

AN ABSTRACT OF THE THESIS OF

Mark J. Hauck for the degree of Master of Science in Forest Science presented on December 1, 2006.

Title: Isotopic Composition of Respired CO₂ in a Small Watershed: Development and Testing of an Automated Sampling System and Analysis of First Year Data

Abstract approved:

Barbara J. Bond

Warming of the terrestrial biosphere due to the anthropogenic addition of carbon dioxide to the earth's atmosphere is becoming a major focus of scientific inquiry. Predictions of the extent of this warming are hampered by uncertainty in the ability of the earth's ecosystems to counteract this effect by sequestering carbon dioxide by increases in the mass of vegetation, soil storage, or storage in the Earth's oceans. Measurement of the carbon isotopic composition of respired CO₂ ($\delta^{13}\text{C}_{\text{R-eco}}$) is becoming increasingly important to ecosystem studies because the information contained in this respiration can be an indicator of ecosystem stress and productivity. This study was conducted as part of a larger research project aimed at developing and testing the capacity for measuring $\delta^{13}\text{C}_{\text{R-eco}}$ in a small, steeply-sloped watershed in western Oregon. The goals of this study were: (1) to develop and test an automated system for sampling nocturnal air to be analyzed in a laboratory for isotopic composition; (2) to collect samples of the

atmosphere from the nocturnal cold air drainage of a steeply sloped watershed in the HJ Andrews Experimental Forest and bring those samples back to a laboratory to be analyzed for CO₂ concentration and carbon isotope composition; and (3) to conduct an initial analysis of the relationship between $\delta^{13}\text{C}_{\text{R-eco}}$ for the first year of deployment and two environmental forcing factors, soil moisture and atmospheric vapor pressure deficit (VPD).

The automated system was designed during 2004 and 2005, tested in the laboratory during the month of April 2005, debugged in the field between August 2004 and May 2005, and deployed during the period from late May, 2005 to November 2005. The design objectives for the automated sampling system were: (1) light weight; (2) portability; (3) high reliability; (4) fast dynamic response; (5) unattended operation; and (6) the ability to capture, transport, and store samples over several days with no loss of data integrity. The automated sampling system proved capable of collecting 15 samples per sampling period and utilized a 16 loop stainless steel sample capture valve (Valcon Instruments, 7806 Bobbitt Houston, TX 77055) for sample containment. The system was designed for automated, labor free operation and uses a Campbell CR 10X datalogger (Campbell Scientific, Inc, 815 West 1800 North, Logan Utah) for both data acquisition and system control. This unique architecture was selected to enhance the ability of the system to capture samples which could be analyzed simultaneously for both CO₂ concentration and $\delta^{13}\text{C}$. The system was deployed in a steeply sloped watershed and proved to have a combination of light weight (approximately 34 kg total weight) and portability which allowed a wide range of field personnel to deploy the system without undue physical stress. Some initial issues with electrical wiring, plumbing connections,

and control program bugs hampered early season performance, but once the system was debugged, it functioned reliably throughout the remainder of the field season with a minimum of operator input. Test results showed that the automated sampler had a maximum sampling frequency of 0.043 hertz, could store air samples for up to 3 days without detectable changes in either isotopic composition or CO₂ concentration and displayed greater precision than the hand sampling process it replaced. Using the automated sampling system, sets of nocturnal cold air drainage samples were collected on nine evenings in 2005. The 15 samples acquired for each of these nine evenings were used to generate Keeling plots to determine a single value of the isotopic composition of ecosystem respiration ($\delta^{13}\text{C}_{\text{R-eco}}$) for each sampling date. The range of $\delta^{13}\text{C}_{\text{R-eco}}$ over the season was 3.9 ‰ and the values of $\delta^{13}\text{C}_{\text{R-eco}}$ varied from -26.2 ‰ on July 13, 2005 to -22.9 ‰ on September 14, 2005. This seasonal pattern of $\delta^{13}\text{C}_{\text{R-eco}}$ was consistent with a forest under drought stress. Trends in ecosystem respiration over the growing season were compared to corresponding measurements of the environmental variables of soil moisture content and VPD. VPD ranged from 2.7 to 1758 Pa, but these patterns of VPD were not significantly correlated to seasonal $\delta^{13}\text{C}_{\text{R-eco}}$ patterns over the study period. Soil moisture content ranged between 7.2 % and 44.3% over the study period and soil moisture content temporal patterns were highly correlated with rain events exceeding 4 mm. The soil moisture content pattern for the south-facing slope was significantly correlated with the seasonal $\delta^{13}\text{C}_{\text{R-eco}}$ pattern, although other groupings of soil moisture content were not.

The automated system designed for this project met all of its design objectives and functioned adequately throughout the sampling period. The carbon isotope patterns

were consistent with a forest under drought stress and the soil moisture content of the plots on the south-facing slopes of the watershed were significantly correlated to these isotope patterns. The future of this system could be enhanced by making adjustments to the supporting hardware, control program and operating procedure to enable larger quantities of samples, more rapid sampling rates, and automated hardware diagnostics.

© Copyright by Mark J. Hauck
December 1, 2006
All Rights Reserved

Isotopic Composition of Respired CO₂ in a Small Watershed: Development and Testing
of an Automated Sampling System and Analysis of First Year Data

by

Mark J. Hauck

A THESIS

submitted to

Oregon State University

in partial fulfillment of
the requirements for the
degree of

Master of Science

Presented December 1, 2006
Commencement June 2007

Master of Science thesis of Mark J. Hauck presented on December 1, 2006.

APPROVED:

Major Professor, representing Forest Science

Head of the Department of Forest Science

Dean of the Graduate School

I understand that my thesis will become part of the permanent collection of Oregon State University libraries. My signature below authorizes release of my thesis to any reader upon request.

Mark J. Hauck, Author

CONTRIBUTION OF AUTHORS

Dr. Barbara Bond assisted with data analysis, thesis editing, and planning of the thesis project. Dr. Thomas Pypker helped design, build, and test the automated sampling system, assisted with data collection and analysis, and thesis editing. Dr. Elizabeth Sulzman advised on general project issues and helped with data collection. Dr. Julia Jones provided guidance in thesis preparation and insight on statistical issues and data reduction. Zachary Kayler, Adam Kennedy, Aiden Padilla, Holly Barnard, Julian Licata, and David Conklin helped with collecting data and providing encouragement and feedback whenever needed.

TABLE OF CONTENTS

	<u>Page</u>
1 INTRODUCTION.....	1
1.1 Introduction.....	1
1.2 References.....	7
2 AN AUTOMATED SAMPLING SYSTEM FOR CAPTURING AIR SAMPLES AT NIGHT IN REMOTE STEEPLY SLOPED, FORESTED WATERSHEDS.....	9
2.1 Abstract.....	9
2.2 Introduction.....	10
2.3 Background.....	13
2.4 Autosampler Design Issues.....	18
2.5 System Description.....	18
2.5 Autosampler Gas Path.....	20
2.6 Autosampler Operation.....	21
2.7 System Testing.....	23
2.8 Results.....	26
2.9 Discussion.....	32
2.10 Conclusions.....	36
2.11 Future Direction.....	37
2.12 Acknowledgements and Credits.....	38
2.13 Table 2.1.....	39
2.14 Figures.....	40
2.15 References.....	52
3 KEELING PLOT ANALYSIS AND ASSOCIATIONS WITH ENVIRONMENTAL VARIABLES.....	54
3.1 Abstract.....	54
3.2 Introduction.....	55
3.3 Materials and Methods.....	58
3.4 Results.....	64
3.5 Discussion.....	67
3.6 Conclusions.....	75
3.7 Acknowledgments and Credits.....	76
3.8 Figures.....	77
3.9 References.....	88
4 SUMMARY AND CONCLUSIONS.....	92
BIBLIOGRAPY.....	99

Chapter 1

ISOTOPIC COMPOSITION OF RESPIRED CO₂ IN A SMALL WATERSHED: DEVELOPMENT AND TESTING OF AN AUTOMATED SAMPLING SYSTEM AND ANALYSIS OF FIRST YEAR DATA

Introduction:

The isotopic composition of forest ecosystem respiration ($\delta^{13}\text{C}_{\text{R-eco}}$) potentially contains a great deal of information about the metabolism of the plants inhabiting that landscape. Although many studies have been carried out to illuminate the linkages between ecosystem respiration and the environmental variables that influence this metabolism (Bowling et al. 2002, Ehleringer et al. 1997), most have been aimed at sites on flat ground, while relatively few have focused on ecosystems situated on steeply sloped, complex terrain. While current techniques for data gathering and analysis are well suited to flat sites with good access, the extension of these techniques to remote, steeply sloped sites has not been advanced. Many of the world's most productive forested ecosystems reside on steeply sloped sites with poor access. New tools are necessary to aid in the study of these ecosystems on complex sites. This thesis is part of a larger project aimed at development of such tools. The goals of this thesis were: (1) to develop and test an automated system for sampling nocturnal air to be analyzed in a laboratory for isotopic composition; (2) to collect samples of the atmosphere from the nocturnal cold air drainage of a steeply sloped watershed in the HJ Andrews Experimental Forest and bring those samples back to a laboratory to be analyzed for CO₂ concentration and carbon isotope composition; and (3) to conduct an initial analysis of the relationship between

$\delta^{13}\text{C}_R$ for the first year of deployment and two environmental forcing factors, soil moisture and atmospheric VPD.

The linkages between $\delta^{13}\text{C}_{R\text{-eco}}$ and plant physiological characteristics have been well established (Farquhar et al., 1989, Flanagan et al., 1996). Photosynthetic discrimination against the heavier stable carbon isotope (^{13}C) is modulated by plant physiological response to certain stresses, such as drought stress (Ehleringer and Cook, 1997), high salinity (Meinzer et al., 1994), lack of nutrients (Meinzer and Zhu, 1998), as well as to high rates of assimilation (Ehleringer, 1993). Seasonal variations in isotopic discrimination have also been reported (Buchman et al. 1997). Because the plant stores a record of these stress states in the form of varying ratios of carbon isotopes in the sugars fixed during photosynthesis, it is possible to relate plant tissues and plant respiration to earlier states of stress. Furthermore, because there appears to be little or no fractionation of isotopes of carbon after the initial photosynthetic fractionation (Lin and Ehleringer, 1997), CO_2 that is respired from plants and plant ecosystems can “carry” an isotopic signal that reflects the physiological status of plants at the point of fixation. Ekblad and Hogberg (2001), for example, used isotopic respiration signals to demonstrate that there was a time lag of 2 to 4 days between the time of carbon fixation and its subsequent release from soils as respiration, with about 70% of the ecosystem respiration being derived from those recently fixed carbon sources. This relationship has been shown to hold in some subsequent studies of forest ecosystems (e.g., Bowling et al. 2002, Knohl, 2005) but not in others (e.g., McDowell et al. 2004).

In almost all previous studies of $\delta^{13}\text{C}_{\text{R-eco}}$, sampling has been restricted to relatively flat sites. This is because the analysis procedure used to determine $\delta^{13}\text{C}_{\text{R-eco}}$ (described later in this thesis) requires a broad range of CO_2 concentrations, and in all previous studies this range of $[\text{CO}_2]$ has been obtained by collecting air samples through a vertical concentration gradient that develops in still air at night on many flat sites. On sloping sites, however, advective airflow patterns at night are likely to “break up” the vertical gradients in CO_2 concentration that are required for sampling and analyses to determine $\delta^{13}\text{C}_{\text{R-eco}}$. However, a pilot study carried out by Pypker et al. (2006, in press) shows there is another possibility for sampling air to determine $\delta^{13}\text{C}_{\text{R}}$ in very steeply-sloped areas. This study demonstrated that nocturnal cold-air drainage is a regular occurrence in a small watershed in western Oregon. Furthermore, the concentration of CO_2 within this airflow increased substantially between dusk and pre-dawn. Thus, it is possible in this watershed to obtain a sufficient range of $[\text{CO}_2]$ for determination of $\delta^{13}\text{C}_{\text{R-eco}}$ by repetitive sampling over time rather than through a vertical gradient. The results of Pypker et al. (2006) suggest an opportunity, for the first time, for intensive studies of $\delta^{13}\text{C}_{\text{R-eco}}$ in a mountainous ecosystem.

Ecologists have developed several tools and techniques for gathering the air samples for isotopic analysis of ecosystem respiration. Pioneering efforts in the field include manual sampling of air in glass flasks of varying size (Ehleringer and Cook, 1997; Ometto et al., 2002; Pataki et al., 2003). These systems proved capable of providing high quality samples but were vexed by demanding equally high quotients of labor; workers typically need to work through the night to gather the necessary samples.

In order to address the issue of excessive, off-hours labor input, others (Shauer et al., 2002; Theis et al., 2004) have developed automated sampling systems which also fill glass flasks. These systems too have proved capable of providing high quality samples and have addressed the labor issue handily. However, for operation in remote, mountainous locations, these systems are too delicate due to the use of glass flasks and too heavy to transport into these challenging sites. In addition, the large flasks are not well suited to the Gas Bench system that is now commonly used as a front-end in isotope ratio mass spectrometry.

A brand new technology involving tunable diode laser (TDL) systems may eventually become the standard for all studies of isotopes in canopy air (Bowling et al., 2003). This technology has several advantages. The analysis device is installed in situ rather than in a laboratory. It is coupled to the sampling mechanism so that the analysis is nearly instantaneous, avoiding the thorny transport issues associated with sampling being separated from analysis, as in traditional flask sampled, mass spectrometer analyzed systems. Another major advantage is that the system operates on a nearly continuous stream of samples and thus produces data sets with large numbers of data points spanning equally large time periods. However, TDL systems have some disadvantages associated with them as well. At the current state of development, these systems are quite large and heavy, making them difficult to transport to locations in rugged terrain. They also require cryogenic cooling, which adds another level of complexity and cost. The power requirements of a TDL system require either the use of a generator or access to line power, both of which are prohibitive at remote sites. Given the above, although

tantalizing for their attributes, the disadvantages of current TDL systems make them unsuitable at the current time for most ecosystem studies.

The attributes of a well-designed air sampling system for use in remote and rugged locations are somewhat mutually exclusive. The ideal system would be highly automated to preserve labor resources and prevent the fatigue associated with operator attended all night sampling, yet be simple enough for troubleshooting and repair by a wide range of field personnel with varying skill sets. It would be small and light, yet big enough to be easily serviced when necessary. It would minimize power consumption to conserve on that precious commodity that is difficult to come by in such locations, yet have sufficient dynamic response necessary to capture samples in a rapidly changing ecosystem. A well-designed system would be highly sensitive to changes in CO₂ concentration so that the desired samples could be captured, yet rugged enough to survive the rigors of the field conditions encountered. The system design and materials would allow samples to be gathered and transported without introducing isotope fractionation so as to provide high fidelity respiration data, yet not so exotic as to increase cost beyond an acceptable limit. The system would be expandable to meet future needs and field modifiable to adapt to changing or unanticipated conditions, yet it would be reliable so that the stream of temporal data is not interrupted. Providing for all of these attributes while avoiding the associated pitfalls requires making some hard choices about where to expend resources.

This study was undertaken to address three major objectives. The first objective was to develop an automated air sampling system to collect air samples for laboratory

analysis of the isotopic composition and the concentration of CO₂. The second objective was to determine, through Keeling plot analysis, the seasonal values for the isotopic composition of ecosystem respired CO₂ for a steeply sloped watershed in the HJ Andrews research forest. The final objective was to establish whether or not relationships between the isotopic ratio of the ecosystem respiration and the environmental variables of soil moisture and VPD exist for this watershed. Chapter 2 of this thesis reports on the results of the automated sampling system development while Chapter 3 presents the data that were produced using the automated sampler as well as the analysis of those data.

References:

- Bowling DR, Sargent SD, Tanner BD, and Ehleringer JR, 2003, Tunable diode laser absorption spectroscopy for stable isotope studies of ecosystem-atmosphere CO₂ exchange, *Agricultural and Forest Meteorology.*, 118, 1-19.
- Buchman N, Guehl, JM, Barigah TS, Ehleringer JR, 1997, Interseasonal comparison of CO₂ concentrations, isotopic composition, and carbon dynamics in an Amazonian rainforest (French Giana), *Oecologia*, 110: 120-131
- Ekblad A, Högberg P. 2001. Natural abundance of ¹³C in CO₂ respired from forest soils reveals speed of link between tree photosynthesis and root respiration, *Oecologia* 127: 305-308
- Ehleringer, JR, 1993. Variation in leaf carbon isotope discrimination in *Encelia farinosa*: implications for growth, competition, and drought survival, *Oecologia*, Volume 95, number 3
- Ehleringer JR, Cook CS, 1997, Carbon and oxygen isotope ratios of ecosystem respiration along an Oregon conifer transect, preliminary observations based upon small-flask sampling, *Tree Physiology*, 18, 513-519
- Farquhar, G.D., Ehleringer, J.R., Hubick, K.T., 1989. Carbon isotope discrimination and photosynthesis. *Annu. Rev. Physiol. Plant Mol. Biol.*, 40, 503-537
- Flanagan, LB, J.R. Brooks, G.T. Varney, S.C. Berry, J.R. Ehleringer Carbon isotope discrimination during photosynthesis and the isotope ratio of respired CO₂ in boreal forest ecosystems, *Global Biogeochemical Cycles* 96GB02345
Vol. 10 , No. 4 , p. 629
- Knohl, A., Werner, R.A., Brand, W.A., Buchman, N., (2005), Short-term variations in $\delta^{13}\text{C}$ of ecosystem respiration reveals link between assimilation and respiration in a deciduous forest, *Oecologia* ,142: 70–82
- Lin, G., Ehleringer, J.R., 1997, Carbon isotopic fractionation does not occur during dark respiration in C₃ and C₄ plants, *Plant Physiology*, Vol 114, Issue 1, 391-394
- Meinzer, F.C., Z. Plaut, and N. Z. Saliendra, 1994, Carbon Isotope Discrimination, Gas Exchange, and Growth of Sugarcane Cultivars under Salinity, *Plant Physiol.* 1994 February; 104(2): 521–526
- Meinzer, F.C, Zhu, J., 1998. Nitrogen Stress reduces efficiency of the C₄CO₂ in the concentrating system, and therefore the quantum yield, in *Saccharum* (sugarcane) species, *Journal of Experimental Botany*, Vol. 49, Number 324

Ometto JHB, Flanagan LB, Martinelli LA, Moreira MZ, Higuuchi N, Ehleringer JR, 2002, Carbon isotope discrimination in forest and pasture ecosystems of the Amazon Basin, Brazil, *Global Biogeochemical Cycles*, Vol. 16, No.4

Pataki DE, Ehleringer JR, Flanagan JB, Yakir D, Bowling DR, Still DJ, Buchman N, Kaplan JO, Berry JA, 2003, The application and interpretation of Keeling plots in terrestrial carbon cycle research, *Global Biogeochemical Cycles*, Vol 17, No.1

Pypker, T G, M H Unsworth, A C Mix, W Rugh, T Ocheltree, K Alstad and B J Bond (2006). Using nocturnal cold air drainage flow to monitor ecosystem processes in complex terrain: a pilot study on the carbon isotopic composition and advection of ecosystem respiration, in press

Schauer, AJ, Lai, CT, Bowling, DR, Ehleringer, JR, 2003, An automated sampler for collection of atmospheric trace gas samples for stable isotope analyses, *Ag. And For. Meteorology* (118), 113-124

Theis D, Saurer M, Blum H, Frossard E, Siegwolf TW, 2004, A portable automated system for trace gas sampling in the field and stable isotope analysis in the laboratory, *Rapid Communications in Mass Spectrometry*, Vol, 18, Issue 18, 2106-2112

Chapter 2

AN AUTOMATED SAMPLING SYSTEM FOR CAPTURING AIR SAMPLES AT NIGHT IN REMOTE STEEPLY SLOPED, FORESTED WATERSHEDS

Abstract:

An automated sampling system was designed, built, tested and deployed in 2005 to collect samples of air for determination of the isotopic composition of ecosystem respiration ($\delta^{13}\text{C}_{\text{R-eco}}$). The design objectives for this automated system included light weight, portability, reliability, fast dynamic response, unattended operation, and the ability to capture, transport, and store samples over several days with no loss of data integrity with respect to the critical measurement values of CO_2 concentration and carbon isotopic composition. The automated sampling system proved capable of collecting 15 samples per sampling period and utilized a 16 loop stainless steel sample capture valve (Valcon Instruments, 7806 Bobbitt Houston, TX 77055) for sample containment. The system was designed for automated, labor free operation and uses a Campbell CR 10X datalogger (Campbell Scientific, Inc, 815 West 1800 North, Logan Utah) for both data acquisition and system control. This unique architecture was selected to enhance the ability of the system to capture samples, which could be analyzed simultaneously for both CO_2 concentration and $\delta^{13}\text{C}$. The system was deployed in a steeply sloped watershed and proved to have a combination of light weight and portability which allowed a wide range of field personnel to deploy the system without undue physical stress. Some initial issues with electrical wiring, plumbing connections, and control program bugs hampered early season performance, but once the system was debugged, it functioned reliably throughout

the remainder of the field season with a minimum of operator input. Test results showed that the automated sampler had a maximum sampling frequency of 0.043 hertz, could store air samples for up to 3 days without detectable changes in either isotopic composition or CO₂ concentration and displayed greater precision than the hand sampling process it replaced. Using the automated sampling system, sets of nocturnal cold air drainage samples were successfully collected on nine evenings in 2005, demonstrating the capability of the automated system to perform as per design intent.

Introduction:

The carbon isotopic composition of forest ecosystem respiration ($\delta^{13}\text{C}_{\text{R-eco}}$) is potentially rich in information related to the health and vitality of plants inhabiting that landscape (Flanagan et al., 1996; Flanagan et al. 1999; Fessenden et al. 2003,). There is promise for using $\delta^{13}\text{C}_{\text{R-eco}}$ as a monitoring tool to evaluate ecosystem responses to environmental variation over a range of time scales (Eckblad and Hogberg 2001; Bowling et al. 2002). In order to capture this information efficiently and accurately, an integrated system for field collection of samples, sample transport, and data analysis is required. Over a two-year period, we developed such a system and methodically tested alternate designs and procedures. This system provides an alternative sampling method to the traditional flask samples that are frequently used to collect samples for Keeling plot analysis (Ehleringer and Cook 1997; Shauer et al. 2003).

Charles Keeling (1958, 1961) developed a technique to identify the isotopic composition of ecosystem-respired CO₂ in samples of air collected in and above a plant

canopy. The Keeling method requires that samples be collected over a range of CO₂ concentrations at night, when photosynthesis is not occurring. These samples are then analyzed to determine their ratio of ¹³C to ¹²C ($\delta^{13}\text{C}$) and CO₂ concentration. The carbon isotopic ratio of ecosystem respiration ($\delta^{13}\text{C}_{\text{R-eco}}$) is determined from this set of samples collected over an evening using a graphical approach to solving a two end-member mixing model (for an example of a Keeling plot, refer to Figure 2.11). To collect samples for a single analysis (one data point), investigators have typically needed to spend three to eight nighttime hours carefully monitoring natural fluctuations in CO₂ concentrations, collecting samples in glass flasks until the required concentration range is obtained (Ehleringer and Cook, 1997; Ometto et al., 2002; Pataki et al., 2003). The samples are quickly returned to a laboratory for analysis to minimize the possibility of fractionation due to slow diffusion into or out of the flask. This is a time-consuming and expensive process, especially when the study site is remote or difficult to access.

The cumbersome method of hand collecting air samples has stimulated interest in automating the sample collection process. Shauer et al. (2002) and Theis et al., (2004) have developed automated sampling systems to fill gas flasks. Shauer's automated sampler uses a pump to move sample air through a manifold to a rotary valve. Glass flasks with a 100ml volume are interfaced to this rotary valve. The sample gas is pumped through one channel of the rotary valve and the valve is cycled when it is desired to capture a sample. The valve is then sequenced to the next position and the cycle repeats itself. Using such automated techniques allows researchers to capture samples without

committing field personnel to continuous duty, which minimizes the labor necessary to get research data.

