *3*, art54. [https://doi.org/10.1890/](https://doi.org/10.1890/ES11-00340.1) [ES11-00340.1](https://doi.org/10.1890/ES11-00340.1)
- Pierson, D., Peter-Contesse, H., Bowden, R. D., Nadelhoffer, K., Kayhani, K., Evans, L., & Lajtha, K. (2021). Competing processes drive the resistance of soil carbon to alterations in organic inputs. *Frontiers in Environmental Science*, *9*, 527803. [https://doi.org/10.3389/](https://doi.org/10.3389/fenvs.2021.527803) [fenvs.2021.527803](https://doi.org/10.3389/fenvs.2021.527803)

258 \blacksquare \blacksquare

- Pisani, O., Lin, L. H., Lun, O. O. Y., Lajtha, K., Nadelhoffer, K. J., Simpson, A. J., & Simpson, M. J. (2016). Long-term doubling of litter inputs accelerates soil organic matter degradation and reduces soil carbon stocks. *Biogeochemistry*, *127*, 1–14. [https://doi.org/10.1007/s1053](https://doi.org/10.1007/s10533-015-0171-7) [3-015-0171-7](https://doi.org/10.1007/s10533-015-0171-7)
- Quinn Thomas, R., Canham, C. D., Weathers, K. C., & Goodale, C. L. (2010). Increased tree carbon storage in response to nitrogen deposition in the US. *Nature Geoscience*, *3*, 13–17. [https://doi.](https://doi.org/10.1038/ngeo721) [org/10.1038/ngeo721](https://doi.org/10.1038/ngeo721)
- R Core Team. (2019). *R: The R project for statistical computing*. [https://](https://www.r-project.org/) www.r-project.org/
- Rasse, D. P., Rumpel, C., & Dignac, M.-F. (2005). Is soil carbon mostly root carbon? Mechanisms for a specific stabilisation. *Plant and Soil*, *269*, 341–356.<https://doi.org/10.1007/s11104-004-0907-y>
- Reichstein, M. (2007). Impacts of climate change on forest soil carbon: Principles, factors, models, uncertainties. In P. H. Freer-Smith, M. S. J. Broadmeadow, & J. M. Lynch (Eds.), *Forestry and climate change* (pp. 127–135). CABI.
- Reimer, P. J., Bard, E., Bayliss, A., Beck, J. W., Blackwell, P. G., Ramsey, C. B., Buck, C. E., Cheng, H., Edwards, R. L., Friedrich, M., Grootes, P. M., Guilderson, T. P., Haflidason, H., Hajdas, I., Hatté, C., Heaton, T. J., Hoffmann, D. L., Hogg, A. G., Hughen, K. A., … van der Plicht, J. (2013). IntCal13 and Marine13 radiocarbon age calibration curves 0–50,000 years cal BP. *Radiocarbon*, *55*, 1869–1887. [https://doi.](https://doi.org/10.2458/azu_js_rc.55.16947) [org/10.2458/azu_js_rc.55.16947](https://doi.org/10.2458/azu_js_rc.55.16947)
- Rousk, J., Brookes, P. C., & Bååth, E. (2011). Fungal and bacterial growth responses to N fertilization and pH in the 150-year 'Park Grass' UK grassland experiment. *FEMS Microbiology Ecology*, *76*, 89–99. <https://doi.org/10.1111/j.1574-6941.2010.01032.x>
- Rousk, J., & Frey, S. D. (2015). Revisiting the hypothesis that fungal-tobacterial dominance characterizes turnover of soil organic matter and nutrients. *Ecological Monographs*, *85*, 457–472.
- Schuur, E. A. G., Druffel, E., & Trumbore, S. E. (Eds.). (2016). *Radiocarbon and climate change*. Springer International Publishing.
- Schwaner, G. W., & Kelly, C. N. (2019). American chestnut soil carbon and nitrogen dynamics: Implications for ecosystem response following restoration. *Pedobiologia*, *75*, 24–33. [https://doi.org/10.1016/j.](https://doi.org/10.1016/j.pedobi.2019.05.003) [pedobi.2019.05.003](https://doi.org/10.1016/j.pedobi.2019.05.003)
- Signorell, A., Aho, K., Alfons, A., Anderegg, N., Aragon, T., Arachchige, C., Arppe, A., Baddeley, A., Barton, K., Bolker, B., Borchers, H. W., Caeiro, F., Champely, S., Chessel, D., Chhay, L., Cooper, N., Cummins, C., Dewey, M., Doran, H. C., … Zeileis, A. (2021). DescTools: Tools for Descriptive Statistics (0.99.42). [https://CRAN.R-project.org/](https://CRAN.R-project.org/package=DescTools) package=[DescTools](https://CRAN.R-project.org/package=DescTools)
- Silva-Sánchez, A., Soares, M., & Rousk, J. (2019). Testing the dependence of microbial growth and carbon use efficiency on nitrogen availability, pH, and organic matter quality. *Soil Biology and Biochemistry*, *134*, 25– 35. <https://doi.org/10.1016/j.soilbio.2019.03.008>
- Simpson, M. J., & Simpson, A. J. (2012). The chemical ecology of soil organic matter molecular constituents. *Journal of Chemical Ecology*, *38*, 768–784. <https://doi.org/10.1007/s10886-012-0122-x>
- Six, J., Frey, S. D., Thiet, R. K., & Batten, K. M. (2006). Bacterial and fungal contributions to carbon sequestration in agroecosystems. *Soil Science Society of America Journal*, *70*, 555–569. [https://doi.](https://doi.org/10.2136/sssaj2004.0347) [org/10.2136/sssaj2004.0347](https://doi.org/10.2136/sssaj2004.0347)
- Smemo, K. A., Zak, D. R., & Pregitzer, K. S. (2006). Chronic experimental NO^{3−} deposition reduces the retention of leaf litter DOC in a northern hardwood forest soil. *Soil Biology and Biochemistry*, *38*, 1340– 1347.<https://doi.org/10.1016/j.soilbio.2005.09.029>
- Smith, P., Fang, C., Dawson, J. J. C., & Moncrieff, J. B. (2008). Impact of global warming on soil organic carbon. In *Advances in agronomy* (pp. 1–43). Academic Press. [https://doi.org/10.1016/S0065-2113\(07\)00001-6](https://doi.org/10.1016/S0065-2113(07)00001-6)
- Song, J., Wan, S., Piao, S., Knapp, A. K., Classen, A. T., Vicca, S., Ciais, P., Hovenden, M. J., Leuzinger, S., Beier, C., Kardol, P., Xia, J., Liu, Q., Ru, J., Zhou, Z., Luo, Y., Guo, D., Adam Langley, J., Zscheischler, J., … Zheng, M. (2019). A meta-analysis of 1,119 manipulative