After samples are collected in the field, they are returned to the laboratory for analysis of isotope composition. In the case of the glass flasks sampled by Schauer et al. (2002), the samples were analyzed by withdrawing subject gases from the flasks with a 500 ml syringe and then injecting these gases into a continuous flow isotope ratio mass spectrometer (Finnigan MAT 252 or DeltaS, Finnigan MAT, San Jose, CA), which is preceded by a pre-condensing unit in series with a gas chromatograph column. This process, although extremely accurate, is both time and labor intensive and is also very costly. In order to automate the sample analysis process (thus lowering cost per sample), many other investigators are currently using the Finnigan Gas Bench II front end for the introduction of samples to the MAT 252 mass spectrometer. The Gas Bench II utilizes an automated robotic arm fitted with a dual flow needle arrangement to extract samples from vials and introduce them into the mass spectrometer. The samples are contained in glass vials (exetainers) which are arranged in a rack. This rack is then visited sequentially by the robotic arm and samples are introduced in that sequence to the mass spectrometer. Thus, the data are ordered by the sequence of sample introduction and can be later coordinated with the sample's meta-data. The automated sampling system developed by Schauer et al. (2002) was designed for glass flasks and this is not compatible with the septa-topped exetainers commonly used with the Gas Bench II. For researchers who endeavor to decrease labor expenses by automating both the sample collection and sample analysis systems, this lack of compatibility presents a difficult dilemma.

The current project was initiated to develop and test an automated air sampling system to be used as part of the HJ Andrews Airshed project. The Andrews Airshed project started in 2002 as a pilot project to determine, among other things, if the $\delta^{13}\text{C}_{\text{R-eco}}$ could be quantified in steeply sloped, forested watersheds (Pypker et al. in press). Isotope analyses for this project were conducted in the laboratory of Dr. Alan Mix, using a Gas Bench II system. Thus, the automated glass flask sampling system developed by Shauer et al. (2002) was not considered a good option for this project. Instead, a new system, which allowed automated transfer from the sample capture containers, was developed.

Background:

Several non-automated air sample collection techniques were tested during the pilot phase of the Airshed Project (before this M.S. project was initiated). Each of these techniques required an operator to monitor a field portable infrared gas analyzer, or “IRGA” (LI-6252, LI-COR Inc., Lincoln, NE) in order to determine the appropriate sample collection times. The overall objective in this decision process is to obtain samples with the largest possible range of CO_2 concentrations. During this phase of the study, the field $[\text{CO}_2]$ measurements were used directly in Keeling plot analyses (described in the introduction). Later, thanks to advances in laboratory analyses, the CO_2 concentration was measured by the mass spectrometer concurrently with the isotopic analysis. Because the CO_2 concentration of the canopy air fluctuates rapidly, and the infrared gas analyzer was necessarily in series with the sample collection tubes, there could be significant differences between the measured CO_2 concentration and the true

concentration of gas sampled in the vials. Thus, it was determined that a buffer volume preceding the sample location was necessary to damp out these rapidly fluctuating CO₂ signals.

To avoid contaminating the samples with operator respiration a “glove bag” was used for filling the glass vials. The operator placed open vials and their caps in the bag, which was “plumbed” to the airflow line. At a desired CO₂ concentration, the operator capped the vials by working through the plastic walls of the bag. Tests were performed as part of the Airshed Pilot project to determine the best combination of glass vial and septum for field sampling. It was determined that either butyl rubber or Kel-F septa coupled to 20ml glass vials supported accurate data retrieval over time spans in excess of 24 hours and through altitude changes between sample point and analysis lab of about 500 feet (Pypker, private communication). The Gas Bench II uses a double-needle sampling system to “sip” gas samples from vials that are capped with septa. The vials are arranged on a grid, and the sampler is engineered to move automatically through this grid. This is a very efficient sampling system, and it is conceptually appealing to use a similar system to automate the collection of air samples in the field. However, the very small hole left by a needle might result in a very minute leak, which could lead to fractionation of isotopes in the sample. A series of tests were conducted as part of the Airshed Pilot Project to evaluate this concern. It was discovered that, during transportation, sample leakage from the punctured septa and associated isotopic fractionation reduced sample fidelity. This problem was confounded when the punctured vials were transported from the relatively high altitude of the field site to the lower

altitude of the laboratory. Tu et al. (2001) found similar results, with isotopic precision decreasing by about 0.07 permil for each day that samples were stored in vials capped with punctured septa.

The next attempt to automate the sample collection system employed the use of a portable automated water collection system (ISCO 6712, Teledyne ISCO, Inc, 4700 Superior Street, Lincoln NE). The ISCO automated sampler uses a pump to circulate water through the body of the sampler. When a signal from the controller indicates that a sample should be taken, this flow is temporarily diverted to one of 99 sample vials. The interface between these vials and the circulation system is from industry standard needle/septa. It was thought that this device could be modified to circulate air rather than water, which would provide the labor saving automation necessary to fill small glass vials with air samples without the development effort necessary to design a customized system. These vials could then be loaded into the sample tray of the Gas Bench II, where the samples could be automatically extracted and analyzed. However, the ISCO design did not prove to be a good candidate for use in collecting gases. There was concern that the pumps might not move adequate volumes of air to keep the sample lines purged and that these pumps might cause isotopic fractionation. Also, as stated earlier, piercing the septa was shown to cause isotope fractionation. In the end, it was determined that too much effort would have to be expended modifying the ISCO system to make it a practical device for sample collection.

At this point in the project, attention turned to a completely different scheme for automating sample collection that avoided the use of vials altogether. This new scheme

involved the use of a Valco 16 port sample collection valve (Valco E2SD16MWE, Valcon Industries, INC, 7806 Bobbitt, Houston, TX). This valve is a specialized sample collection valve used extensively in laboratory instruments such as mass spectrometers and gas chromatographs. There are 16 sets of inlet and outlet ports, which are addressed by a spring loaded indexing distribution hub. This hub seals each set of ports from the rest of the system and opens one set at a time for sample collection. The sample passes through the distribution hub to an inlet port. Connected between the inlet port and the outlet port is a length of stainless steel tubing (or “loop”), which serves as the collection volume. The sample is then passed through this tubing to the outlet port, back through the distribution hub, and exits the collection system. When the electronic valve controller receives a signal from the system controller to take a sample, the hub is indexed by an electric stepper motor, isolating and sealing the particular set of inlet and outlet ports thus trapping the sample in the loop. Sequential indexing of the distribution hub traps samples in each set of the 15 loops, which comprise the valve. Dawson et al. (Rugh, private communication) tested a similar sample collection scheme, but found that isotope fractionation occurred during sample storage. However, our own initial tests showed no indication of fractionation in the Valco valve collection. It is possible that there may have been some residual organic cleaning compounds retained in the sample loops in the Dawson et al. tests, which may have contributed to the fractionation signal observed. This problem was avoided by flushing all sample loops for 48 hours with dry N₂ prior to starting fractionation testing and all subsequent field deployments.

Due to the highly negative impact on $\delta^{13}\text{C}_{\text{R-eco}}$ of large errors in measurement of $[\text{CO}_2]$ (Zobitz et al., 2004), it was not considered adequate to rely on field measurements of $[\text{CO}_2]$ recorded when samples were captured. The Oregon State University College of Oceanic and Atmospheric Sciences Stable Isotope Mass Spectrometry Facility (OSU/COAS SIMSF) came up with a unique solution to increase the precision of measurement of $[\text{CO}_2]$ by using the integral of the mass spectrometer voltage output signal to determine $[\text{CO}_2]$. This technique provided the advantage that the CO_2 concentration and the isotopic signature for a particular measurement came from the same gas volume.

The objective of this part of the thesis was to design and test an automated system for collecting air samples based on the Valco valve to be used for isotope analyses. The design objectives for this automated system included light weight, portability, reliability, fast dynamic response, unattended operation, and the ability to capture, transport, and store samples over several days with no loss of data integrity with respect to the critical measurement values of CO_2 concentration and carbon isotopic composition. Laboratory and field testing, and their associated results, are presented. Discussion of the results and recommendations for system improvements are also provided.

Autosampler Design Issues:

The design of the automated sampling system for the HJ Andrews Airshed Project was guided by the need to provide air samples that accurately and precisely represented the $\delta^{13}\text{C}_{\text{R-eco}}$ within the watershed. It was desired that the system operate automatically to minimize labor, especially as this labor is required for sampling between 1800 and 0200 hours. The location of the pilot study sample site is somewhat remote and difficult to access, so light weight and portability were also highly desired. Design and operation simplicity was sought to increase the range of personnel able to deploy the system. As this system was to be the model for future systems, which would be semi-permanently installed to monitor ecosystem function, it was desired that this system be both flexible and field modifiable so that it could easily be adapted to various locations.

As mentioned above, a Finnigan Gas Bench II front end was coupled to a Finnigan Delta Plus MAT 252 mass spectrometer to analyze the air samples provided by the automated sampling system. Typically, samples are collected in small glass vials, or exetainers, for use on the Gas Bench II. The sample loops used in the Valco valve system created some unique constraints on the Gas Bench system. These issues were addressed by the laboratory technical staff, resulting in a custom interface between the sample containment valve and the Gas Bench II inlet.

System Description:

The automated sampling system is contained in two separate enclosures, a “Control Enclosure” and a “Sample Valve Enclosure” (Figure 2.1). Both enclosures were

constructed of 0.95cm marine grade plywood with overall dimensions of 61cm x 61cm x 61cm. Wood was chosen for the initial construction of the enclosures due to the ease of field modification and low cost. As the design matures and final component choices are made, a transition to electrical grade fiberglass enclosures would increase the durability of the system and provide better protection of interior components.

Two enclosures were needed to separate the sample-trapping valve from the system control components. The Sample Valve Enclosure contains the 16 port Valco sample containment valve and its electronic controller. A 155 cm³ polycarbonate cylinder filled with magnesium perchlorate is mounted to the outside of this enclosure and is used to remove water from the sample stream before it enters the delicate internals of the sample valve. All plumbing connections inside this box are 1/8" stainless steel tubing with Swagelok fittings. The Control Enclosure contains a LI-COR 6252 infrared gas analyzer (IRGA), a Campbell CR10X data logger (Campbell Scientific, Inc, 815 West 1800 North, Logan Utah), a six port Valco selector valve and its controller (Valcon E2SD6MWE, Valcon Instruments, 7806 Bobbitt Houston, TX), the sample and purge pumps, the solenoid valves that isolate the calibration gases, and the miscellaneous electrical connectors that are necessary to provide power and control signals to the various devices. The Sample Valve Enclosure weighs 14.3 kg, while the Control Enclosure weighs 20.6 kg. See Table 2.1 for a detailed list of the components, which make up the automated sampling system.

Autosampler Gas Path:

The automated sampling system retrieves air samples drawn through three ¼ inch diameter tubes that are affixed to an adjacent 37 m tower, allowing air intake from heights of 3m, 10m, and 28m, respectively. Each of these separate tubes has a 5 micron filter attached so that fine particulates and water droplets are removed from the air stream. Each of these tubes enters a bulkhead fitting mounted to the Control Enclosure. From the bulkhead fitting, 1/8" diameter nylon tubing carries the air sample to the Valco 6 port selector valve. The automated calibration gases enter the Control Enclosure in a similar manner, running through separate solenoid valves on their route to the Valco selector valve. The purge pump and the sample pump are also connected to the Valco selector valve. The sample pump outlet exits the Control Enclosure through a bulkhead fitting and connects to a 200cm³ water trap filled with magnesium perchlorate mounted to the outside of the Sample Valve Enclosure. The outlet of the water trap is connected to the 16 Port Valco Sample Valve through a bulkhead fitting, which is mounted to the Sample Valve Enclosure. The output of the 16 Port Sample Valve exits the enclosure via a bulkhead fitting and is connected to an infra-red gas analyzer (IRGA) (LI-6252, LICOR Inc., Lincoln, NE, USA) through a bulkhead fitting mounted to the Control Enclosure. The output of the IRGA is exhausted to the exterior of the Control Enclosure via another bulkhead fitting. The Campbell Data Logger, mounted inside the Control Enclosure, receives input signals from the IRGA and sends output signals to the sample pump, purge pump, Selector Valve, field calibration gas solenoids, and Sample Valve

based on the control sequence dictated by the Logger's programming. See the schematic diagram of the Automated Sampling System for details (Figure 2.1).

Autosampler Operation:

The Autosampler is connected to a CR10X Campbell data logger. The automated functions of the sampling system are triggered by software written by another member of our research team for the datalogger. This software is responsible for providing the timing of the control signals, which activate the sample and purge pumps, selector valve, and sample valve. Initially, the control program starts the system pumps and infrared gas analyzer (IRGA) one hour before sunset so that the IRGA has a sufficient warm-up period before sampling begins. During this warm-up period, a technician performs a leak test of all of the connections of the automated sampler and the IRGA is manually calibrated for CO₂ measurements using three field calibration gases. One hour after sunset, the control program triggers the system to begin taking samples. This typically occurs around 9pm PST in the summer months and is intended to be coincident with the onset of cold air drainage in the watershed. The program is parameterized by the technician to achieve a minimum range of CO₂ concentration (usually around 30ppm), which it uses to establish set intervals of CO₂ concentration for each height sampled. A minimum of 5 minutes must pass between each sample cycle so that flushing of all connection tubing between the sample introduction point on the tower and the sample collection valve is assured. The air sample passes through the selector valve from the tower tubing and is pumped by the sample pump through the water trap. From the water

trap, the air sample passes through the 16 Port Sample Valve and into the IRGA. The IRGA sends a continuous CO₂ concentration signal to the data logger. When a pre-determined CO₂ concentration value is achieved, the control program triggers the data logger to signal the sample valve to index, trapping an air sample in one of its loops. The pre-determined concentration value is a product of a calculation between two operator-defined values in the control program. To arrive at this concentration, the desired CO₂ range is divided by the number of samples at each height (the total number of samples divided by the number of sample heights). The results of this calculation determine the concentration that must be achieved before the sample can be collected. Once this pre-determined concentration is achieved and the sample collection valve has indexed, the logger then signals the selector valve to select a new height and the program repeats its operation. A purge pump continuously purges the two tower tubes not being currently sampled with the aim of keeping the air moving through those tubes fresh. After the loops are filled the program progresses the system to an idle state, shutting off the pumps and the IRGA. If all the loops are not filled by one hour before sunrise the system enters “panic” mode. The system then sequentially fills the remaining loops regardless of the CO₂ concentration of the air. The program records the CO₂ concentration (from the IRGA) each time the Sample Valve is directed to obtain a sample, along with the time of sample. The temperature of the interior of the Control Enclosure is also recorded to monitor the operating temperature of the IRGA.

System Testing:

Testing of the automated sampling system was performed in both laboratory and field situations to quantify the ability of the sampler to deliver high fidelity air samples. All of the isotopic analyses for this project were performed at the OSU/COAS SIMSF using the Finnegan MAT 252 Mass Spectrometer coupled with the automated Gas Bench II sample introduction process. Shauer et al. (2003) performed extensive testing of different materials of construction of their automated sampling system, looking for isotope fractionation effects attributable to these materials. Their results indicated that stainless steel tubing, Viton o-rings, and Valcon M grade (high density polyethelene) seal interface materials performed well. We chose to use the materials recommended by Shauer et al. (2003) and focused our testing on how the system performed with respect to isotopic fractionation and leakage or diffusion of CO₂ concentration. A series of laboratory tests and field tests were conducted to determine: (1) the maximum time samples could be retained in the system before they no longer represented the gases from which the samples were produced; (2) the comparison of accuracy and precision of both isotopic composition and CO₂ concentration of the automated sampling system as compared to hand sampling; (3) the system's dynamic response; and (4) the integrity of the system's plumbing from week to week.

Storage Test

This test was performed in the laboratory by filling each loop of the Valco sample storage valve (see illustration in Appendix B) with gas from a NOAA certified tank (nominal CO₂ concentration 959 ppm \pm 3.5ppm). This high [CO₂] gas was chosen for the

test because it is markedly different from the ambient laboratory gas, which makes detection of concentration changes more apparent. On each of five sequential days, three loops were filled with gas. Each sample was taken after gas flowed through a loop at 50 ± 1 ml/min for 5 minutes. The entire system was flushed with the same NOAA gas for 15 minutes prior to the test at a rate of 100 ± 1 ml/min prior to capturing samples to ensure that no trace gases contaminated the samples. After filling three loops each day for five days the samples were run on the MAT 252 Mass Spectrometer, with the most recently filled loops run first.

Hand Sampling comparison to Automated Sampling

To compare the hand and automated sampling techniques we simultaneously filled the automated sampler sample loops and glass exetainers with CO₂ of known concentration. Similar to the Storage Test outlined above, the automated sampling system was flushed with the field calibration gas for 15 minutes prior to capturing samples. Once the flush process was completed, gas was allowed to flow through the automated sampling system. For each loop sampled, an exetainer glass vial was filled just prior to capturing the loop sample by flushing the exetainer vial from the exhaust of the automated system and then quickly attaching a cap, which contained an un-punctured butyl rubber septa. This technique was similar to the one used in the field to capture hand samples used in the pilot study (Pypker et al., in press).

Dynamic Response Test

The Dynamic Response test was performed in the field, with the automated sampling system connected as it is during normal operation. This testing arrangement provided realistic conditions for determining system dynamic response, including proper length and orientation of air supply tubing, potential kinks and bends, and plugging of inlet filters, all of which would be difficult to simulate in a laboratory setting. The program that controls the automated sampler was modified from its field sampling form to allow continuous operation with operator controlled valve switching. Measurements of CO₂ concentration stabilization time after switching heights were recorded for each of 4 different heights with 5 data points being taken per height. This stabilization time represents the minimum amount of time the system must be allowed to flush for a particular height before a sample can be captured and is an indication of how quickly the system can respond to changes in CO₂ concentration.

Weekly System Integrity Tests

During the field season, canopy gas samples were drawn each week. To be certain that the air samples captured by the automated sampling system were not being compromised during transport back from the field site to the laboratory, two of the sample loops were filled with field calibration gas at the field sampling site. These loops were filled at random with two different field calibration gases, each with different values of CO₂ concentration and isotopic ratios. The field calibration gas samples were analyzed in the laboratory simultaneously with the rest of the field air samples. After laboratory

analysis of all of the samples, all of the loops were then flushed and re-filled in the laboratory with the two field calibration gases used to fill the two loops in the field (half the loops with one gas and half with the other). The resulting analysis of these laboratory-filled loops provided a statistical basis for comparison of the weekly field filled loop samples. Any large differences between the field filled loops and the laboratory filled loops would trigger an internal investigation of the integrity of the loops or the field data. Inspection of the standard deviation of the laboratory data would also indicate outliers, which serve as a second check of loop integrity.

Results:

Storage Test

The mean values for $\delta^{13}\text{C}$ of each of the timed storage samples up to 75 hours fell within the 95% confidence interval for the NOAA gas tested, while the values for the 97-hour samples fell just outside the same confidence interval (Figure 2.2). The CO_2 concentration within the loops fell within the 95% confidence interval (within $\pm 7\text{ppm}$) of the NOAA gas used for the test for the entire five-day test period, although there was some indication of a slight positive drift in $[\text{CO}_2]$ for the final two days (Figure 2.3). The confidence interval for the NOAA gas was established by repeatedly analyzing samples that had been flush/filled into exetainers. The 95% confidence interval values produced by this flush/fill process are outside those provided by NOAA for the standard gas; however, they represent a more realistic accounting of how well our mass spectrometer can measure the concentration (or isotopic signature) of the gas in question. Thus, this is

a better test of the ability of the automated sampling system to retain these data over time. In any event, it is instructive to realize that, for a given Keeling plot analysis the change in the $\delta^{13}\text{C}_{\text{R-eco}}$ for a constant offset variation of $\pm 2\text{ppm}$ (the amount of variation of the storage test values) would be about $\pm 0.09\text{‰}$ (Figure 2.12). Given that the variation in $\delta^{13}\text{C}_{\text{R-eco}}$ over a season is in the range of 3‰ and week to week changes can be as high as 1.5‰ , the 0.09‰ variation produced by this potential error in CO_2 concentration is relatively small.

Hand Sampling comparison to Automated Sampling

In all testing, whether the subject was isotopic information or CO_2 concentration, the precision of measurements from the automated sampler was consistently better than that of the hand sampling process (Figures 2.4 to 2.7). The standard deviation of the $\delta^{13}\text{C}$ values varied between 0.039 to 0.093 and 0.077 to 0.3649 ‰ for the automated sampling system and hand sampling process. The standard deviations of the CO_2 concentration for the automated sampling system were 0.58 to 0.94 ppm while the corresponding values for the hand sampling process were 1.94 to 8.63 ppm. Clearly, the automated sampling system is capable of producing more consistent results than the hand sampling process.

With respect to accuracy of the automated sampling system and the hand sampling system, the picture was less clear. The results indicated that for the lower CO_2 concentration samples tested, the accuracy of the hand sampling process was better for both the isotopic signature and the CO_2 concentration (Figures 2.4 and 2.5). The

difference between the hand sample process mean and the flush/fill sample mean was 0.27 ppm for CO₂ concentration and 0.18 ‰ for δ¹³C while the corresponding values for the difference between the automated sampler process mean and the flush/fill sample mean were 7.24 ppm and -0.39 ‰, respectively. For the higher absolute CO₂ concentration samples, the situation is reversed (Figures 2.6 and 2.7). For these samples, the difference between the hand sample process mean and the flush/fill sample mean was 0.03 ppm for CO₂ concentration and -1.97 ‰ for δ¹³C while the corresponding values for the difference between the automated sampler process mean and the flush/fill sample mean were 0.03 ppm and -0.05 ‰, respectively. These results are somewhat surprising and non-intuitive. It is suspected that there may have been some errors associated with either the experimental technique or the mass spectrometer analysis for this test. The expected result, given that the precision of the automated sampler is clearly much higher than that of the hand sampling system, would be that the accuracy of the automated sampler would also be higher than that of the hand sampling process. This is especially true when one realizes that the loops of the automated sampler are quite small in diameter and thus don't allow much (or any) mixing of the helium gas which pushes the sample gas into the mass spectrometer, whereas the hand sampling process, using a concentric needle apparatus, is quite likely to have helium diluting the sample. This dilution should contribute to lower accuracy of the sample gas due to the drop in output voltage of the mass spectrometer reading associated with the lower concentration, yet there is no consistent indication that the automated sampler outperforms the hand sampling process with respect to accuracy in this regard. Another possible explanation for this lower

accuracy at low concentrations is that the mass spectrometer has shown, in general, poorer performance as the CO₂ concentration drops. Again, since the output voltage of the mass spectrometer is proportional to CO₂ concentration and at concentrations below about 350ppm this voltage signal is below 500mV (too low for stable results), it could be that this experiment simply presents confirmation that 300ppm CO₂ concentrations can't be accurately measured with a mass spectrometer unless a pre-conditioning stage is employed. It is recommended that this test be repeated at the end of this field season to establish the true increases in both accuracy and precision afforded by the automated sampler over the hand sample process using higher CO₂ concentration test gases.

Dynamic Response Test

Results of this test show that the system's maximum response rate is about 0.043 Hz (or about one cycle every 23.5 seconds) with no tower tubing connected to the sampler (corresponding to a zero height measurement) (Figure 2.8). As the tower tubing length is increased, the dynamic response of the system decreases at a rate of about 0.001 Hz m⁻¹ (the cycle time increases by 0.58 seconds per meter of tubing). The data indicate that the resistance to air flow imposed by the small diameter internal passages of the components of the automated sampler coupled with the large volume of the desiccant chamber are responsible for most of the time necessary for the system to be flushed between cycles, with the relatively large diameter tower tubing contributing little to overall system air flow resistance. The apparent lack of linearity may be due to varying

resistances within each of the separate sampling tubes, various degrees of contamination on the surface of the inlet filter or kinking of the tubing.

Weekly System Integrity Tests

On a weekly basis the isotopic signatures of the field calibration gases were consistently within the confidence interval established in the laboratory, whereas CO₂ concentrations were not (Figures 2.9 and 2.10). During the course of this field season, we consistently had difficulty maintaining proper calibration of our CO₂ concentration, due in large part to lack of availability of a suitable NOAA standard gas. This situation caused the lab to rely, temporarily, on using a de facto gas standard derived from a set of ordinary, non-NOAA calibrated tanks that were calibrated using the same LI-COR 6252 that was used in the automated sampling system. These tanks were also used to fill two of the sample loops during weekly sampling. The values for the CO₂ concentration of these field calibration gases were established on a weekly basis using the IRGA (LI-6252), which was calibrated using a NOAA gas of known concentration (397.7 ± 0.1 ppm). The LI-COR 6252, when working at peak capability, has a stated variance in output of ± 1 ppm (one standard deviation, as per the LI-COR 6252 specification manual). This value shows up in the data for this experiment as the 95% confidence interval of the weekly field calibration gas readings from the mass spectrometer and is in fair agreement with those values (LI-COR 95% confidence interval ± 2 ppm, mass spectrometer weekly field calibration gas laboratory readings 95% confidence interval ± 2.75 ppm). The purpose of this test was to alert the operators and field personnel of any situations where the Valco

valve storage loops might be leaking or somehow damaged. The date DOY 249 (9/6/2005) is a good example of how these data might be used. On DOY 249, the isotopic signatures of the field calibration gases were within the 95% confidence interval of the ongoing laboratory results, but the CO₂ concentration values for the field calibration gases were significantly different from the values produce in the laboratory. This situation, if properly illuminated, should raise concerns about either individual loop integrity, the integrity of the entire automated sampling system, or the calibration of the CO₂ concentration of the mass spectrometer. For this case, there was no evidence to suggest that either the individual loops were leaking or that the automated sampling system had malfunctioned, leaving the conclusion that there were problems with the mass spectrometer calibration. An ongoing graph depicting seasonal results along with the last field results could work to alert those using the system that a problem exists. This is analogous to a control chart used on automated production lines in industry and is a standard practice at most manufacturing facilities. Although there would be some overhead costs associated with setting up and maintaining such a practice, the effort would be justified by the value of early detection of problematic situations. Unfortunately for this data set, we were not able to maintain such a system and thus relied on the mass spectrometer technician to spot such problems and to announce these issues to the group. We were not always successful in catching problems, as witnessed by the number of field calibration gas data points that are outside the 95% confidence interval for CO₂ concentration. One goal of future work should be to devise and implement an early problem detection control chart.

Discussion:

The automated sampling system designed and built for this project provided significant advantages over previous methods of sample collection. The testing described above, and the results of those tests, show that the automated sampling system provides for storage of samples in the sample loops for at least three days without significant change in $[\text{CO}_2]$ or $\delta^{13}\text{C}$. This is a clear advantage for researchers who endeavor to set up such a system and then allow it to perform its function without supervision. With the addition of a wireless interface and transmitter, the system could be made to collect data independently and report back to the experimenter that the system is ready for retrieval. If there were specific requirements of the dataset (say, a particular range of CO_2 concentrations), the system could continue to produce and discard nightly samples until this threshold had been reached. Once the threshold had been reached, the system could then alert the researcher of the data capture. The ability to store the data for up to 3 days is a significant advantage for the researcher at this point, allowing flexibility in deployment of individuals to retrieve the samples for analysis. Also key is the ability of the system to simply reset itself and take another set of samples after a sample event, which again allows flexibility in achieving the desired, predetermined data thresholds. Although the present system is not currently enabled for such operation, the addition of a wireless interface, a transmitter, and some relatively minor programming changes would allow such an operational scheme. The hand sampling method, which preceded this automated system forced, at minimum, two field technicians to be present during

sampling for long hours under adverse conditions. Since the capture of samples under the hand sampling system was subject to the particular operator's skill and judgment, significant variation could take place between data sets produced by different sets of field technicians. The automated system does not suffer this weakness.

The results of the test comparing automated sampling to hand sampling clearly show that the automated system has a much higher capability with respect to precision of the dataset than that of the hand sampling process. The standard deviation of the $\delta^{13}\text{C}$ values varied between 0.04 to 0.09 ‰ for the automated sampling system while the corresponding values for the hand sampling hand system were 0.08 to 0.36 ‰, which is at minimum a two-fold increase in precision. The corresponding values for one standard deviation of CO_2 concentration were 0.58 to 0.94 ppm for the automated sampling system and 1.94 to 8.63 for the hand sampling system, a nearly four-fold increase in precision for the automated sampling system over the hand sampling system. Shauer et al. (1997) report more precise values for CO_2 concentration (0.2ppm) but less precision for the isotope results (0.12 ‰) while Buchman et al. (1997) report the opposite (1ppm for CO_2 concentration and 0.03 ‰ for $\delta^{13}\text{C}$). Both of these authors used glass flasks for sample storage, which was not considered feasible given the objectives of our automated system. In summary, the automated sampling system was from two to four times more precise than the hand sampling system it replaced, was equal to or better than those of previous automated systems, and supported the project goals of light weight and system portability.

Although the accuracy of the datasets produced by the automated system appears to only exceed that of the hand sampling process at higher CO₂ concentrations, it is felt that this result was probably in error. There is simply no reason to suspect that the value of CO₂ concentration should affect the ability of the automated sampler to accurately reproduce the input data from either a field or laboratory test. Given that there were significant CO₂ concentration calibration issues experienced during the course of the field season, it is likely that these results have been skewed by this variation. Unfortunately, at the time of this writing the automated sampling system is in steady use during the current field season and is thus not available for follow-up testing. This follow-up test should be completed during the winter of this year when the equipment is not in use, but there is high confidence on the part of the author that the system will prove to be at least as accurate in data reproduction as that of the hand sampling process and will likely far exceed that of the performance of the hand sampler. It should be noted that, by judicious use of standard gases during the mass spectrometer analysis of the samples that the accuracy of the results for field testing can be improved by making constant offset adjustments in the field data to match the sample standards data. If it were to turn out in subsequent testing that the accuracy of the automated sampling system was not adequate, this situation could be remedied by the use of the data from the weekly field calibration gas tests to provide the proper data offsets. Precision improvements, however, are not impacted by manipulation of constant offsets and therefore can't be improved by post-analysis manipulation. For this reason, one of the most valuable advantages of the automated sampling system over the hand sampling process is the increase in precision

the automated system provides. Neither Shauer et al. (1997) or Theis et al. (2004) indicate any values for accuracy of their measurements of either $\delta^{13}\text{C}$ or CO_2 concentration, so a direct comparison of accuracy between this automated sampling system and previous efforts is not currently possible.

The dynamic response test performed in the laboratory shows the automated sampler can capture individual samples at a maximum rate of about 0.043 hertz, or about 23.5 seconds per sample. The time lag between each measurement is a function of the volume of the fluid passages in the system, the flow rate of the fluids moving through the system, and the amount of mixing which takes place between sample capture events. The hand sampling process, due to the instability of readings of CO_2 concentration produced by the LI-COR and the response time of the operators who used that reading to determine sample capture points, required that a large buffer volume be inserted in the fluid path to stabilize these readings. This large buffer volume required that much longer periods of time elapse between capturing samples (on the order of 5 minutes) to ensure that sample mixing was not taking place and reducing the quality of individual samples. Clearly, the automated system, which does not require a buffer volume, has a much higher capability with respect to sample capture rate and is therefore much better equipped to respond to environments where there is rapid fluctuation in sample content. The automated systems described by Theis et al. (2004), Shauer et al. (1997) and Buchman et al. (1997) all use large volume glass flasks (1/2 to 2 L) to store samples. Although no dynamic response was reported for any of these systems, it is highly likely that their maximum sample capture rate would be similar to the hand sampling system described thus far (about 5

minutes per sample). Although the current software program requires the automated system to sample at a maximum rate of 5 minutes per cycle, the dynamic response test results indicate that this time period could be significantly reduced, as low as 23.5 seconds per sample. This could be a significant advantage in certain environments or where rapid sampling is desired, especially if the capability of the system to store higher numbers of samples were increased in conjunction with this change. It was observed in the field that that actual rate of fluctuation of CO₂ concentration was both highly varied and extremely rapid (with changes of 0.01ppm per second not uncommon). It is suspected that these rapid changes were not based on biologically driven processes but instead were more likely to be caused by atmospheric turbulence and the associated mixing of the air stream. If it was desired to capture this turbulent change, a system capable of rapid dynamic response would be desirable.

Conclusions:

The design, construction and testing of this automated sampling system was the result of group effort aimed at increasing the quality of the data and reducing the required labor to capture that data for the HJ Andrews Airshed Project. The system was to be employed remotely in a steeply sloped watershed, requiring light weight, portability, and rugged construction. It was desired to build a system that was simple to operate to increase the range of personnel able to use the system effectively. The system needed to be able to interface effectively to the Finnigan Gas Bench II IRMS, which is rapidly becoming the new standard in gas measurement for isotopic analysis. It was also desired

to produce an automated system which could be the working model for future automated systems which would be called upon to operate virtually independently of supervision, having the capability to report back to researchers when it had collected samples of a predetermined quality. The ability of the system to capture samples in a highly dynamic environment was also sought. With regard to all of the above objectives, the automated system which resulted from these efforts performed admirably.

Future Direction:

Although the precision improvements of the automated sampling system described above over the hand sampling process being clearly demonstrated by laboratory testing, some additional testing to confirm the capability of the system with respect to dataset accuracy should be performed. It is strongly believed that the results of this testing will confirm that the automated system is as highly accurate as it is precise.

Weekly datasets produced by this system for analysis over the course of the field season included field calibration gases inserted in the field to verify the soundness of the system and to help analyze system performance. It is recommended that a control chart strategy, or a functional equivalent of the same, be implemented and tracked by both field personnel and laboratory technicians to increase the ability of system problems to be detected and rectified before too many questionable data are captured.

The system has proved capable of performing its intended function without operator intervention. The system's capacity to function independently could be improved by increasing the number of samples the system can store, removing the 5

minute settling time forced between sample capture, adding wireless communication and transmission capability, and field supply of calibration gases. Adding these components, along with the control program modifications necessary to make them function seamlessly, would extend the utility of the system to new environments.

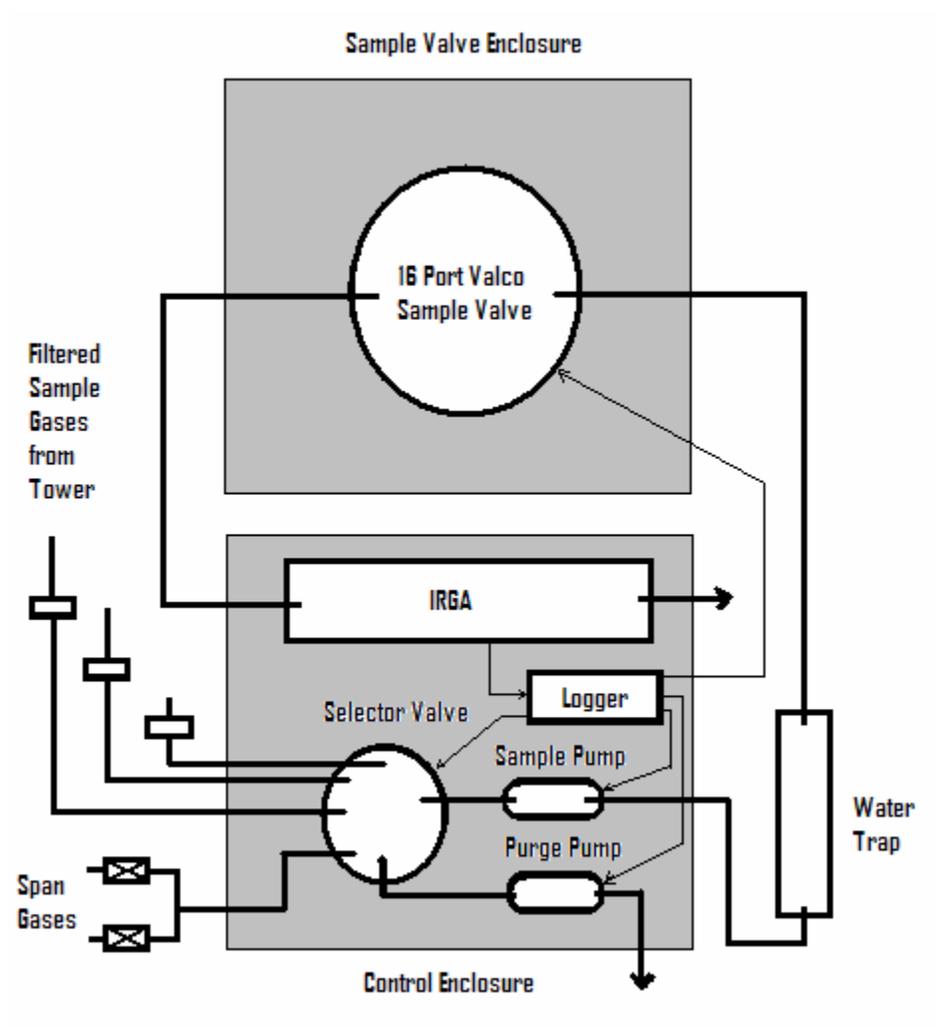
Acknowledgements and Credits:

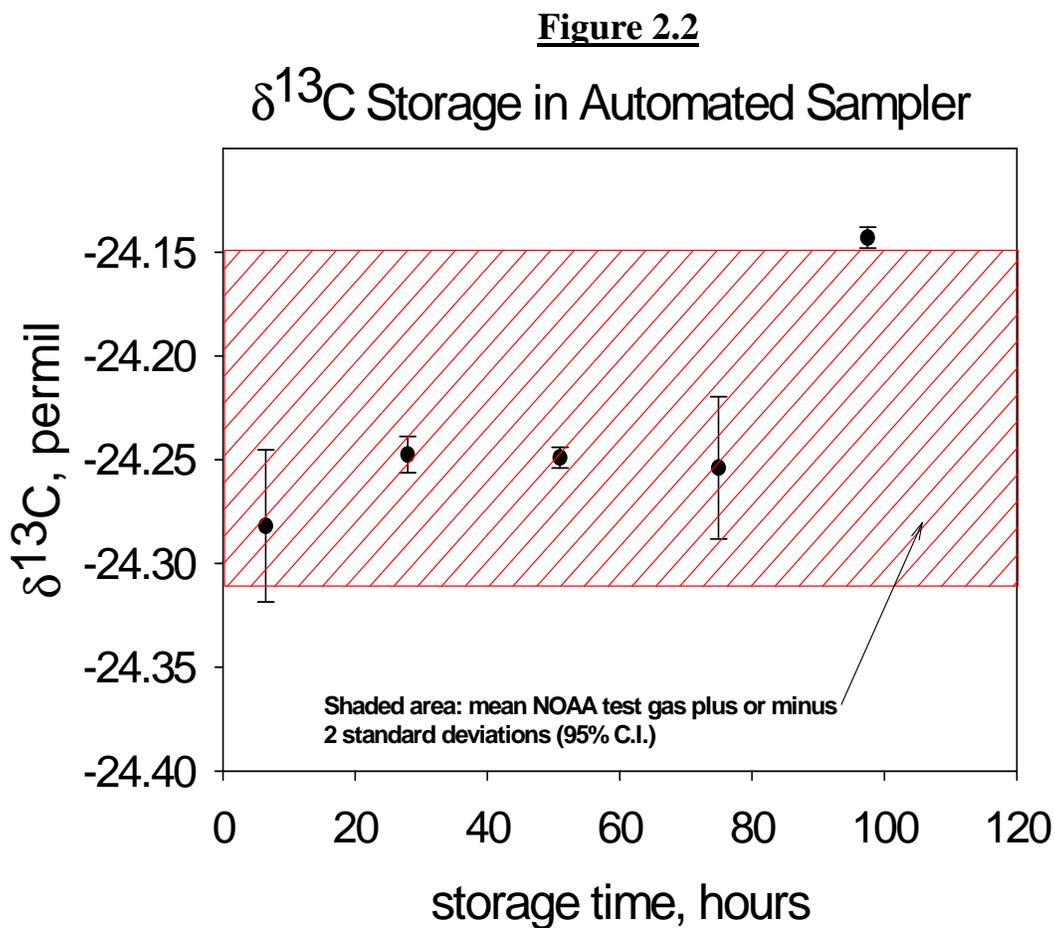
Tom Pypker and Bill Rugh were invaluable for their assistance in design, construction and testing of the automated sampler. Tom spent many hours in the field helping with both sample collection and system debugging as well as explaining various aspects of the overall project, the effort for which I am eternally grateful. Bill did the majority of the analysis of the mass spectrometer data, often working late nights and weekends to complete the analysis on time. The concept for the project was born from the work on the Pilot Project for the HJ Andrews Airshed Project under the direction of Dr. Barbara Bond, along with Co-Principle Investigators Dr.'s Elizabeth Sulzman, Michael Unsworth, and Alan Mix. General advice and assistance were given by Zachary Kayler, Holly Barnard, and Nicole Czarnomski.

Table 2.1: Autosampler Materials of Construction List

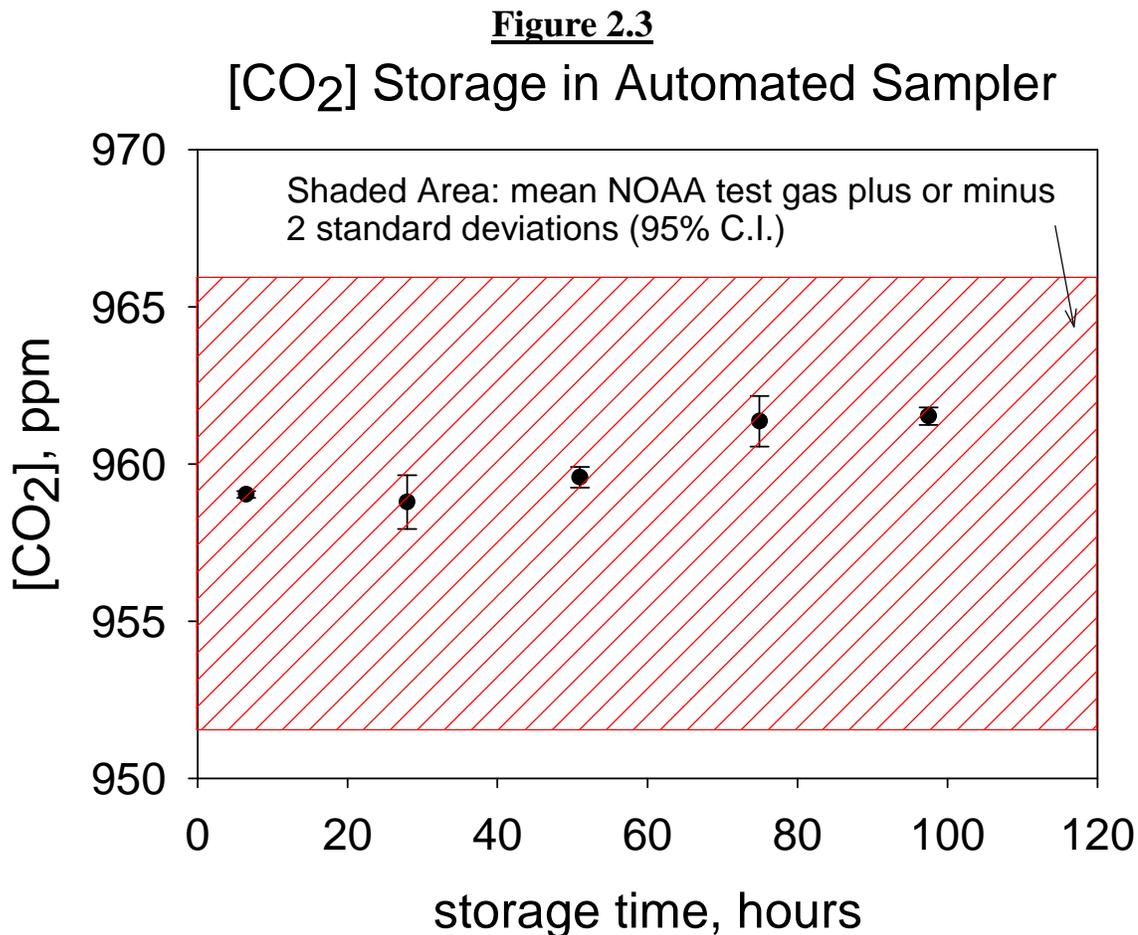
Item	Qty	Description	Manufacturer	Model
Gas Analyzer	1	Infrared Gas Analyzer; provides control signal of CO ₂ concentration to system controller	LI-COR	LI-6252
Selector Valve	1	6 port Selector Valve; provides selectable path for multiple source air samples	Valcon	E2SD6 MWE
Sample Valve	1	16 Port Sample Valve; provides sealed storage for air samples	Valcon	E2SD1 6MWE
Data Logger	1	System Controller/Data Logger; provides control signals to various system components and stores critical data	Campbell Scientific	CR 10X
Water Trap	1	155 cm ³ magnesium perchlorate cylinder which traps water to keep IRGA and Sample Valve free of contamination		
Sample Pump	1	Air pump; motive force for moving air samples through system	Neuberger, Inc	UNMP50 -KNDC, 12vdc
Purge Pump	1	Air pump; motive force for keeping tower sample lines purged	Brailsford	TD-42N, 12vdc
Solenoid Valve	2	Isolates field calibration gas cylinders from Selector Valve	Cole Parmer	625E
Inlet Filter	3	5 μm prefilters; removes fine contaminants and water drops from incoming air samples		
Bulkhead Fitting	9	Stainless steel tubing fittings; interface between inside and outside of both enclosures		Swag-Lok
Nylon Tube	Various lengths	Connects bulkhead fittings of both Enclosures and interior air sample paths in Control Enclosure		
Stainless Tube	Various Lengths	Connects Sample Valve to bulkhead fittings inside the Sample Valve Enclosure		
Solid State Relays	6	Relays between control signals and high power requirement components	Crydom	D1D07
Inverter	1	Converts 12vdc to 110vac for Valcon Valve controllers	Radio Shack	300W High Eff., Cat # 22-146

Figure 2.1: Autosampler System Diagram





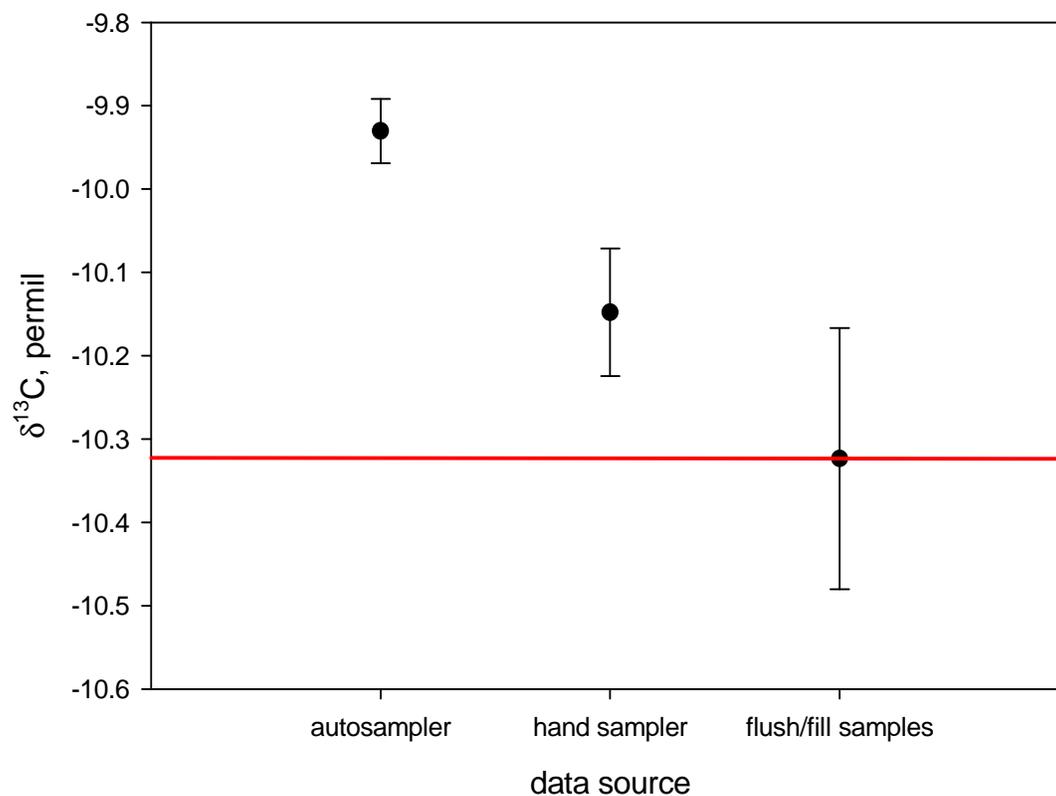
Results from a storage test performed in the laboratory. Three loops of the Valco sample storage valve (see Appendix B, Autosampler System Diagram) were filled each day for 5 days with a NOAA certified gas (nominal CO_2 concentration 958.9 ppm). After the final fill process, the contents of all of the loops were analyzed using the COAS mass spectrometer.