experiments on terrestrial carbon-cycling responses to global change. *Nature Ecology & Evolution*, *3*, 1309–1320. [https://doi.](https://doi.org/10.1038/s41559-019-0958-3) [org/10.1038/s41559-019-0958-3](https://doi.org/10.1038/s41559-019-0958-3)

- Spielvogel, S., Prietzel, J., Leide, J., Riedel, M., Zemke, J., & Kögel-Knabner, I. (2014). Distribution of cutin and suberin biomarkers under forest trees with different root systems. *Plant and Soil*, *381*, 95–110. <https://doi.org/10.1007/s11104-014-2103-z>
- Stanek, M., Piechnik, Ł., & Stefanowicz, A. M. (2020). Invasive red oak (*Quercus rubra* L.) modifies soil physicochemical properties and forest understory vegetation. *Forest Ecology and Management*, *472*, 118253. <https://doi.org/10.1016/j.foreco.2020.118253>
- Su, X., Su, X., Zhou, G., du, Z., Yang, S., Ni, M., Qin, H., Huang, Z., Zhou, X., & Deng, J. (2020). Drought accelerated recalcitrant carbon loss by changing soil aggregation and microbial communities in a subtropical forest. *Soil Biology and Biochemistry*, *148*, 107898. [https://doi.](https://doi.org/10.1016/j.soilbio.2020.107898) [org/10.1016/j.soilbio.2020.107898](https://doi.org/10.1016/j.soilbio.2020.107898)
- Sulzman, E. W., Brant, J. B., Bowden, R. D., & Lajtha, K. (2005). Contribution of aboveground litter, belowground litter, and rhizosphere respiration to total soil CO₂ efflux in an old growth coniferous forest. *Biogeochemistry*, *73*, 231–256.<https://doi.org/10.1007/s10533-004-7314-6>
- Terrer, C., Phillips, R. P., Hungate, B. A., Rosende, J., Pett-Ridge, J., Craig, M. E., van Groenigen, K. J., Keenan, T. F., Sulman, B. N., Stocker, B. D., Reich, P. B., Pellegrini, A. F. A., Pendall, E., Zhang, H., Evans, R. D., Carrillo, Y., Fisher, J. B., van Sundert, K., Vicca, S., & Jackson, R. B. (2021). A tradeoff between plant and soil carbon storage under elevated CO₂. Nature, *591*, 599–603. <https://doi.org/10.1038/s41586-021-03306-8>
- Tian, P., Liu, S., Wang, Q., Sun, T., & Blagodatskaya, E. (2019). Organic N deposition favours soil C sequestration by decreasing priming effect. *Plant and Soil*, *445*, 439–451. [https://doi.org/10.1007/s1110](https://doi.org/10.1007/s11104-019-04331-3) [4-019-04331-3](https://doi.org/10.1007/s11104-019-04331-3)
- Torn, M. S., Swanston, C. W., Castanha, C., & Trumbore, S. E. (2009). Storage and turnover of organic matter in soil. In N. Senesi, B. Xing, & P. M. Huang (Eds.), *Biophysico-chemical processes involving natural nonliving organic matter in environmental systems* (pp. 219–272). John Wiley & Sons, Inc.
- Trumbore, S. (2009). Radiocarbon and soil carbon dynamics. *Annual Review of Earth and Planetary Sciences*, *37*, 47–66. [https://doi.](https://doi.org/10.1146/annurev.earth.36.031207.124300) [org/10.1146/annurev.earth.36.031207.124300](https://doi.org/10.1146/annurev.earth.36.031207.124300)
- UC Davis Soil Resource Laboratory. (2021). *SoilWeb: An online soil survey browser| California Soil Resource Lab*. [https://casoilresource.lawr.](https://casoilresource.lawr.ucdavis.edu/gmap/) [ucdavis.edu/gmap/](https://casoilresource.lawr.ucdavis.edu/gmap/)
- Vanderbilt, K. L., Lajtha, K., & Swanson, F. J. (2003). Biogeochemistry of unpolluted forested watersheds in the Oregon Cascades: Temporal patterns of precipitation and stream nitrogen fluxes. *Biogeochemistry*, *62*, 87–117. <https://doi.org/10.1023/A:1021171016945>