Results of a storage test performed in the laboratory. Three loops of the Valco sample storage valve (see Appendix B, Autosampler System Diagram) were filled each day for 5 days with a NOAA certified gas (nominal CO₂ concentration 958.9 ppm). After the final fill process, the contents of the loops were analyzed for CO₂ concentration and $\delta^{13}\text{C}$ ratio. The contents of the loops filled last represent day 0, the CO₂ concentration reported for that time period was adjusted to agree with the nominal value of the NOAA gas. All of the values for the other days were adjusted using the same offset required to bring the day 0 value to alignment with the NOAA standard.

Figure 2.4

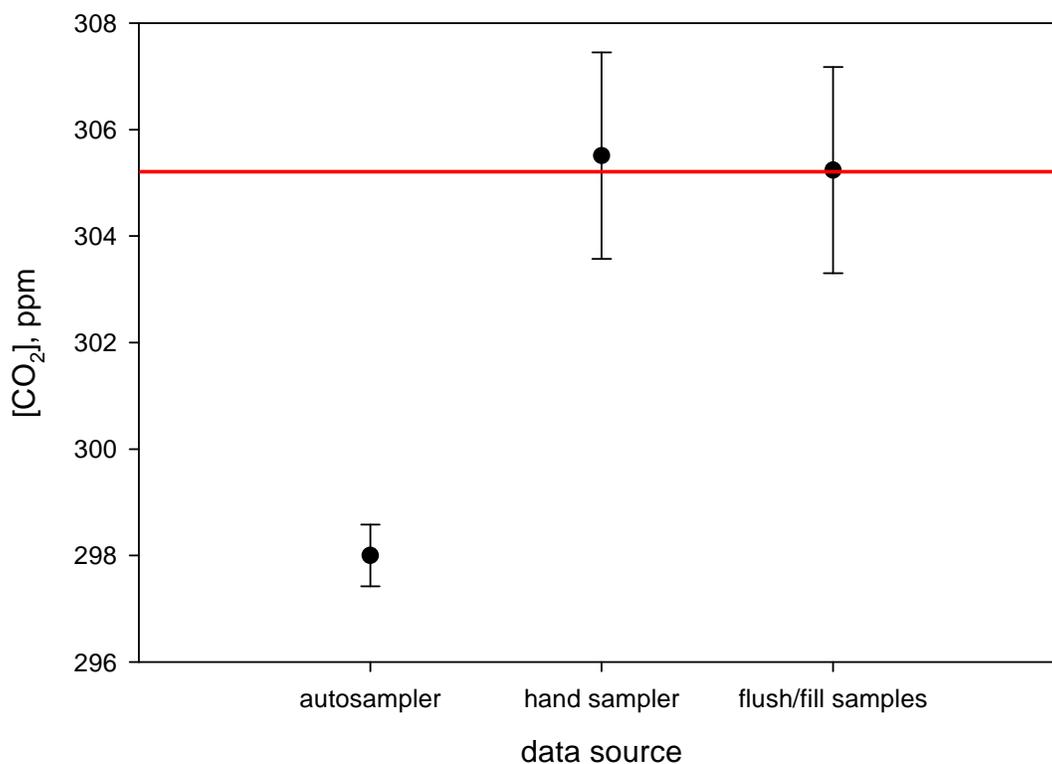
$\delta^{13}\text{C}$ measurements for low CO_2 concentration samples
June 28, 2005



Comparison of the isotopic accuracy and precision of two sampling techniques (autosampler and hand sampling). The test used a reference gas to fill both the autosampler loops and the hand sampled extainer glass vials. The “flush/fill” values represent the best technique available for determining the true value of $\delta^{13}\text{C}$ ratio of the reference gas and thus are used in this instance as a standard of comparison (“truth”).

Figure 2.5

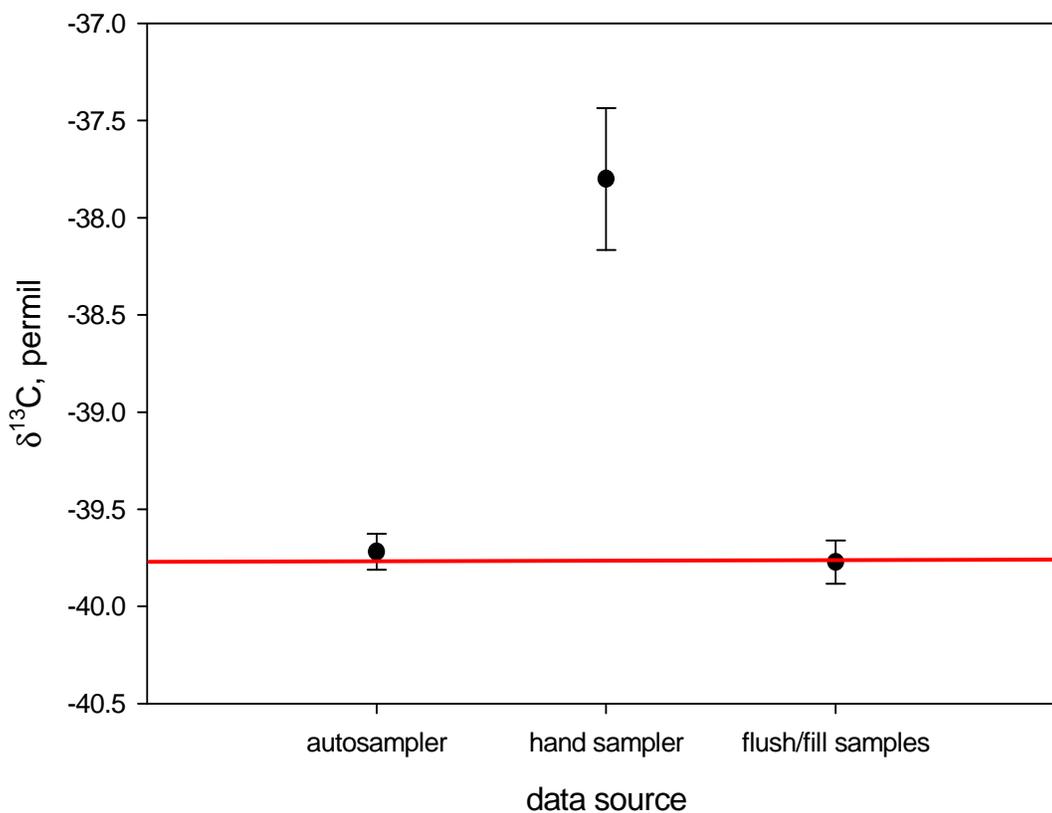
CO₂ concentration data for low CO₂ concentration samples
June 28th, 2005



Comparison of the performance of the automated sampler to hand sampling for capture of CO₂ concentration. The test uses a field calibration gas to fill both the autosampler loops and the hand sampled exetainer glass vials. The “flush/fill” values represent the best technique available for determining the true value of $\delta^{13}\text{C}$ ratio and CO₂ concentration for a given gas and thus are used in this instance as a standard of comparison (“truth”). As can be seen from the data above, the precision of the Automated Sampling System data is much better than that of the hand sampling technique, although it appears that the hand sampling output has higher accuracy.

Figure 2.6

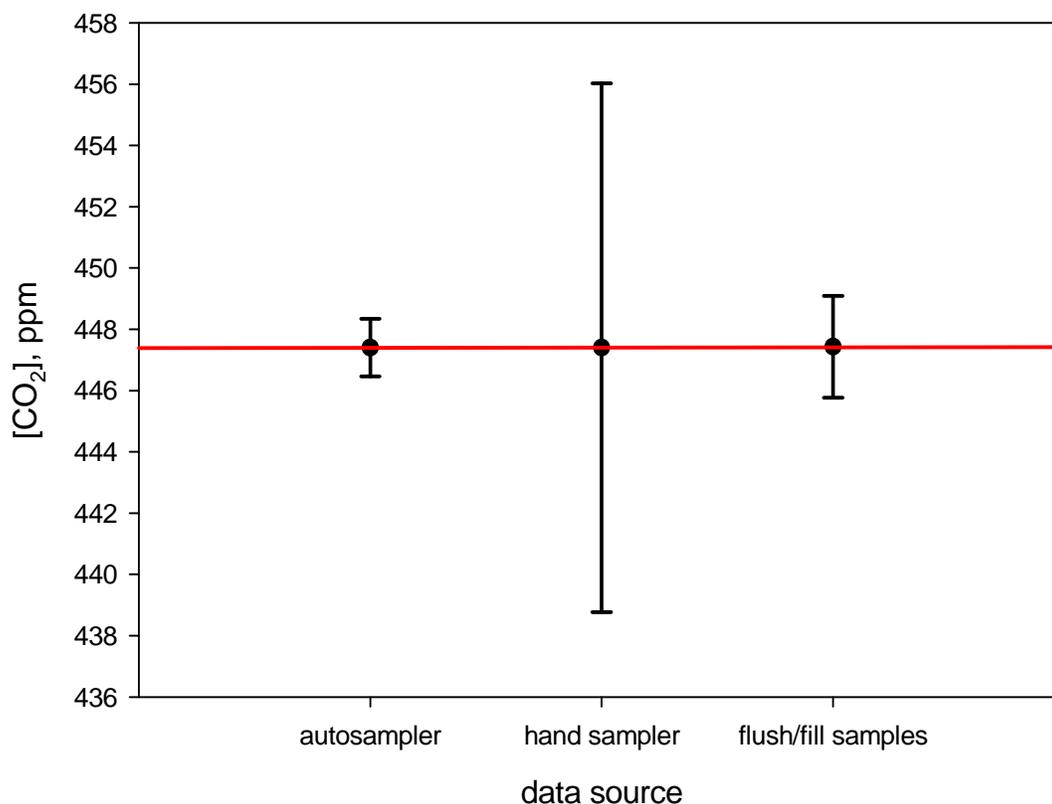
$\delta^{13}\text{C}$ measurements for high CO_2 concentration samples
June 28, 2005



Comparison of the performance of the automated sampler to hand sampling for capture of $\delta^{13}\text{C}$ ratio. The test uses a field calibration gas to fill both the autosampler loops and the hand sampled extainer glass vials. The “flush/fill” values represent the best technique available for determining the true value of $\delta^{13}\text{C}$ ratio and CO_2 concentration for a given gas and thus are used in this instance as a standard of comparison (“truth”). As can be seen from the data above, both the precision and accuracy of the Automated Sampling System data is much better than that of the hand sampling technique.

Figure 2.7

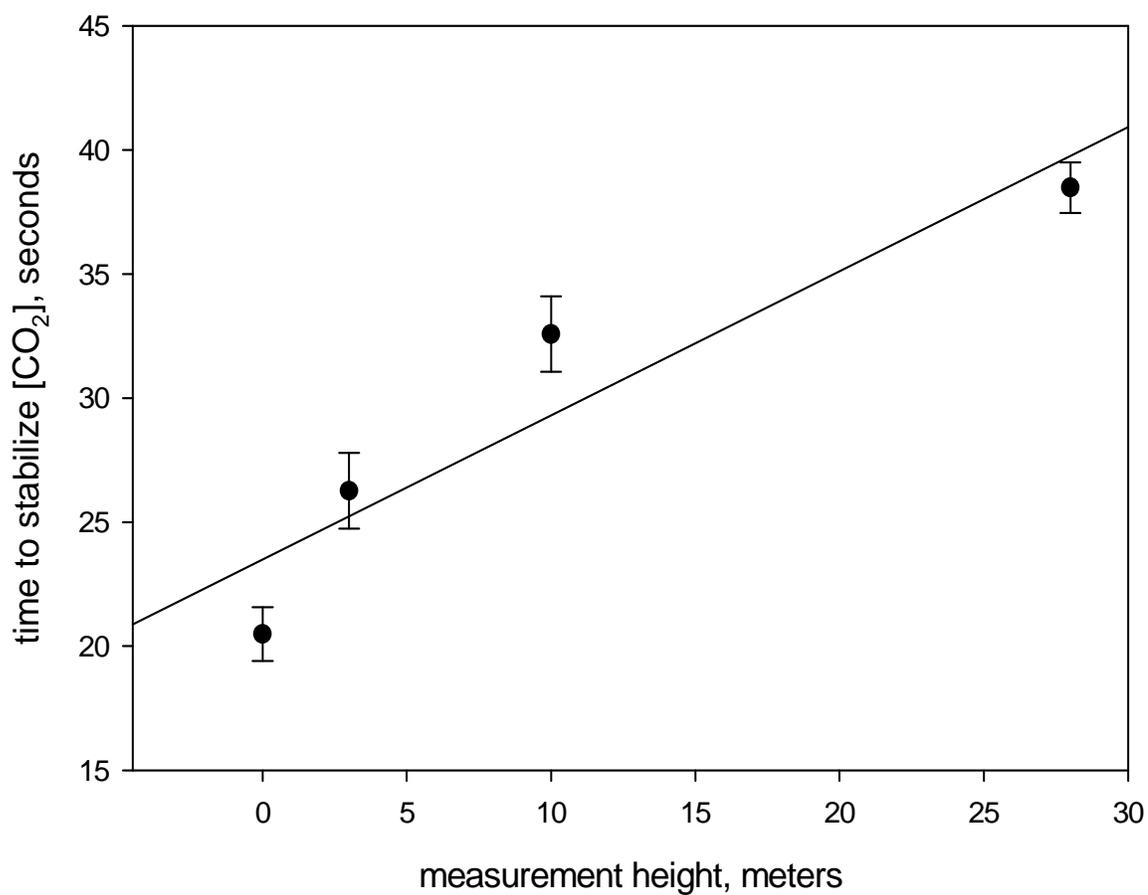
CO₂ concentration data for high CO₂ concentration samples
June 28th, 2005



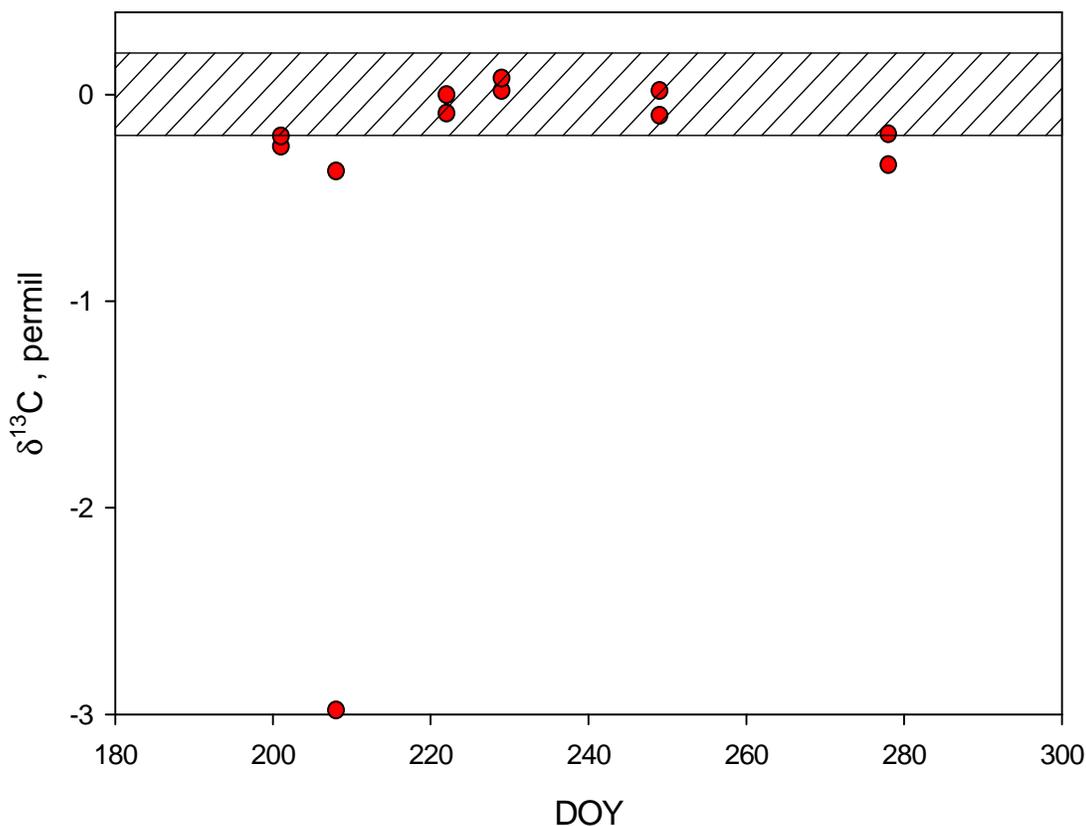
Comparison of the performance of the automated sampler to hand sampling for capture of CO₂ concentration. The test uses a field calibration gas to fill both the autosampler loops and the hand sampled exetainer glass vials. The “flush/fill” values represent the best technique available for determining the true value of $\delta^{13}\text{C}$ ratio and CO₂ concentration for a given gas and thus are used in this instance as a standard of comparison (“truth”). From the above results, for this test the accuracy of both techniques was extremely high while the precision of the automated sampling system far exceeded that of the samples taken by hand.

Figure 2.8

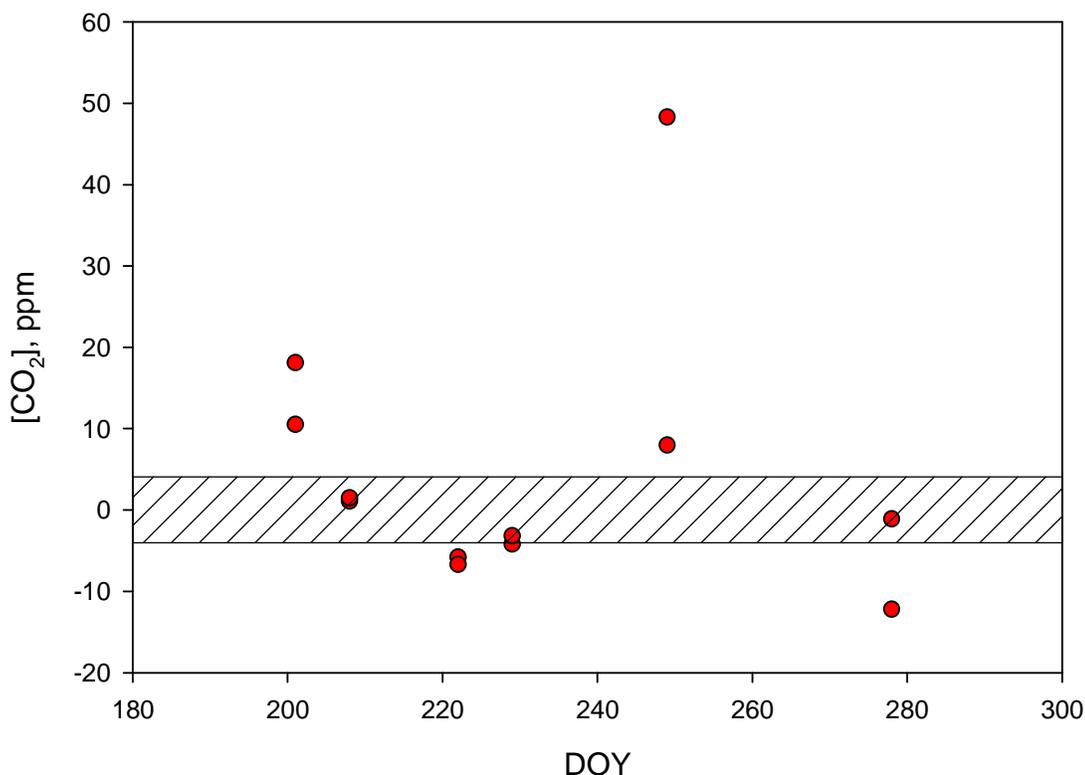
Automated Sample System Dynamic Response



Dynamic response of the automated field sampling system. The least squares regression of the data indicates that the system's maximum response rate is about 0.043 Hz (or about one cycle every 23.5 seconds) with no tower tubing connected to the sampler (corresponding to a zero height measurement). As the tower tubing length is increased, the dynamic response of the system decreases at a rate of about 0.001 Hz per meter (the cycle time increases by 0.58 seconds per meter of tubing).

Figure 2.9Weekly Span Gas Field Filled Trap Data for $\delta^{13}\text{C}$ 

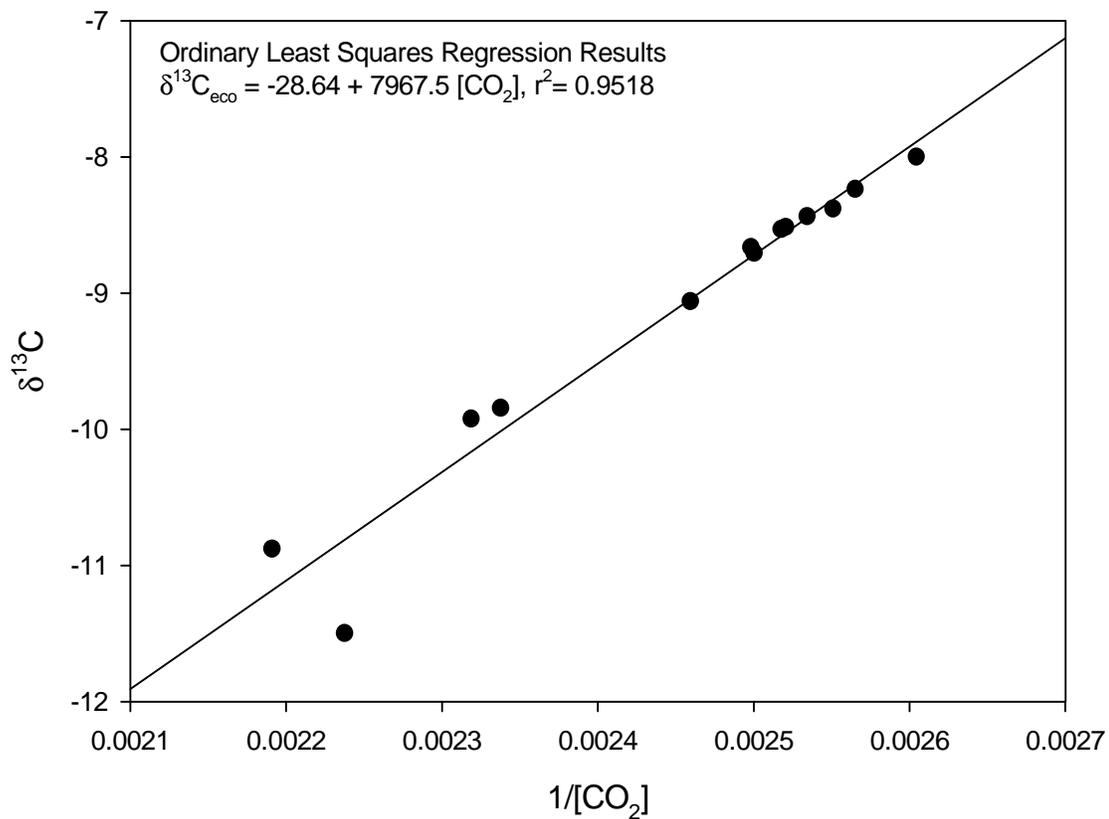
Differential $\delta^{13}\text{C}$ values of the weekly field filled plumbing integrity test. Each week two separate gases which were used to field calibrate the automated sampler IRGA were stored in two loops of the sample collection valve. After the mass spectrometer analysis of all of the field samples were completed, $\frac{1}{2}$ of the loops were refilled with each of these two gases. The results of the field filled loops were compared to the average of the lab filled loops of the same gas (the comparison is a differential value). The shaded portion of the graph represents the 95% confidence interval for the lab filled loops using the pooled standard deviation for these data throughout the season.

Figure 2.10Weekly Span Gas Field Filled Trap Data for CO₂ Concentration

Differential [CO₂] values of the weekly field filled plumbing integrity test. Each week two separate calibration gases which were used to field calibrate the automated sampler IRGA were stored in two loops of the sample collection valve. After the mass spectrometer analysis of all of the field samples were completed, ½ of the loops were refilled with each of these two calibration gases. The results of the field filled loops were compared to the average of the lab filled loops of the same calibration gas (the comparison is a differential value). The shaded portion of the graph represents the 95% confidence interval for the lab filled loops using the pooled standard deviation for these data throughout the season.

Figure 2.11

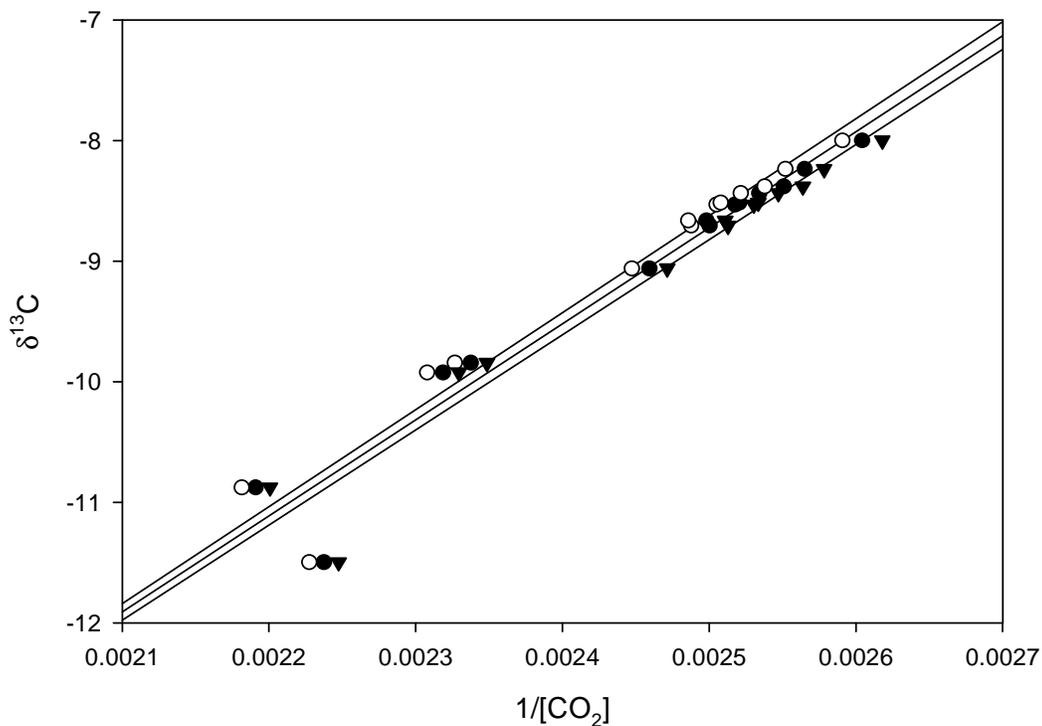
Keeling Plot of Airshed Tower Data
August 10, 2005



Typical Keeling plot. Data from August 10, 2005.

Figure 2.12

Keeling plot variation with constant offset of CO₂ concentration of +/- 2 ppm



Keeling plot variation with constant CO₂ offset. A constant offset variation of the CO₂ concentration of ± 2 ppm causes the value of δ¹³C_{R-eco} to vary from -28.73 to -28.55 ‰, with the center value of -28.64 (corresponding to 0 offset).

References:

Bowling, D.R., McDowell, N.G., Bond, B.J., Law, B.E., Ehleringer, J.R., ^{13}C content of ecosystem respiration is linked to precipitation and vapor pressure deficit, *Oecologia* (2002), 131:113-124

Buchman N, Guehl, JM, Barigah TS, Ehleringer JR, 1997, Interseasonal comparison of CO_2 concentrations, isotopic composition, and carbon dynamics in an Amazonian rainforest (French Giana), *Oecologia*, 110: 120-131

Ekblad A, Högberg P. 2001. Natural abundance of ^{13}C in CO_2 respired from forest soils reveals speed of link between tree photosynthesis and root respiration, *Oecologia* 127: 305-308

Ehleringer JR, Cook CS, 1997, Carbon and oxygen isotope ratios of ecosystem respiration along an Oregon conifer transect, preliminary observations based upon small-flask sampling, *Tree Physiology*, 18, 513-519

Fessenden, JE, Ehleringer, JR, 2003, Temporal variation in $\delta^{13}\text{C}$ of ecosystem respiration in the Pacific Northwest: links to moisture stress, *Oecologia* 136:129-136

Flanagan, LB, J.R. Brooks, G.T. Varney, S.C. Berry, J.R. Ehleringer Carbon isotope discrimination during photosynthesis and the isotope ratio of respired CO_2 in boreal forest ecosystems, *Global Biogeochemical Cycles* 96GB02345
Vol. 10 , No. 4

Flanagan, L.B., Kubien, D.S., Ehleringer, J.R., 1999, Spatial and temporal variation in the carbon and oxygen stable isotope ratio of respired CO_2 in a boreal forest ecosystem, *Tellus B*, Vol 51

Keeling CD, 1958, The concentration and isotopic abundances of atmospheric carbon dioxide in rural areas, *Geochim Cosmochim Acta.* 13: 322-334

Keeling CD, 1961, The concentration and isotopic abundance of carbon dioxide in rural and marine air, *Geochim Cosmochim Acta.* 24: 277-298

Miller JB, Tans PP, White JWC, Conway TJ, Vaughn BW, 2003, The atmospheric signal of terrestrial carbon isotopic discrimination and its implication for partitioning carbon fluxes, *Tellus* 55B, 197-206

Ometto JHB, Flanagan LB, Martinelli LA, Moreira MZ, Higuuchi N, Ehleringer JR, 2002, Carbon isotope discrimination in forest and pasture ecosystems of the Amazon Basin, Brazil, *Global Biogeochemical Cycles*, Vol. 16, No.4

Pataki DE, Ehleringer JR, Flanagan JB, Yakir D, Bowling DR, Still DJ, Buchman N, Kaplan JO, Berry JA, 2003, The application and interpretation of Keeling plots in terrestrial carbon cycle research, *Global Biogeochemical Cycles*, Vol 17, No.1

Schauer, AJ, Lai, CT, Bowling, DR, Ehleringer, JR, 2003, An automated sampler for collection of atmospheric trace gas samples for stable isotope analyses, *Ag. And For. Meteorology* (118), 113-124

Theis D, Saurer M, Blum H, Frossard E, Siegwolf TW, 2004, A portable automated system for trace gas sampling in the field and stable isotope analysis in the laboratory, *Rapid Communications in Mass Spectrometry*, Vol, 18, Issue 18, 2106-2112

Trolier M, White JWC, Tans PP, Masarie KA, Gemery PA, 1996, Monitoring the isotopic composition of atmospheric CO₂: Measurements from the NOAA Global Air Sampling Network, *J. Geophys. Res.*, Vol. 101, No. D20

Tu KP, Brooks PD, Dawson TE. 2001. Using septum-capped vials with continuous-flow isotope ratio mass spectrometric analysis of atmospheric CO₂ for Keeling plot applications. *Rapid Commun. Mass Spectrom.* 15:952–56

Zobitz JM, Keener JP, Bowling DR, 2004, Sensitivity analysis and quantification of uncertainty for isotopic mixing relationships in carbon cycle research, American Geophysical Union, Fall Meeting 2004, in press

Chapter 3

KEELING PLOT ANALYSIS AND ASSOCIATIONS WITH ENVIRONMENTAL VARIABLES

Abstract:

Variation in the isotopic composition of ecosystem respiration ($\delta^{13}\text{C}_{\text{R-eco}}$) was studied over the 2005 growing season at Watershed 1 of the HJ Andrews Experimental Forest in western Oregon with the goal of understanding relationships between environmental factors, especially atmospheric humidity and soil moisture, and $\delta^{13}\text{C}_{\text{R-eco}}$. The watershed has a pronounced nocturnal air drainage system, and concurrent atmospheric studies have revealed that the air passing the tower at night contains a well-mixed sample of respired CO_2 from the entire 96 ha basin. Samples of air were collected at night from various heights of a 37m tower erected at the base of the watershed on nine occasions between July 13, 2005 and October 5, 2005. The automatic sampler collected 15 air samples for each sampling period, and the concentration and carbon isotope composition of CO_2 in these samples was determined in the COAS laboratory at Oregon State University. $\delta^{13}\text{C}_{\text{R-eco}}$ was determined from these measurements using the Keeling Plot approach. Instrumentation affixed to the tower continuously measured microclimatic conditions during the study period; microclimate as well as soil moisture was also continuously monitored along a transect of eight plots arrayed across the watershed. Measured values of $\delta^{13}\text{C}_{\text{R-eco}}$ increased from -26.2 ‰ on 13th of July to -22.9 ‰ on 14th of September. VPD, measured by instruments affixed to the tower and averaged over 24 hour periods, ranged from 2.7 to 1758 Pa, but VPD patterns were not correlated to

seasonal $\delta^{13}\text{C}_{\text{R-eco}}$ patterns over the study period. Soil volumetric water content ranged between 7.2 % and 44.3% over the study period; the variation in soil volumetric water content over the study period was highly correlated with rain events exceeding 4 mm. Soil volumetric water content on south-facing slopes was significantly correlated ($r = -0.77$, $p < 0.01$) with seasonal $\delta^{13}\text{C}_{\text{R-eco}}$.

Introduction:

Ecosystem scientists have developed many methods and tools to study metabolic processes of forested ecosystems on flat terrain. Remote sensing (Martin et al. 1997), micrometeorology (Baldocchi 1997; Yakir et al. 2000; Baldocchi 2003), and flask sampling of air (Buchmann et al. 1998; Bowling et al. 2001) have been used successfully to uncover important ecosystem functions on flat land, but most of these techniques are inadequate when addressing ecosystem processes for forests lying on mountainous terrain. Micrometeorological techniques (eddy covariance) require a homogenous forest where horizontal fluxes of mass and momentum sum to zero over short timescales. In complex terrain, advection of air into the system via nighttime cold air drainage violates this requirement, thereby introducing unacceptable levels of uncertainty (Black et al. 1996, Lavigne et al. 1997, Staebler and Fitzjarrald 2004). Due to the inherent difficulty in studying forested systems in mountainous terrain, these systems are poorly represented in the scientific literature. As some of the most productive ecosystems in the world lie on such terrain (Harmon et al. 2004), the study of ecosystem metabolism and associated environmental conditions are among today's most important ecological problems (Schimel et al. 2002).

Convection and advection account for much of the air movement in areas with mountainous terrain. Disproportionate heating of air by radiative energy from sunlight at the ridge of a mountain during the day causes that air to rise (convection), lowering local pressure. Air from the valley flows into this low pressure zone and an up-valley wind results. During the nighttime hours, radiative cooling of the surface of the slopes in this terrain causes local air temperature to decrease, which increases that air's density. Gravity then drives this air down the slope, creating down slope winds. Air from above the valley surface moves in (advection) to replace this lost air, setting up airflow patterns that mimic stream flows. This airflow is strongest on cloudless nights, when radiative losses are highest (Monteith and Unsworth 1990). This nighttime airflow pattern in mountainous terrain has been termed "cold air drainage" (Fleagle 1950, Manins and Sawford 1979, Mahrt and Larsen 1982, Whiteman 1990).

Nocturnal cold air drainage flows entrain any gases released by the vegetation and soils on the slopes of the mountainous terrain. Of particular interest to ecosystem scientists studying the carbon cycle in mountainous regions is the capture and transport of respired CO₂ from vegetation and soils. Environmental variables such as VPD, temperature, and soil moisture have been linked to the isotopic composition of this ecosystem respiration (Bowling et al. 2002, Barbour et al. 2005, Ekblad and Hogberg 2001). The isotopic composition of respired CO₂ varies as a result of discrimination against the heavier isotope of carbon by vegetation during photosynthesis (Farquhar et al. 1989, Ehleringer et al. 1992). As a result of this discrimination, the respiration of plants and soils are depleted in ¹³C (Pataki et al. 2003), and this respiration appears to be

dominated by recently fixed carbon (Eckblad and Hogberg, 2001). By sampling the air exiting the cold air drainage at night, ecosystem scientists may be able to monitor ecosystem respiration of the watershed and use this information to study environmental effects on ecosystem processes.

The isotopic composition of ecosystem respiration ($\delta^{13}\text{C}_{\text{R-eco}}$) is typically measured using a graphical approach developed by Keeling (1958). This method requires a relatively wide range of CO_2 concentrations to ensure a low value of standard error of the Keeling intercept (Pataki et al. 2003; also see the detailed description in the Methods section of this thesis).

Almost all previous studies of $\delta^{13}\text{C}_{\text{R-eco}}$ have focused on sites with relatively flat topography. Because the CO_2 profile is usually vertically stratified at these flat sites, air samples are taken from different heights in the canopy to produce the required range of CO_2 concentrations. During cold air drainage events in the watershed investigated in our study, the canopy air tends to be well mixed, negating the advantage of sampling at multiple heights to produce the requisite large range of CO_2 concentrations. Previous studies of this watershed (Pypker et al., in press) have shown that it is possible to collect the necessary range of CO_2 concentrations from samples taken over the course of the evening rather than by sampling vertically. An advantage of sampling in mountainous terrain with this technique is that the sampling likely represents a larger area, or footprint, than the corresponding measurements on flat ground. This method is showing promise as not only a means of collecting samples that represent the state of the watershed

respiration signal, but as well a method to monitor ecosystem processes within a steeply sloped watershed.

The goal of this study was to investigate the relationships between the isotopic concentration of ecosystem respiration and two important environmental factors that affect carbon isotope discrimination by plants: VPD and soil volumetric water content. The specific objectives were to: (1) use an automated sampling system to produce weekly samples of air for isotopic and CO₂ concentration laboratory analysis; (2) determine weekly values of $\delta^{13}\text{C}_{\text{R-eco}}$ from the results of these laboratory analyses; (3) attempt to correlate this pattern of $\delta^{13}\text{C}_{\text{R-eco}}$ to the environmental variables of VPD and soil volumetric water content.

Materials and Methods:

Study Site

All measurements were made in Watershed #1 (WS1) of the HJ Andrews Experimental Forest in the western Cascades of central Oregon, USA (44.2 °N, 122.2 °W). This watershed covers 96 ha and was harvested in the late 1960s. Douglas-fir (*Psuedostuga menziessii*) dominates this closed canopy forest, with secondary components of bigleaf maple (*Acer macrophyllum*), red alder (*Alnus rubra*) and western hemlock (*Tsuga hererophylla*). This site is typical of those in the western Cascades, with wet, mild winters and warm, dry summers. The soils of this watershed are gravelly clay loams (Swanson and James 1975). A 37m tower was erected at the base of this watershed to serve as the focal point for collection of air samples and meteorological data. A

transect consisting of eight plots was installed across the watershed slightly upstream from the tower and normal to the stream, which generally follows a northwesterly path. The plots are identified by numbers; plots 501, 502, 504 and 505 are arrayed from the ridge top to the lower slope of the southwest-facing side of the watershed and plots 507, 508, 510 and 512 are arrayed from the lower slope to the ridge top of the northwest facing slope (the numbering system was established to correspond to nearby long-term vegetation plots. The discontinuity in the numbers occurs because not all of the vegetation plots have associated measurement plots). Soil volumetric water content, soil temperature, and micrometeorological variables were measured at 15 minute intervals by sensors located at these plots for correlation to seasonal values of $\delta^{13}\text{C}_{\text{R-eco}}$. This transect contained four plots on each slope face, which were distributed on each slope with roughly equal differences in elevation. For the purposes of comparison of soil volumetric water content to $\delta^{13}\text{C}_{\text{R-eco}}$, these plots were grouped by sun exposure (i.e., aspect) as well as by elevation. The sun exposure groupings were northwest-facing plots versus southwest facing plots (hereafter simplified as simply “north-facing” and “south-facing”) while the elevation groupings pitted the two highest plots on each slope against the two lowest plots on each slope. The rationale for the sun exposure groupings is that there should be differences in productivity and soil moisture content based on how much sun exposure a tract of forest receives, while the rationale for elevation groupings is that there should be differences between total water volume available for transpiration based on elevation due to gravity drainage of soil moisture. These differences in productivity and total water availability may manifest themselves in variation of vegetation respiration

signal and thus should be discernable in the analysis of $\delta^{13}\text{C}_{\text{R-eco}}$. (Note that one plot on the north-facing side of the transect (plot 508) was removed from these analyses because the data stream was not continuous over the sample period.) See Figure 3.1 for a map of the study site.

Air Samples for Keeling Plot Analysis

An automated sampling system was designed and built for this study. This system pumped air from 3 heights of the tower (3m, 10, and 28m) at the rate of about 1.5 liter min^{-1} through an infra-red gas analyzer (IRGA) (LI-6252, LI-COR Inc., Lincoln, NE, USA) to a 16 port Valco sample capture valve (E2SD16MWE, Valco Instruments Company, Inc, 7806 Bobbit, Houston, TX). See chapter 2 of this thesis for a full description of this system. Each week during the study period (May 4, 2005 to November 17, 2005), 15 air samples were captured at night (between 2000 PST and 0300 PST) by the system and brought back to the College of Oceanography and Atmospheric Sciences Stable Isotope Mass Spectrometry Facility (Oregon State University) (OSU/COAS SIMSF lab) for analysis of the CO_2 concentration and $\delta^{13}\text{C}$ of the samples. A continuous flow isotope ratio mass spectrometer (Finnigan/MAT DeltaPlus XL) coupled to a GasBench II automated headspace sampler was used to analyze the samples for both CO_2 concentration $[\text{CO}_2]$ and $\delta^{13}\text{C}$ analysis. Due to the highly negative impact on the accurate estimation isotopic ratio of ecosystem respiration ($\delta^{13}\text{C}_{\text{R-eco}}$) caused by large errors in measurement of $[\text{CO}_2]$ (Zobitz et al., 2004), it was not considered adequate to rely on LI-COR 6252 field measurements of $[\text{CO}_2]$ recorded when samples were captured. The

OSU/COAS SIMSF lab came up with a unique solution to increase the accuracy of measurement of $[\text{CO}_2]$ by using the integral of the mass spectrometer voltage output versus time signal to determine $[\text{CO}_2]$.

Keeling Plot Analysis

The value of $\delta^{13}\text{C}_{\text{R-eco}}$ was determined from each set of weekly air samples using a two end-member mixing model, or Keeling plot (Keeling 1958, 1961, Buchmann and Ehleringer 1998, Flanagan and Ehleringer 1998). The Keeling method requires plotting the $\delta^{13}\text{C}$ value for each sample against the inverse of its CO_2 concentration on a graph. The y-axis intercept of a geometric mean regression for all air samples of a particular sample period is the estimation of $\delta^{13}\text{C}_{\text{R-eco}}$ for that dataset. The geometric mean regression technique has been chosen over ordinary least squares regression because it is considered more accurate for cases in which there are errors in both the x and y variables (Pataki et al. 2003). The standard errors of $\delta^{13}\text{C}_{\text{R-eco}}$ measurements were estimated using the bootstrapping method (Chernick 1999) of standard error determination. Bootstrapping is a technique which uses the original data from a set to build a large number of datasets (in this case, 5000) by randomly re-sampling the original dataset. The Keeling plot technique was then applied to each of these re-sampled datasets and a distribution of intercepts was determined. The standard error of this distribution was then chosen as the error bounds for that particular dataset's intercept. These errors were then minimized by successively applying a jackknife algorithm (Chernik 1999), which identifies and removes the data point which has the most negative effect on the resulting standard error

and then recalculating the standard error based on the new dataset. Using the jackknife technique, potential outliers were identified. These potential outliers were then subjected to a set of criteria to determine whether or not they should be eliminated from the dataset. The final datasets were constrained to have a minimum value of CO₂ range and minimum number of regression data points (20ppm and 8 data points, respectively) to make certain that the results retained statistical significance (Pataki et al. 2003).

Vapor Pressure Deficit

The seasonal values of $\delta^{13}\text{C}_{\text{R-eco}}$ were compared to the vapor pressure deficit (VPD) of the canopy air to determine if a correlation existed between these two variables. VPD was calculated from air temperature measurements at 15 minute intervals using a curve fitted to predict saturation vapor pressure from air temperature and then multiplying the result by the value (1-Rh) (where Rh is the relative humidity of the canopy air). Air temperature and relative humidity data for the canopy air were measured using shielded probes (HMP45C, Vaisala, Helsinki, Finland) at the tower and at each of the eight transect plots. Maximum and daily average VPD were determined from these 15 minute VPD compilations.

Soil Volumetric Water Content

Continuous soil moisture measurements were made at three depths (5cm, 30cm and 100cm) at each of four locations in each of the eight transect sites using ECH₂O soil moisture probes (EC-10, Decagon Devices, Pullman WA). These instruments were

placed in service at different times during the sampling season, resulting in gaps in soil moisture data for some plots. The data from these probes were stored as 15 minute averages on a datalogger (CR 10x, Campbell Scientific Inc.). Some individual sensors malfunctioned over the course of the sampling period and their data was removed from the dataset. After the datasets were cleaned of the data from malfunctioning sensors, daily average values for all depths were determined.

VPD and Soil Volumetric Water Content to Keeling Intercept Correlation Statistics

Relationships between seasonal $\delta^{13}\text{C}_{\text{R-eco}}$ patterns with both VPD and soil volumetric water content were estimated using Pearson's product-moment correlation (Barbour et al. 2005, Bowling et al. 2002). The potential for temporal lag was also explored (Bowling et al. 2002). The lag analysis was performed by offsetting the daily averages of the independent variable (VPD or soil volumetric water content) to the dependent variable (seasonal $\delta^{13}\text{C}_{\text{R-eco}}$ pattern) for the target number of days (between 0 days and 10 days for these analyses). For these analyses, it was assumed that the dependent variable response would lag the independent variable change, so the offset days all preceded the dependent variable patterns. The significance of the correlations was determined using the method described by Helsel and Hirsch (1992). The Helsel and Hirsch method of determining correlation significance compares the calculated value t_r to the corresponding value of t from a student's t -test for the desired level of confidence in the outcome (in this case, 95%) and the number of degrees of freedom using the equation:

$$t_r = (r * (n-2)^{0.5}) / ((1-r^2)^{0.5})$$

where r is the Pearson's correlation coefficient and n is the number of data points in the data set. This equation is then solved for the critical value of the Pearson's correlation coefficient which yields a t_r equal to the desired t from the student's t -test. Values of r exceeding this critical value (or lower than the critical value, if the correlation is negative) are considered significant.