- Wallenstein, M. D., McNulty, S., Fernandez, I. J., Boggs, J., & Schlesinger, W. H. (2006). Nitrogen fertilization decreases forest soil fungal and bacterial biomass in three long-term experiments. *Forest Ecology and Management*, *222*, 459–468.<https://doi.org/10.1016/j.foreco.2005.11.002>
- Wang, J.-J., Pisani, O., Lin, L. H., Lun, O. O. Y., Bowden, R. D., Lajtha, K., Simpson, A. J., & Simpson, M. J. (2017). Long-term litter manipulation alters soil organic matter turnover in a temperate deciduous forest. *Science of the Total Environment*, *607–608*, 865–875. [https://](https://doi.org/10.1016/j.scitotenv.2017.07.063) doi.org/10.1016/j.scitotenv.2017.07.063
- Wang, X., Tian, Q., Li, Q., Liao, C., He, M., & Liu, F. (2018). Lignin characteristics in soil profiles in different plant communities in a subtropical mixed forest. *Journal of Plant Ecology*, *11*, 560–568. [https://doi.](https://doi.org/10.1093/jpe/rtx028) [org/10.1093/jpe/rtx028](https://doi.org/10.1093/jpe/rtx028)
- Wei, X., Shao, M., Gale, W., & Li, L. (2014). Global pattern of soil carbon losses due to the conversion of forests to agricultural land. *Scientific Reports*, *4*, 4062.<https://doi.org/10.1038/srep04062>
- Xu, S., Liu, L. L., & Sayer, E. J. (2013). Variability of above-ground litter inputs alters soil physicochemical and biological processes: A metaanalysis of litterfall-manipulation experiments. *Biogeosciences*, *10*, 7423–7433. <https://doi.org/10.5194/bg-10-7423-2013>
- Xu, S., Sayer, E. J., Eisenhauer, N., Lu, X., Wang, J., & Liu, C. (2021). Aboveground litter inputs determine carbon storage across soil

profiles: A meta-analysis. *Plant and Soil*, *462*, 429–444. [https://doi.](https://doi.org/10.1007/s11104-021-04881-5) [org/10.1007/s11104-021-04881-5](https://doi.org/10.1007/s11104-021-04881-5)

- Yan, G., Dong, X., Huang, B., Wang, H., Hong, Z., Zhang, J., Xing, Y., & Wang, Q. (2020). Effects of nitrogen deposition on litter decomposition and nutrient release mediated by litter types and seasonal change in a temperate forest. *Canadian Journal of Soil Science*, *100*, 11–25.<https://doi.org/10.1139/cjss-2019-0037>
- Yano, Y., Lajtha, K., Sollins, P., & Caldwell, B. A. (2005). Chemistry and dynamics of dissolved organic matter in a temperate coniferous forest on andic soils: Effects of litter quality. *Ecosystems*, *8*, 286– 300. <https://doi.org/10.1007/s10021-005-0022-9>
- Zelles, L. (1999). Fatty acid patterns of phospholipids and lipopolysaccharides in the characterisation of microbial communities in soil: A review. *Biology and Fertility of Soils*, *29*, 111–129.
- Zhao, Y., Ding, Y., Hou, X., Li, F. Y., Han, W., & Yun, X. (2017). Effects of temperature and grazing on soil organic carbon storage in grasslands along the Eurasian steppe eastern transect. *PLoS One*, *12*, e0186980.<https://doi.org/10.1371/journal.pone.0186980>

SUPPORTING INFORMATION

Additional supporting information can be found online in the Supporting Information section at the end of this article.

How to cite this article: Castañeda-Gómez, L., Lajtha, K., Bowden, R., Mohammed Jauhar, F. N., Jia, J., Feng, X., & Simpson, M. J. (2023). Soil organic matter molecular composition with long-term detrital alterations is controlled by site-specific forest properties. *Global Change Biology*, *29*, 243–259. <https://doi.org/10.1111/gcb.16456>