Results:

Carbon Isotopic Composition of CO₂

The values of $\delta^{13}\text{C}_{\text{R-eco}}$ over the sampling period of this study ranged from -26.83 to -22.96‰ (Figure 3.3), with standard errors ranging from 0.27 to 1.83‰. A typical Keeling plot for one of the sample days (Figure 3.2) shows a high degree of correlation ($r^2 = 0.997$) between the $\delta^{13}\text{C}$ values of the air and the inverse of the associated CO₂ measurements. The pattern of the isotopic composition of CO₂ (Figure 3.3) shows a relatively flat response to the dry summer months, rising later in the season when significant precipitation was recorded. Attempts to capture air samples earlier in 2005 (May and June) were unsuccessful because of automated sampling system failure. After these collections were made, a closer investigation of the data revealed that samples collected during the early evening were not representative of the watershed because the canopy air did not become well-mixed until later in the evening. Therefore, individual data points for these early-evening samples were removed from the datasets before the Keeling analyses were conducted.

Correlations between VPD and $\delta^3 C_{R-eco}$

VPD was calculated from temperature and relative humidity data taken at all plots along the transect (Figure 3.1) as well as at a height of 10m on the tower. Minimum and maximum values of 15-min averages of VPD were 5.0 and 1672 Pa for the plot averages, while the corresponding values for the tower were 2.7 and 1758 Pa. Patterns of VPD for all plots were well correlated (Pearson correlation ranged from 0.98 to 1.0) with one another, indicating that the average VPD of all plots accurately represented the variation in VPD for the entire transect (Figure 3.4). A correlation of this average plot VPD with that of the VPD measured at the tower also yields a high correlation (Pearson correlation = 0.99), indicating that tower VPD can be used as accurate proxy for the VPD of the transect (Figure 3.5). For simplicity in analysis, and justified by the high degree of correlation between plots and plot average and Tower, only the Tower VPD values were used to correlate to Keeling intercepts. This correlation of Keeling intercepts to Tower VPD for 0 to 10 day lags yielded Pearson correlation coefficients between -0.49 and 0.11, which were not significantly different from 0 using the Helsel and Hirsch significance criteria ($r = \pm 0.58$) (Figure 3.6).

Soil Volumetric Water Content Correlation

Average soil volumetric water content from the 30cm depth ranged from 7.2 % at plot 507 on September 12, 2005, to 44.3% at plot 501 on November 15, 2005. The 30cm depth was chosen for correlation to Keeling intercepts due to its dominance as the source of water for forests composed of predominantly Douglas-fir (Brooks et al. 2006). A

comparison of plots by elevation showed a high correlation between soil volumetric water content for upslope versus downslope plots ($r = 0.90$), with the higher elevation plots (upslope) having an average of $6.3 \pm 1.2\%$ (1 s.d.) higher soil volume than that of the lower elevation (downslope) plots (Figure 3.7). A comparison of plots grouped by sun exposure did not show a strong soil volumetric water content correlation ($r = 0.58$) (Figure 3.8), with south-facing plots having an average of $7.8 \pm 3.7\%$ (1 s.d.) higher soil volumetric water content than that of the north-facing plots. Both of these groupings showed a strong correlation, with virtually no lag, to significant rainfall events (greater than approximately 4mm, Pypker et al. 2005).

Soil volumetric water content on plots grouped by elevation was not significantly correlated to Keeling intercepts (Helsel and Hirsch, 1992). The values of correlation ranged from -0.09 to -0.46 for upslope plots and -0.15 to -0.47 for downslope plots (Figure 3.9). Soil volumetric water content on plots grouped by sun exposure showed a significant correlation between Keeling intercepts and south-facing plot soil moisture ranging from -0.77 to -0.69, but there was no significant correlation for north-facing plots ($r = -0.08$ to 0.29). It was expected that the isotopic signature of the ecosystem respiration would lag behind the soil moisture changes, but the data showed no significant lag response (Figure 3.10). The lagged response of ecosystem respiration is thought to be a consequence of the fact that about 70% of ecosystem respiration comes from soil microbial respiration (Eckblad and Hogberg, 2001). The sugars which supply fuel for these microbes are supplied by vegetation via phloem transport. There is a delay between when the sugars are fixed in the leaves and when they are liberated from the soil

and become part of the ecosystem respiration signal because of: (1) the phloem transport time; (2) the time it takes sugars to diffuse into the soil and reach the microbes; (3) the time it takes for the microbes to digest and respire these sugars; and (4) the time it takes the respiration to diffuse from the microbial source to the soil surface. All of these delay times add in sequence to produce a lag between the soil moisture signal which produces the change in isotopic composition of the sugars produced by assimilation due to stomatal modulation and the time it takes this signal to reach the ecosystem as soil respiration.

Discussion:

The seasonal pattern of isotopic composition of ecosystem respiration observed in this study is consistent with the patterns observed in other studies for forests of similar vegetative compositions and climates (Bowling et al. 2002, Fessenden and Ehleringer 2003). During the initial stages of the growing season, while the soils of the forest are still relatively moist, vegetation has adequate access to water. Accordingly, the vegetation will respond by maintaining high levels of stomatal conductance to increase their rate of carbon uptake. The ratio of carbon dioxide concentration in the leaf interior to that of the atmosphere (c_i/c_a) remains relatively high in this situation, which promotes discrimination against the heavier isotope of carbon (Farquhar et al. 1989). As the summer progresses and precipitation drops, soils begin to dry out, limiting vegetation's access to water. In order to avoid cavitation and the associated damage to xylem cells, trees begin to close their stomata, lowering stomatal conductance. This drop in conductance decreases the c_i/c_a ratio, which reduces discrimination against the heavier

carbon isotope. As a result of this mechanism, as the season progresses and access to moisture is reduced, the newly-fixed sugars should show an increase in the ratio of ^{13}C to ^{12}C . Recent studies have shown that a large amount of recently-fixed carbon may cycle quickly through an ecosystem (Eckblad and Hogberg, 2001), with the result that the isotopic composition of respired CO_2 may be heavily influenced by recently-fixed sugars. Thus, the pattern of change in isotopic composition of recently fixed sugars is visible in the respiration signal, as observed in this study. In the early part of the summer, values of $\delta^{13}\text{C}_{\text{R-eco}}$ were relatively depleted. As the soils began to dry out, $\delta^{13}\text{C}_{\text{R-eco}}$ became enriched.

If the only component of ecosystem respiration were to come directly from the use by vegetation of recently fixed sugars, the respiration isotopic composition signal could be used as a direct proxy for the state of vegetative drought stress. However, respiration from soils and vegetation may be fueled by stored carbon, such as starches or other complex biochemicals, as well as newly-fixed carbon, and this mixed source may confound the signal (Mortazavi et al. 2002, Buchman 1997). In soils, the partitioning of respiration between heterotrophic and autotrophic components is a topic of intense study (Chambers et al. 2004, Rochette et al. 1999, Bhupinderpal et al. 2003) and is a topic of interest in the overall HJ Andrews Airshed project. The partitioning of respiration signals is beyond the scope of this thesis, however.

Many studies have been aimed at connecting the seasonal pattern of Keeling intercepts with drought stress variables such as VPD (Bowling et al. 2002, Sanford et al. 1986, Condon et al. 1992), soil volumetric water content (Fessenden et al. 2002) and

precipitation (Ometto et al. 2002, Pataki et al. 2003, Bowling et al. 2002). The results of this study indicate that VPD was not significantly correlated with seasonal Keeling intercepts, unlike results from many previous studies (Bowling et al. 2002). The correlations between soil volumetric water content of south-facing plots and seasonal Keeling intercepts were strong, but there were no significant correlations between the isotopic values and soil volumetric water content of north-facing slopes. When soil volumetric water content was grouped by elevation, no significant correlations were found. Past research has found lags of between one and four days between VPD and $\delta^{13}\text{C}$ of CO_2 respired by roots (Eckblad and Hogberg 2001) and five to ten days between VPD and ecosystem respiration (Bowling et al. 2002), yet the data from this study do not show significant lags between either VPD or soil volumetric water content and seasonal Keeling plot intercepts.

The lack of correlation between VPD and seasonal Keeling intercepts may lie in the timing of the sampling period. Although attempts were made to procure air samples for respiration analysis in May and June of 2005, these data points were rejected due either to equipment malfunction or lack of sufficient CO_2 range. The physiological connection between soil volumetric water content and VPD with respect to isotope discrimination is rooted in how those variables affect stomatal function. During the early part of the growing season, soil volumetric water content is relatively high, so that vegetation's access to a pool of water for transpiration is plentiful. During this period of time, isotopic discrimination will be dictated by the rise and fall of VPD. As the growing season progresses and soil moisture begins to decline due to the cumulative effects of

transpiration, vegetation's access to water begins to diminish. At this time, leaf water potentials must also fall in order to provide the "lift" necessary to extract the water from the soil for transpiration. At some point in this continuum of soil volumetric water content depletion, the ability of trees to extract water from the soil is becomes dominated by soil water availability and not by VPD. This point comes later in the growing season, after significant dry down of the soil. I believe that the data set for this study does not begin early enough to include the period of time for which VPD is the dominant force in dictating stomatal closure and associated carbon isotope discrimination. Had Keeling plot intercepts been available for earlier periods in the growing season, correlations between Keeling intercepts and VPD would likely have been much greater. Sapflow data from 2000 (Moore et al. 2004) support this hypothesis, indicating that changes in sapflow were well coupled to changes in VPD up to about July 19th, after which the those variables were poorly correlated. Sapflow changes again appeared to be well coupled to changes in VPD beyond about September 7th, coincident with the onset of significant rainfall in that year. Since sapflow can be considered a proxy for transpiration, this sapflow data indicates that transpiration was only coupled to VPD during the early growing season (when soil moisture content is relatively high) and late growing season (after the rains started, again coincident with high soil moisture content). This pattern is consistent with the notion that when soils have plenty of water, VPD dominates transpiration.

The high degree of correlation between soil moisture patterns for upper and lower elevation plots ($r = 0.90$) implies that if there were a correlation between upper plot soil volumetric water content and seasonal Keeling intercept pattern, then there should also be

a correlation between the lower plot soil volumetric water content and the seasonal Keeling intercept pattern. The reverse is also true: if the lower elevations did not show a correlation, then neither should the upper elevations. This was indeed the case- no significant correlation was found for either upper or lower elevation plots and seasonal Keeling intercepts. It is worth noting that grouping all soil moisture plots together for the entire transect also produces no significant correlation with seasonal Keeling intercepts. Indeed, the pattern of correlation over time lag was quite similar for upper elevation, lower elevation, and all plots combined.

The correlation Keeling intercepts and the soil water content of south-facing plots was significant, whereas the correlation for north-facing plots was not. It has been shown by Pypker et al. (2006) that the nighttime air drainage from which these samples were taken is well mixed, so stratification or preferential sampling of one side of the watershed is not likely, yet the ecosystem respiration signal appears to be correlated with the south-facing slope soil volumetric water content. I propose three hypotheses to explain this finding. The first hypothesis pertains to airflow patterns. It is possible that air may enter the watershed predominantly from the south face, flooding the basin with air from the south side. If this were the case, then the ecosystem respiration signal from the north-facing slope may not be making it down to the site of the tower and thus was not represented in the air sampling conducted at the base of the watershed. However, this hypothesis is not considered likely. The theory behind cold air drainage is that cooling at the ground level causes surface air to increase in density, which then flows downhill due to gravity, where it collects in the bottom of the watershed. Continued cooling of the

slopes above this pool of air in the bottom of the watershed causes further cold air drainage off of the slope, pushing the collected air in the bottom out of the base of watershed. The tower at the base of the watershed lies in this drainage path. Since the flow of air is turbulent (Pypker et al. in press), the air in this drainage is well mixed. The north slope of the watershed is shaded earlier in the day than the south-facing slope and therefore should begin draining earlier. If there were to be an uneven representation of one slope's respiration signal over the other, this geographic layout should favor the north slope, not the south slope.

The second hypothesis is that the volumetric water content of the south-facing slope represents the overall state of the entire watershed better than that of the north-facing slope. If this were true, then the respiration contributions of both slopes of the watershed could be represented in the dataset, yet the correlation between $\delta^{13}\text{C}_{\text{R-cco}}$ and north-facing slope soil moisture would be poor. Essentially, the soil volumetric water content of the thin line of the transect is being used as a proxy for the entire surface area of the watershed (or at least of the portion of the watershed contributing respiration to the drainage, also termed the "footprint" of the watershed). It is impossible to say, given current the dataset, whether or not the conditions of the plots on the south-facing slope better represent the entire watershed "footprint" than the plots on the north-facing slope. To test this hypothesis, a sampling of multiple locations in the watershed would need to be compared to corresponding transect data to see determine what best represents the entire watershed's soil moisture signal. Such an endeavor is beyond the scope of this paper.

The third hypothesis is that the north-facing slope of the watershed had significantly lower respiration rates than the south-facing side due to greater moisture stress (Zavitkowski et al. 1968; Burton et al. 1998). As a result of this reduction in respiration rate, the north-facing slope would contribute less to the total amount of respired CO₂ captured in the samples collected at the tower. The patterns for soil volumetric water content for the north-facing slope versus the south-facing slope lend support to this argument. During the early part of the sampling period, the north-facing slope soil moisture values were consistently nearly 12% lower than those of the south-facing slope. Average soil depth (measured to the bottom of the B horizon) on the south-facing slope was about 105cm, whereas the equivalent measurement for the north slope is 81cm (Padilla, private communication). Assuming that both slopes started the year with similar water content (on a percentage basis), the larger volume of soil on the south-facing slope would provide a larger pool of water for the vegetation on that slope to draw from, compared to the north side with less soil volume. Assuming the vegetation on both slopes would begin the growing season with ample soil water available and assuming roughly equal distributions and size classes of vegetation, the slope with the smaller pool of water available should reach drought stress conditions ahead of the slope with the larger pool. The smaller soil volume of the north-facing slope and the lower soil volumetric water content of that slope at the beginning of the sampling season lend support to this hypothesis. Unfortunately, data for the early part of the growing season (May and June 2005) was not available for this study. It may have been possible to extract the data necessary to confirm this hypothesis by looking for declining values of

CO₂ concentration as drought stress on the north-facing slope developed. As the north-facing slope was heading into drought stress during this timeframe, the isotopic composition of the north-facing slope air drainage would have reflected this transition. Had this early data been available, this signal should have also been detectable, which may have provided confirming evidence of this hypothesis. Streamflow data for WS-1 from the period 1997-2002 (Jones and Post, 2004) lend support to this hypothesis: continuously declining streamflow starting in May and continuing until September indicate that the watershed was drying out during this period. The loss of stream flow in the dry months would be indicative of a general decline in soil moisture content, which supports the hypothesis that respiration declined on the north-facing slope due to loss of soil moisture and associated drought stress.

Other data that lend more direct support to the third hypothesis. For the current year the value of foliar respiration for the south-facing plots averaged $-7.8 \mu\text{molm}^{-2}\text{s}^{-1}$ whereas for the north-facing plots the corresponding value was $-4.3 \mu\text{molm}^{-2}\text{s}^{-1}$ (Bond, private communication) (figure 3.11). Soil respiration values for south-facing and north-facing were -7.0 and $-3.0 \mu\text{molm}^{-2}\text{s}^{-1}$, respectively (Kayler, private communication) (figure 3.11). With respect to both foliar and soil respiration, the south-facing slope respired at nearly twice the rate of the north-facing slope, again reinforcing the notion that the respiration signal seen at the base of the watershed during the sample period was dominated by respiration from the south-facing slopes.

Why would soil accumulation be lower in the north-facing slope? South-facing slopes generally dry out earlier in the year than north-facing slopes, limiting vegetative

carbon assimilation and detritus production. As a general rule, this should favor soil development and vegetation growth for north-facing slopes. However, differences in the topography of individual sites may alter the outcome of these generalized mechanisms. The average incline of the north-facing slope at the study site is about 93% whereas that of the south slope is about 65% (Padilla, private communication). It is possible that this difference in slope creates soil instability on the north-facing slopes, promoting slides and soil movement towards the base of the watershed and is thus responsible for the north-facing slope having less soil volume than the south-facing slope. There is a great deal of blowdown and apparent soil instability on the north-facing slope, which also supports this position.

Conclusions:

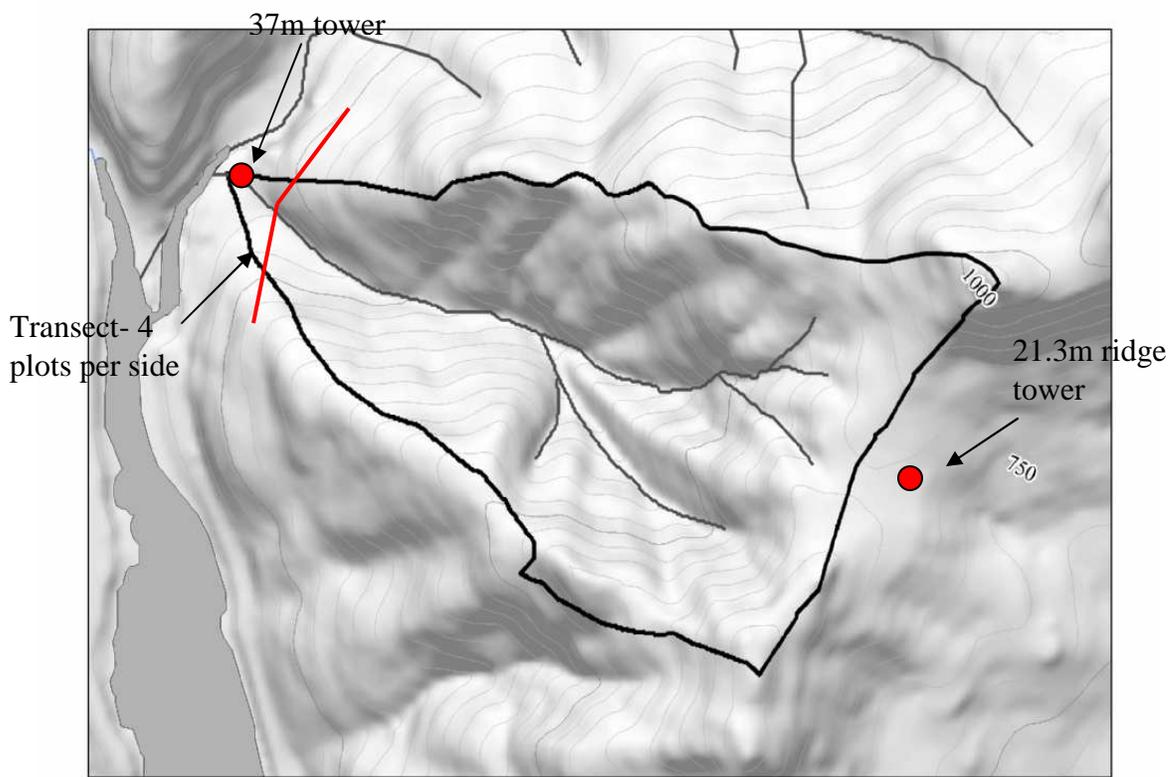
The pattern of isotopic composition of ecosystem respiration at WS-1 of the HJ Andrews Experimental Forest from July to September 2005 is consistent with expectations of a forest in the midst of drought stress. Contrary to results of many other researchers, the seasonal Keeling intercept pattern did not correlate significantly to the pattern of VPD during the sample period, but did correlate significantly to the patterns of soil volumetric water content on the south-facing slope of the watershed. The lack of correlation with VPD is likely due to the fact that the sampling period did not include the early part of the growing season, where soil moisture is plentiful and VPD dominates stomatal modulation. The correlation of Keeling intercept pattern with south-facing soil moisture is likely due to the fact that soil volumetric water content on that face was

sufficient to allow the vegetation to continue to assimilate carbon and maintain higher respiration rates in the later part of the season.

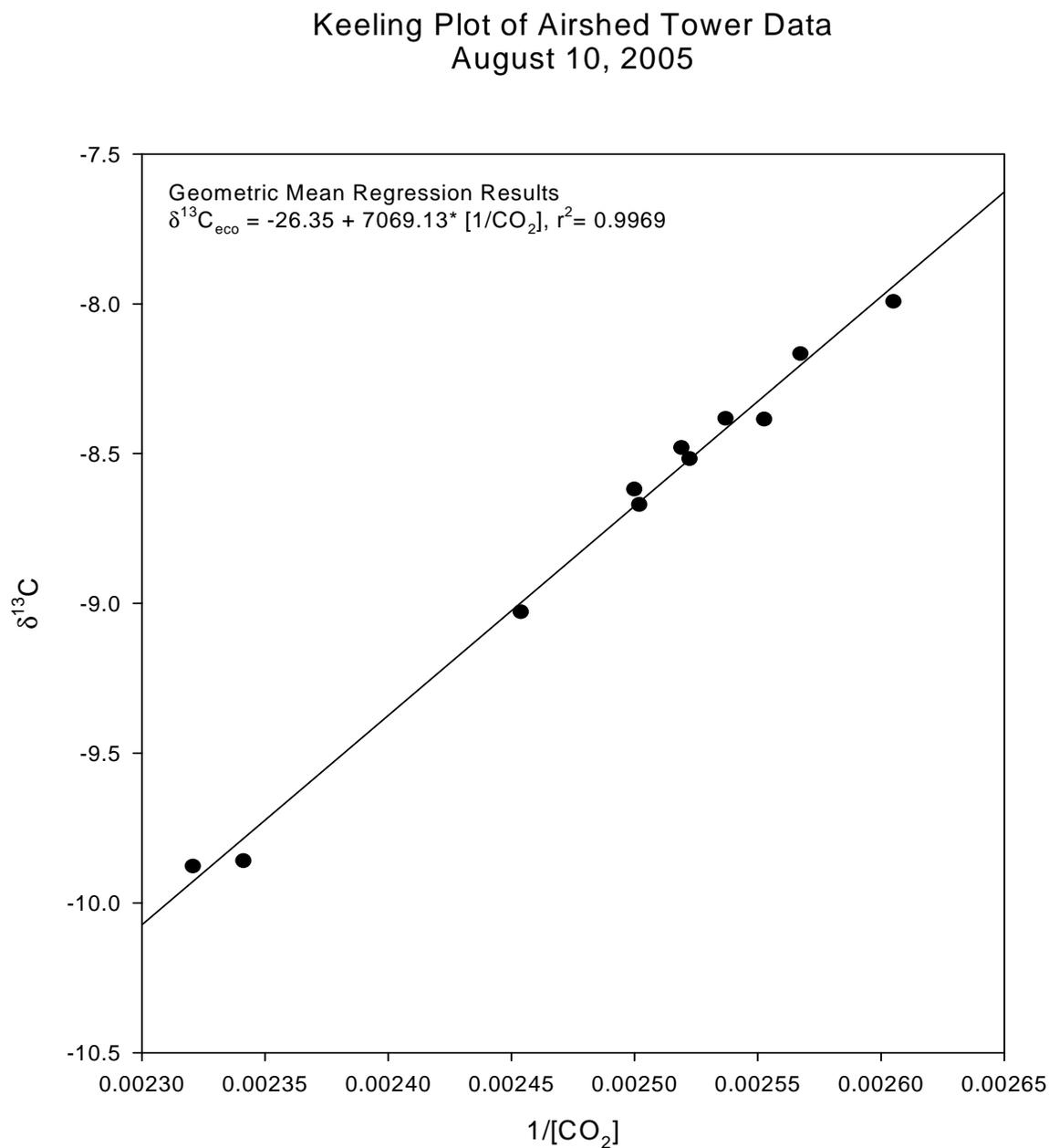
Acknowledgements and Credits:

Tom Pypker was immensely helpful in all aspects of this project, from design and construction of the automated sampler to data acquisition and manuscript editing. This project literally could not have been completed without his help. Advice and coaching on field work techniques and assistance in data acquisition from Zach Kayler, Holly Barnard, Claire Phillips, Julian Licata and Dave Conklin were greatly appreciated. Zac Kayler, Aidan Padilla, and Barb Bond provided data to support hypotheses on soil depth. My advisers Barb Bond, Elizabeth Sulzman, and Julia Jones provided education, encouragement, and advice at critical junctures and were of great aid in the production of this thesis.

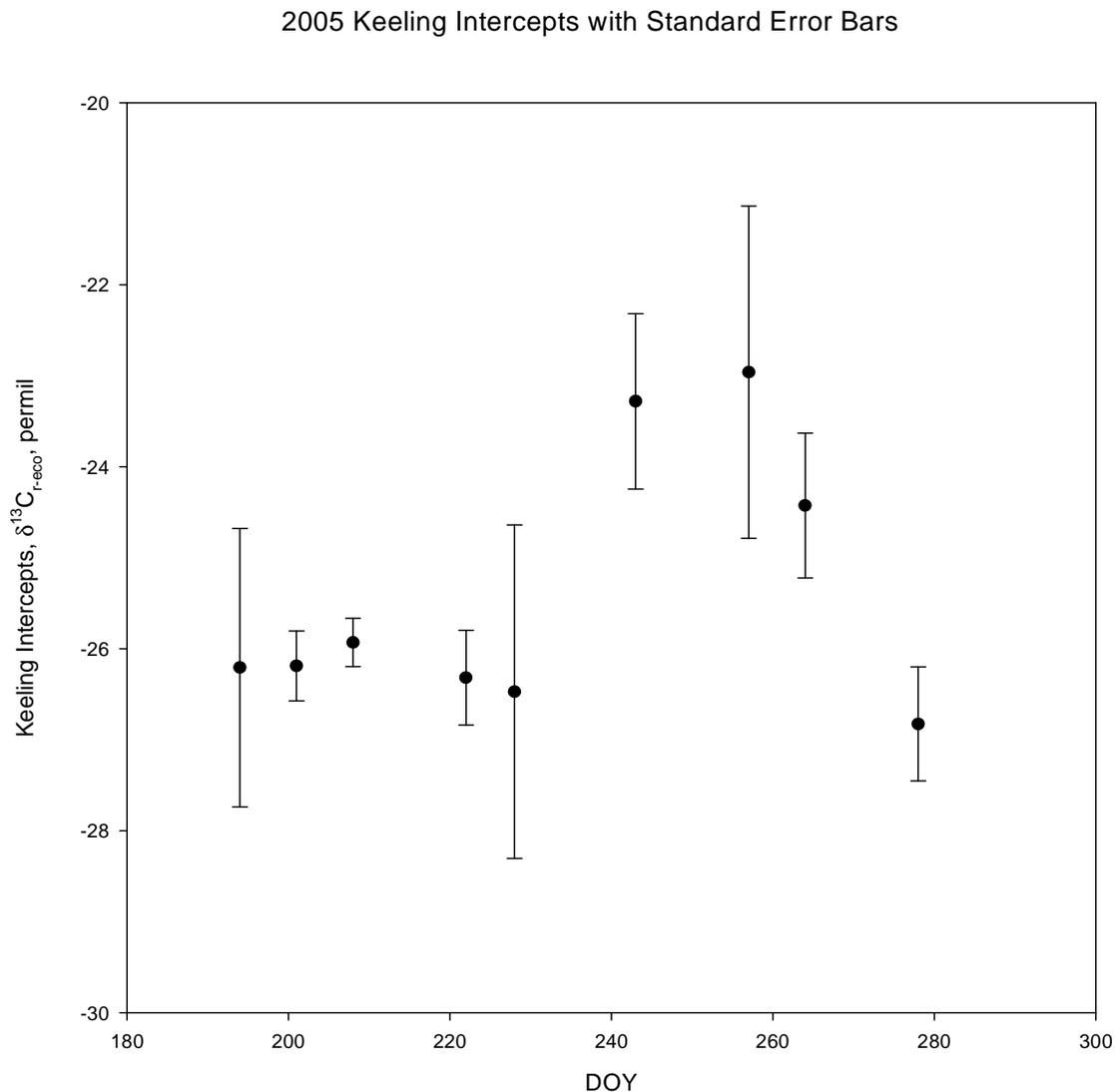
Figure 3.1: HJ Andrews Watershed 1 Topographic Map



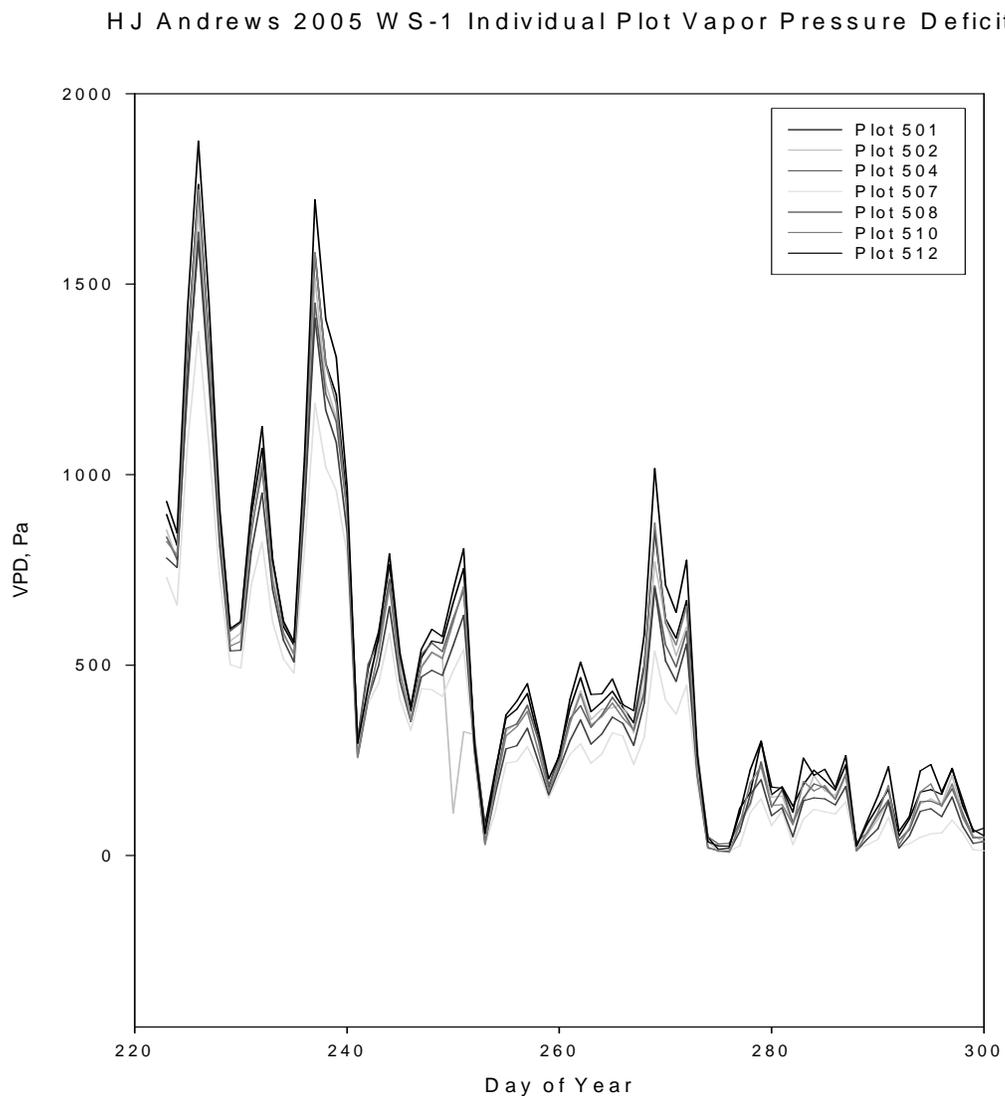
Topographic map of Watershed 1 at HJ Andrews Experimental Forest. The 37m tower at the base of the watershed and the 8 plots which make up the transect serve as the data collection points for the isotopic composition of ecosystem respiration, VPD, and soil moisture used as variables in this study. The 21.3m ridge tower has been used in subsequent studies to determine values of wind speed, wind direction, and isotopic composition of the air above the canopy.

Figure 3.2: Typical Keeling Plot

Typical Keeling plot for HJ Andrews Airshed project. This data is from August 10, 2005. Geometric mean regression analysis of the data for this date resulted in an intercept of -26.35 ‰ .

Figure 3.3: Seasonal graph of Keeling intercepts

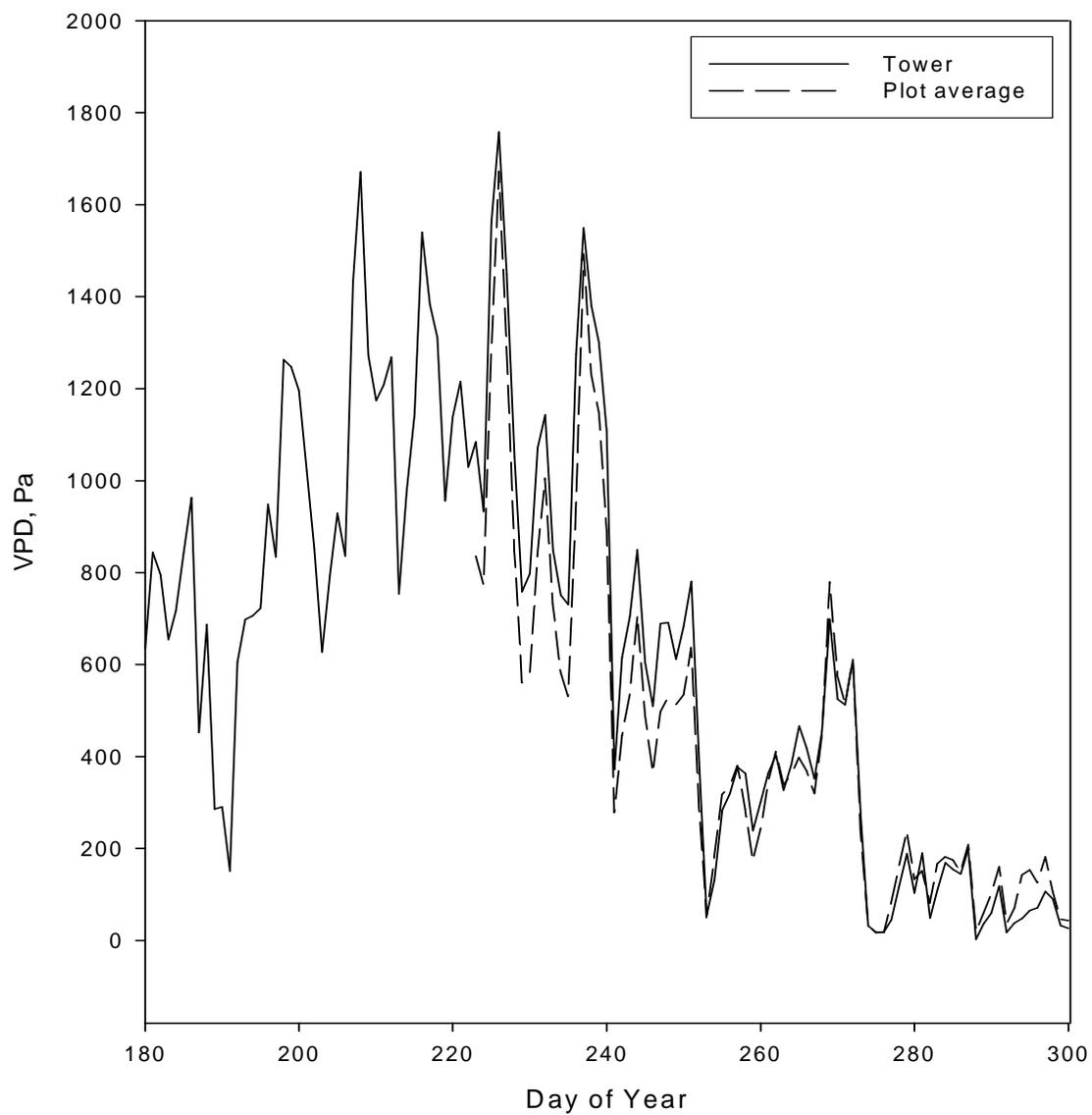
Keeling plot intercept graph for 2005 data from WS-1 of HJ Andrews Experimental Forest. DOY's 315 and 236 were removed from the dataset because the air sample data was collected before the airshed drainage became well mixed. DOY 249 was removed because the data was corrupted by a malfunction of the collection storage valve. DOY 315 was removed because the CO₂ range was too small to support a statistically valid intercept. The first trap sample was removed from each dataset because the sample timing was premature. For the 2005 growing season, the intercepts ranged between -26.83 and -22.960/00. Error bars represent +/- 1 standard error.

Figure 3.4: Seasonal values of VPD for all plots on transect

2005 VPDs for individual plots across transect in WS-1 of HJ Andrews Experimental Forest. Data for temperature and relative humidity at each plot were used to determine VPD. Note that the pattern for each plot is virtually identical, indicating that for the purposes of establishing correlations with isotopic composition of ecosystem respiration, any one plot could be used in place of the others without a change in the correlation statistics (Pearson correlation ranged from 0.98 to 1.0).

Figure 3.5: Comparison of Tower VPD to Plot Average VPD

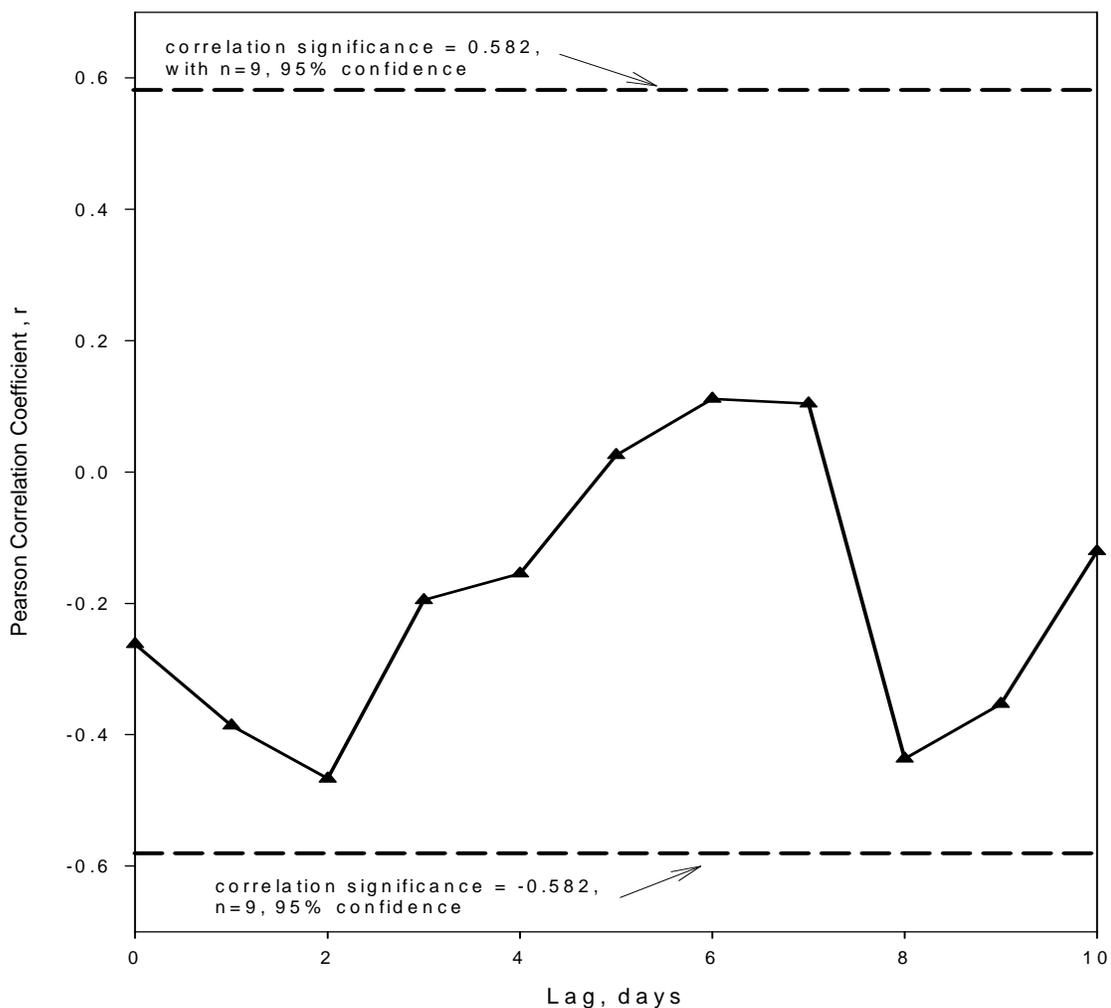
HJ Andrews 2005 WS-1 VPD Comparison: Tower to Plot Average



Comparison of VPD of tower with average of VPD of transect plots. The patterns are obviously highly correlated (Pearson's correlation coefficient = 0.99)

Figure 3.6: Keeling Intercepts correlated to Tower VPD

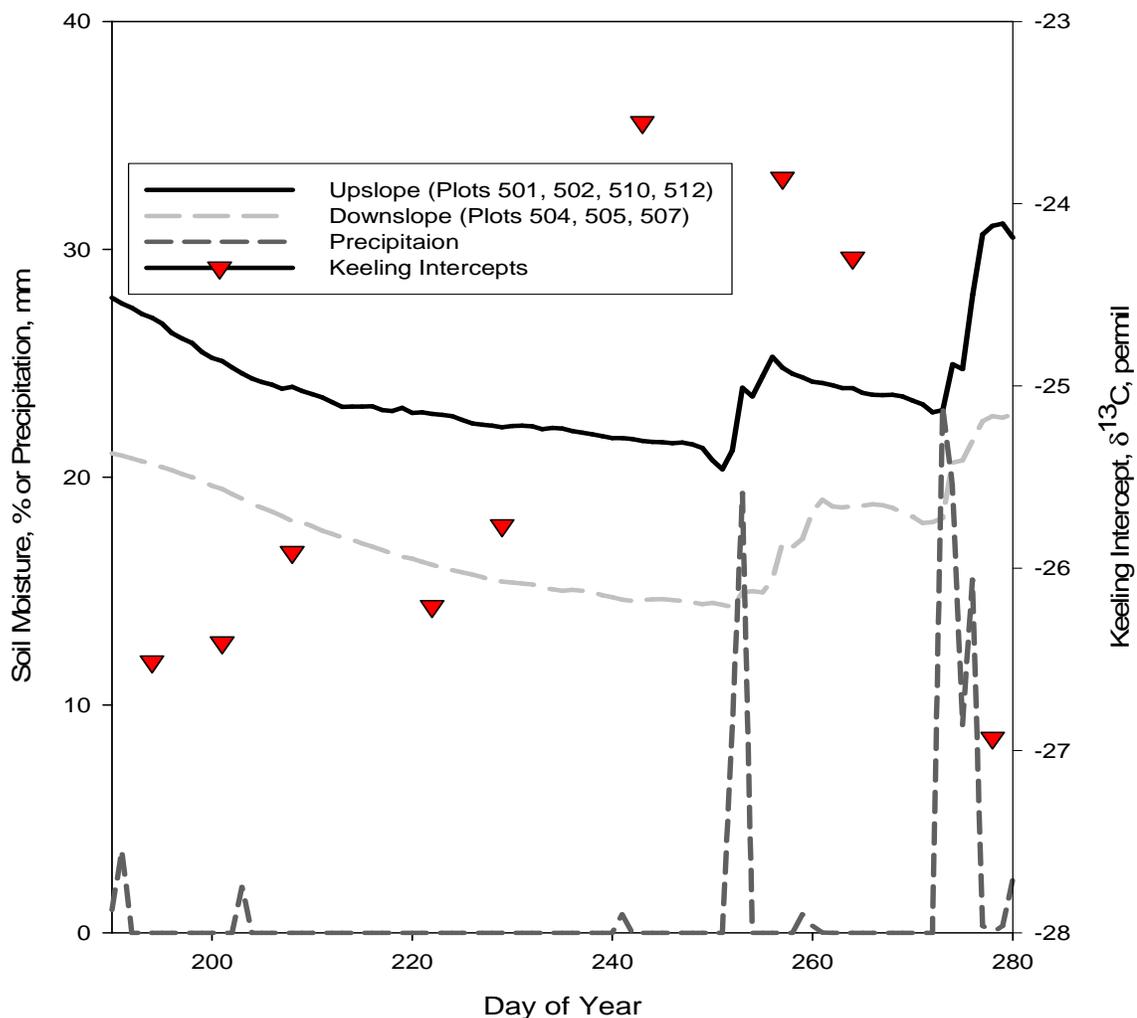
Keeling Intercepts correlated to vapor pressure deficit



Growing season Keeling intercepts correlated to Tower VPD with temporal lag. The Pearson's correlation coefficient for Keeling intercept was correlated to the corresponding values of Tower VPD, starting with a lag of 0 days and ending with a 10 day lag (Keeling intercept lag behind VPD, assuming that the isotopic response of the ecosystem respiration lagged behind the VPD which may have been responsible for the isotopic changes in respiration). Note that there are no values of the Pearson's correlation coefficient which are significant after applying Helsel and Hirsch's significance test.

Figure 3.7: Comparison of Soil Moisture on upslope plots versus downslope plots

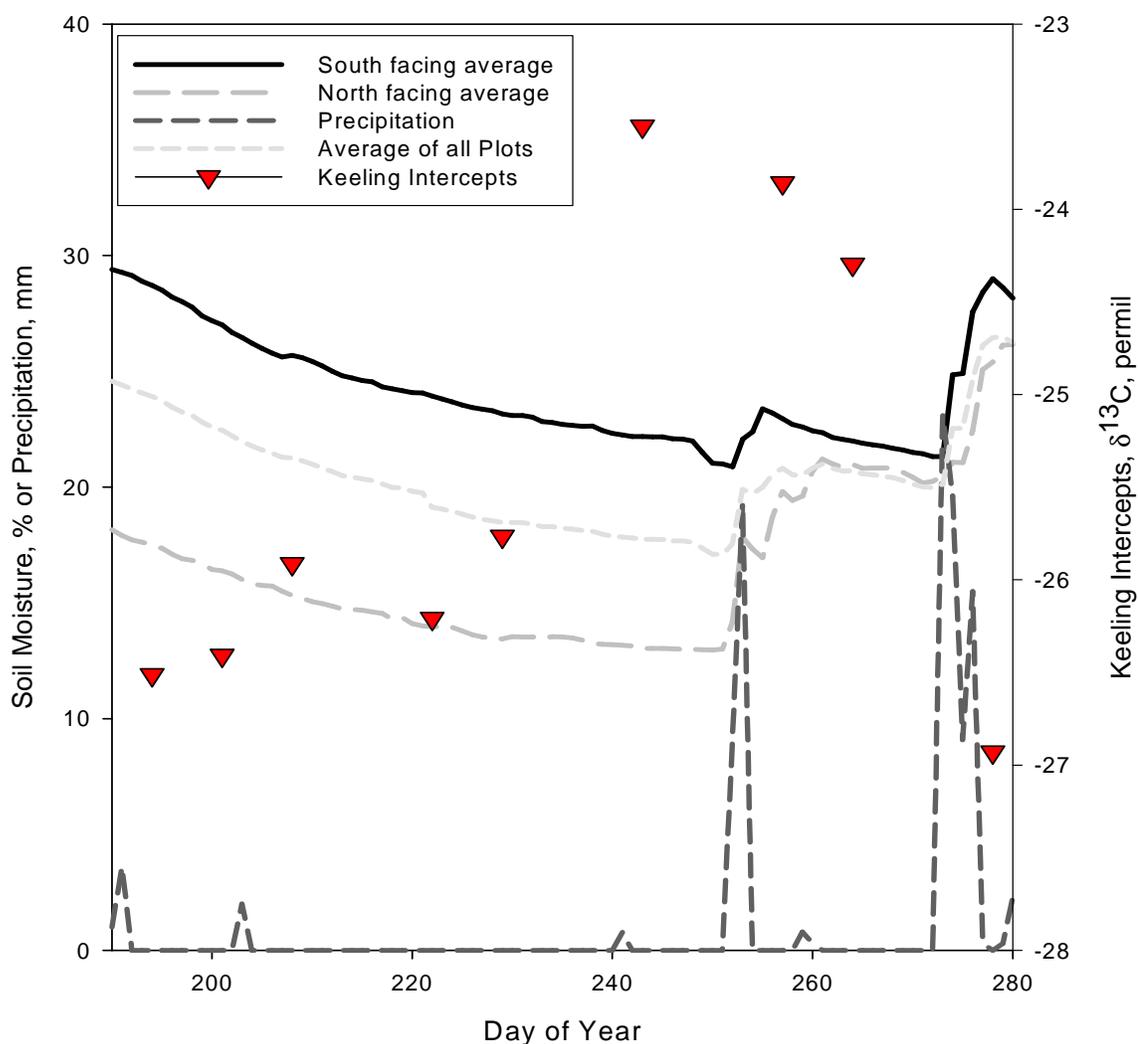
HJ Andrews 2005 WS-1 Soil Moisture: Plot Elevation Comparison



Comparison of soil moisture at plots on grouped by slope position. Soil moisture patterns for upslope and downslope plots show high degree of correlation with each other (Pearson correlation =0.90). Note response of soil moisture to larger precipitation inputs is immediate, whereas for small values of precipitation, no effect on soil moisture is obvious. This is probably due to the fact that the canopy did not saturate in the smaller precipitation events, thus effectively no rainfall reached the soil.

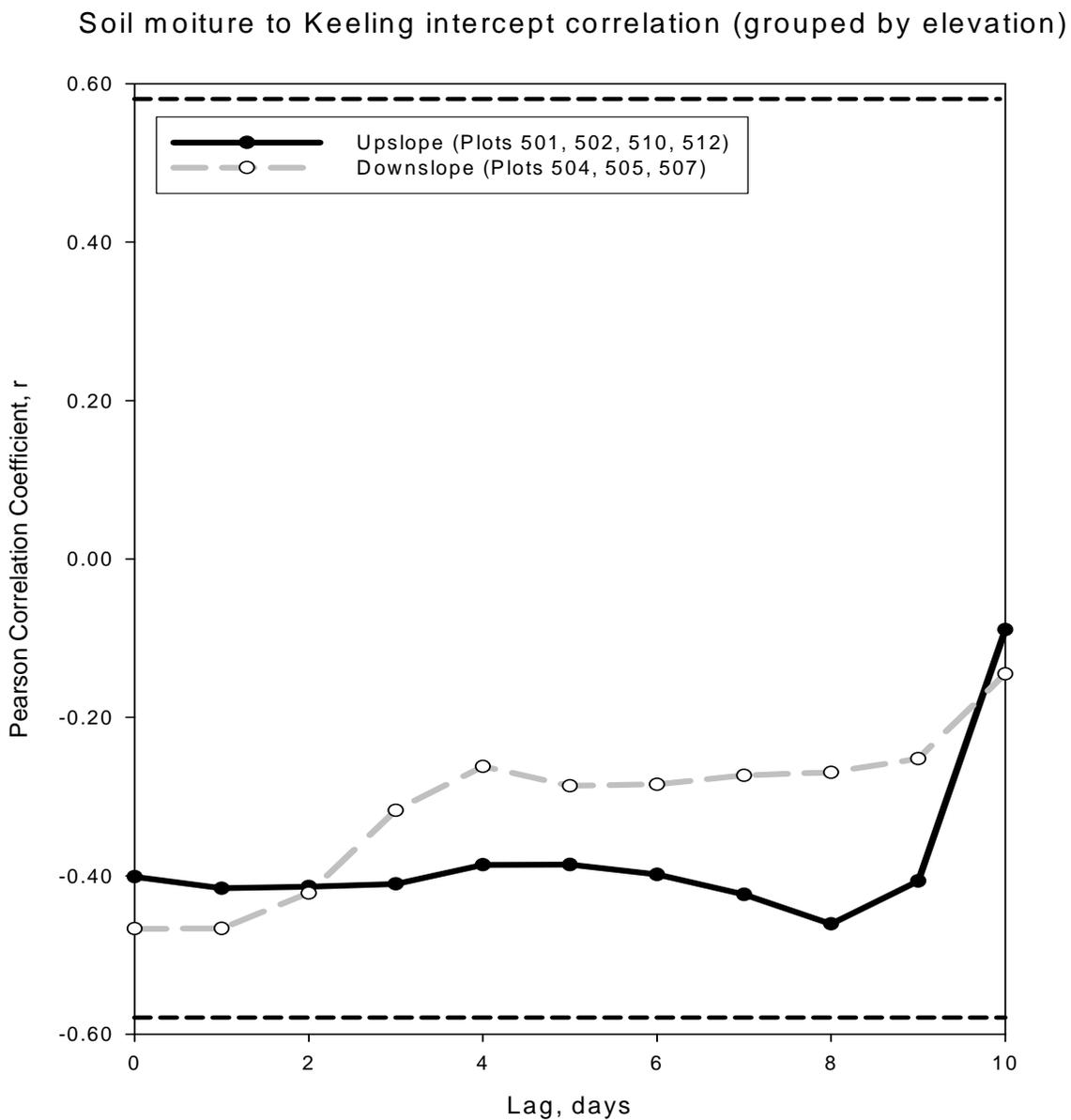
Figure 3.8: Comparison of Soil Moisture on north-facing to south-facing plots

HJ Andrews 2005 WS-1 Soil Moisture: Plot Direction Comparison



Comparison of soil moisture at plots on grouped by slope directional orientation. Soil moisture patterns for south-facing and north-facing plots show low degree of correlation with each other (Pearson correlation = 0.58). Again, the same pattern of soil moisture changes with magnitude of precipitation events is seen.

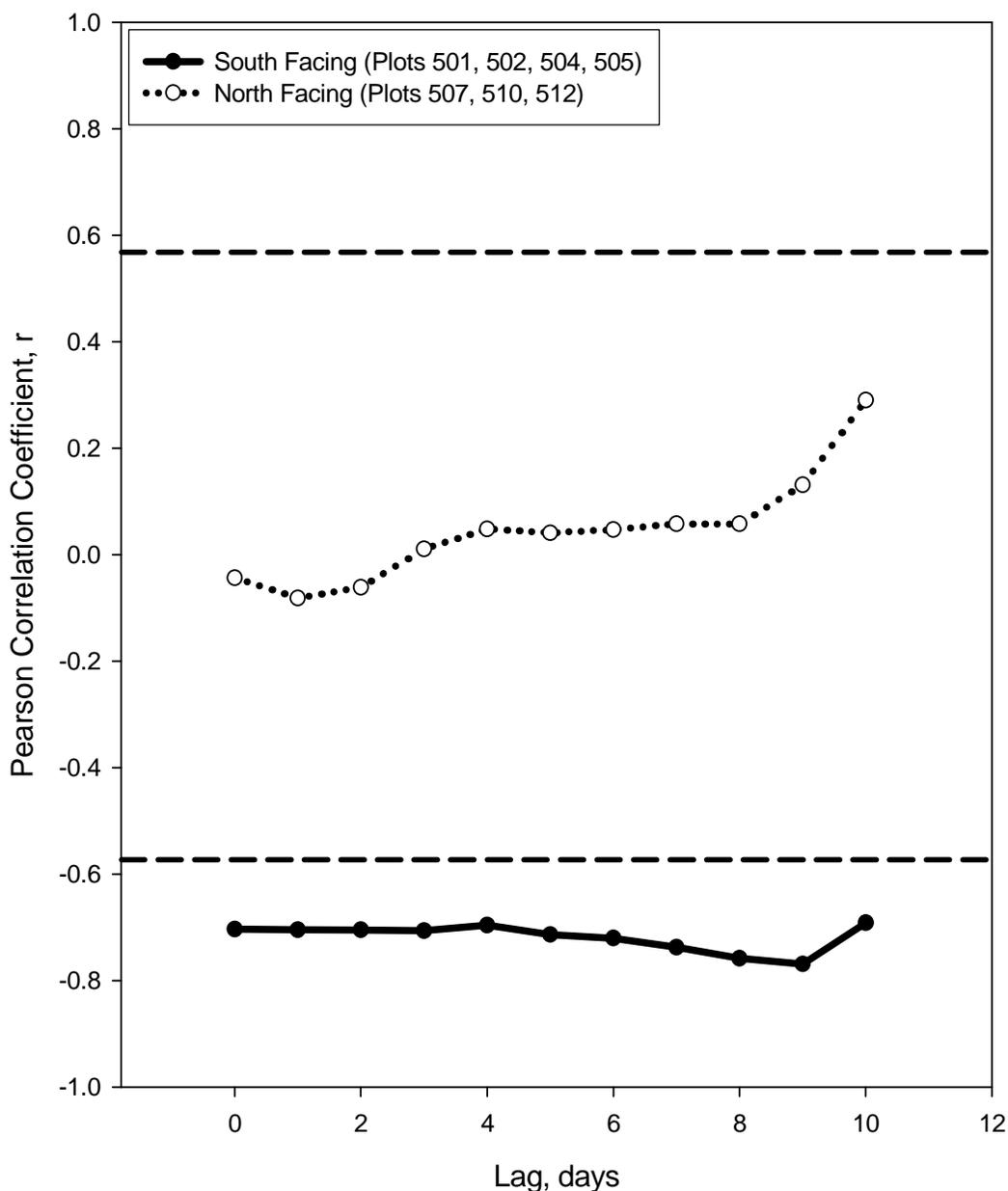
Figure 3.9: Keeling intercept correlations to soil moisture for upslope and downslope plots



Keeling intercept correlated to soil moisture grouped by elevation. Note that there are no significant correlations for either upslope grouped plots or downslope grouped plots after applying the Helsel and Hirsch significance test.

Figure 3.10: Keeling intercept correlations to soil moisture for north-facing and south-facing plots

Soil moisture to Keeling intercept correlations (grouped by Plot direction)



Keeling intercept correlated to soil moisture grouped by plot direction. Note that there are no significant correlations for the north-facing grouped plots, whereas all of the correlations for the south-facing grouped plots are significant. Also, the data does not support a significant lag effect. This lack of lag effect is likely an artifact of the slow rate of change of soil volumetric water content.

Figure 3.11: Table of soil depth and ecosystem respiration

Measurement	South Facing	North Facing
Soil Depth to bottom of B Horizon	105cm	81cm
Foliar Respiration	-7.8 mmolm ⁻² s ⁻¹	-4.2 mmolm ⁻² s ⁻¹
Soil Respiration	-7.0 mmolm ⁻² s ⁻¹	-3.0 mmolm ⁻² s ⁻¹

Table of soil depth and ecosystem respiration values. These data were supplied by other members of the HJ Andrews Airshed Project team. The foliar respiration values are averages of 2006 measurements while the soil respiration measurements are averages of combined 2005 and 2006 values. The soils depth measurements were averaged from measurements made in 2005 from 4 soil pits dug at 2 sites per slope.

References:

Baldocchi, D. 1997. Measuring and modeling carbon dioxide and water vapour exchange over a temperate broad-leaved forest during the 1995 summer drought. *Plant, Cell and Environment* **20**:1108-1122.

Baldocchi, D. 2003, Assessing the eddy covariance technique for evaluating carbon dioxide exchange rates of ecosystems: past, present, and future, *Global Change Biology*, Vol 9, Page 479

Barbour, M.M., Hunt, J.E, Dungan, R. J., Turnbull, M.H., Brailsford, G.W., Farquhar, G.D., Whitehead, D., Variation in the degree of coupling between d13C of phloem sap and ecosystem respiration in two mature *Nothofagus* forests, *New Phytologist*, Volume 166, Page 497, May 2005.

Bhupinderpal-Singh, Nordgren A, Ottosson Lofvenius M, Hogberg MN, Mellander P-E, Hogberg P, 2003. Tree root and soil heterotrophic respiration as revealed by girdling of boreal Scots pine forest: extending observations beyond the first year, *Plant, Cell, Environ.*, 26, 1287-1296

Black, T. A., G. d. Hartog, H. H. Neumann, P. D. Blanken, P. C. Yang, C. Russell, Z. Nesic, X. Lee, S. G. Chen, R. Staebler, and M. D. Novak. 1996. Annual cycles of water vapour and carbon dioxide fluxes in and above a boreal aspen forest. *Global Change Biology* **2**:219-229.

Bowling, D.R., McDowell, N.G., Bond, B.J., Law, B.E., Ehleringer, J.R., 13C content of ecosystem respiration is linked to precipitation and VPD, *Oecologia* (2002), 131:113-124.

Brooks, J.R., Meinzer, F.C., Warren, J.M., Domec, J.C., Coulombe, R., 2006, Hydraulic redistribution in a Douglas-fir forest: lessons from system manipulations, *Plant, Cell and Environment*, (2006) 138-150

Bowling, D.R., Tans, P.P., Monson, R.K., 2001, Partitioning net ecosystem carbon exchange with isotopic fluxes of CO₂, *Global Change Biology*, Vol 7, Page 127

Buchman, N., and J. A. Ehleringer. 1998, CO₂ concentration profiles, and carbon and oxygen isotopes in C₃ and C₄ crop canopies, *Agricultural and Forest Meteorology* 89:45-58.

Buchman, N, Kao, W, Ehleringer, J, 1997, Influence of stand structure on Carbon 13 of vegetation, soils and canopy air within deciduous and evergreen forests in Utah, United States, *Oecologia*, Vol 10, number 1

- Burton, A.J., Pregitzer, K.S., Zogg, G.P., Zak, D.R., 1998, Drought reduces root respiration in sugar maple forests, *Ecological Applications*, Vol. 8, No.3
- Condon, A.G., Richards, R.A., Farquhar, G.D., 1992, The effect of variation in soil water availability, vapour pressure deficit and nitrogen nutrition on carbon isotope discrimination in wheat, *Australian Journal of Agricultural Research* 43(5) 935 - 947
- Chambers, J.Q., Tribuzy, E.S., Toledo, L.C., Crisp, B.F., Higuchi, N., Dos Santos, J., Araujo, A.C., Kruijt, B., Nobre, A.D., Trumbore, S.E., 2004, Respiration from a tropical forest ecosystem: Partitioning of sources and low carbon use efficiency, *Ecological Applications*, 14(4) Supplement 2004, 72-88
- Chernick, M.R., *Bootstrap methods: a practitioner's guide*, 1999, John Wiley & Sons.
- Ehleringer, J. R., A. E. Hall, and G. D. Farquhar. 1992. Stable isotopes and plant carbon-water relations. Academic Press, San Diego, CA.
- Eckblad A, Högberg P. 2001. Natural abundance of ^{13}C in CO_2 respired from forest soils reveals speed of link between tree photosynthesis and root respiration, *Oecologia* 127: 305-308
- Farquhar, G.D., Ehleringer, J.R., Hubick, K.T., 1989. Carbon isotope discrimination and photosynthesis. *Annu. Rev. Physiol. Plant Mol. Biol.*, 40, 503-537
- Flanagan, L. B., and J. A. Ehleringer. 1998. Ecosystem - atmosphere CO_2 exchange: interpreting signals of change using stable isotope ratios, *Trends in Ecology and Evolution* 13:10-14.
- Fessenden, J.E., Ehleringer, J.R., 2003, Temporal variation in $\delta^{13}\text{C}$ of ecosystem respiration in the Pacific Northwest: links to moisture stress, *Oecologia* 136:129-136
- Fleagle, R. G. 1950. A theory of air drainage. *Journal of Meteorology* 7:227-232.
- Gifford, R.M., 1994, The Global Carbon Cycle: a Viewpoint on the Missing Sink, *Australian Journal of Plant Physiology* 21(1) 1 - 15
- Harmon, M. E., K. Bible, M. G. Ryan, D. C. Shaw, H. Chen, J. Klopatek, and X. Li. 2004. Production, respiration, and overall carbon balance in an old-growth *Pseudotsuga-Tsuga* forest ecosystem. *Ecosystems* 7:498-512.
- Helsel, D.R., and R.M. Hirsch, *Statistical Methods in Water Resources*, Elsevier, Science Publishers, Amsterdam, The Netherlands, 1992.

Jones, J.A. and Post, D.A., 2004, Seasonal and successional streamflow response to forest cutting and regrowth in the northwest and eastern United States, *Water Resources Research*, Vol. 40, W05203

Keeling, C. D. 1958. The concentration and isotopic abundances of atmospheric carbon dioxide in rural areas, *Geochimica et Cosmochimica Acta* 13:322-334.

Manins, P. C., and B. L. Sawford. 1979. Katabatic winds: A field study. *Quarterly Journal of the Meteorological Society* **105**:1011-1025.

Martin, M.E., Aber, J.D., 1997, High Spectral Resolution Remote Sensing of Forest Canopy Lignin, Nitrogen, and Ecosystem Processes, *Ecological Applications*, Vol. 7, No. 2, pp. 431-443

Monteith, J. L., and M. H. Unsworth. 1990. *Principles of Environmental Physics*, 2nd edition. Edward Arnold, New York.

Moore, G.W., Bond, B.J., Jones, J.A., Phillips, N., Meinzer, F.C., 2004, Structural and compositional controls on transpiration in 40 and 450-year-old riparian forests in western, Oregon, USA, *Tree Physiology* 24, 481-491

Lavigne, M. B., D. E. Anderson, M. G. Ryan, D. E. Anderson, D. D. Baldocchi, P. M. Crill, D. R. Fitzjarrald, M. L. Goulden, S. T. Gower, J. M. Massheder, J. H. McCaughey, M. Rayment, and R. G. Striegl. 1997. Comparing nocturnal eddy covariance measurements to estimates of ecosystem respiration made by scaling chamber measurements at six coniferous boreal sites. *Journal of Geophysical Research* **105**:28977-28985.

Mortazavi, B, Chanton, JP, 2002, Carbon isotopic discrimination and control of nighttime canopy $\delta^{18}\text{C-CO}_2$ in a pine forest in the southeastern United States, *Global Biogeochemical Cycles*, Vol. 16, No. 1

Ometto, J.P., Flanagan, L.B., Martinelli, L.A., Moreira, M.Z., Ehleringer, J.R., 2002, Carbon isotope discrimination in forest and pasture ecosystems of the Amazon Basin, Brazil, *Global Biogeochemical Cycles*, Vol.16, No. 4, 1109

Pataki, D. E., J. A. Ehleringer, L. B. Flanagan, D. Yakir, D. R. Bowling, C. Still, N. Buchmann, J. O. Kaplan, and J. A. Berry. 2003, The application and interpretation of Keeling plots in terrestrial carbon cycle research, *Global Biogeochemical Cycles* 17:1022.

Pypker, T.G., Unsworth, M.H., Mix, A.C., Rugh, W., Ocheltree, T., Alstad, K., Bond, B.J., 2006, Using nocturnal cold air drainage flow to monitor ecosystem processes in

complex terrain: a pilot study on the carbon isotopic composition and advection of ecosystem respiration, in press (Ecol. Apps)

Pypker, TG, Bond, BJ, Unsworth, MH, The role of epiphytes in the interception and evaporation of rainfall in old-growth Douglas-fir forests in the Pacific Northwest, *Agricultural and Forest Meteorology* 130 (2005) 113-129

Rochette, P., L.B. Flanagan, E.G. Gregorich. 1999. Separating soil respiration into plant and soil components using analyses of the natural abundance of carbon-13. *Soil Science Society of America Journal* 63: 1207-1213

Sanford A.P., Jarvis, P.F., 1986, Stomatal responses to humidity in selected conifers, *Tree Physiol* 2:89-103

Schimel, D. S., T. G. F. Kittel, S. W. Running, R. K. Monson, A. A. Turnipseed, and D. E. Anderson. 2002. Carbon sequestration studies in Western US mountains. *EOS* **83**:445-448.

Staebler, R. M., and D. R. Fitzjarrald. 2004. Observing subcanopy CO₂ advection. *Agricultural and Forest Meteorology* **122**:139-156.

Swanson, F. J., and M. E. James. 1975. Geology and geomorphology of the H.J. Andrews Experimental Forest, Western cascades, Oregon. Pacific Northwest Forest and Range Experiment Station, Forest Service, U.S.D.A Research Paper PNW-188:1-14.

Whiteman, C. D. 1990. Observations of thermally developed wind systems in mountainous terrain. Pages 5-42 in W. Blumen, editor. *Atmospheric processes over complex terrain*. American Meteorological Society, Boston.

Yakir, D., Sternberg, L., 2000, The use of stable isotopes to study ecosystem gas exchange, *Oecologia*, Vol 123, No. 3

Zavitkowski, J., Ferrell, W.K, 1968, Effect of drought upon rates of photosynthesis, respiration, and transpiration of seedlings of two ecotype of Douglas-fir, *Botanical Gazette*, Vol. 129, No. 4 (Dec., 1968), pp. 346-350

Chapter 4

SUMMARY AND CONCLUSIONS

Warming of the terrestrial biosphere due to the anthropogenic addition of carbon dioxide to the earth's atmosphere is becoming a major focus of scientific inquiry. Predictions of the extent of this warming are hampered by uncertainty in the ability of the earth's ecosystems to counteract this effect by sequestering carbon dioxide by increases in the mass of vegetation. Measurement of the carbon isotopic composition of respired CO₂ ($\delta^{13}\text{C}_{\text{R-eco}}$) is becoming increasingly important to ecosystem studies because the information contained in this respiration is a good indicator of ecosystem stress and productivity. This study was conducted as part of a larger research project aimed at developing and testing the capacity for measuring $\delta^{13}\text{C}_{\text{R-eco}}$ in a small, steeply-sloped watershed in western Oregon. The goals of this study were: (1) to develop and test an automated system for sampling nocturnal air to be analyzed in a laboratory for isotopic composition; (2) to collect samples of the atmosphere from the nocturnal cold air drainage of a steeply sloped watershed in the HJ Andrews Experimental Forest and bring those samples back to a laboratory to be analyzed for CO₂ concentration and carbon isotope composition; and (3) to conduct an initial analysis of the relationship between $\delta^{13}\text{C}_{\text{R-eco}}$ for the first year of deployment and two environmental forcing factors, soil moisture and atmospheric VPD.

The automated sampling system designed and built for this project provided significant advantages over previous methods of sample collection. The design objectives for this automated system included light weight, portability, reliability, fast dynamic

response, unattended operation, and the ability to capture, transport, and store samples over several days with no loss of data integrity with respect to the critical measurement values of CO₂ concentration and carbon isotopic composition. The system was deployed in rugged terrain on a weekly basis throughout the 2005 sampling system by a variety of personnel, demonstrating its light weight and portability. With respect to maximum sampling rate, the dynamic response testing of the automated sampling system provides evidence that the system can exceed the performance of systems which use glass flasks to contain samples and should be able to support rapid sampling in research situations which demand such capability. For this capability to be fully exploited, some changes to the control architecture would need to be implemented, but those changes are relatively minor and could be accomplished without much difficulty. The system reliably operated throughout the nighttime sampling period without technician intervention. Testing showed that the system provides for storage of samples in the sample loops for at least three days without significant change in [CO₂] or $\delta^{13}\text{C}$. This is a clear advantage for researchers who endeavor to set up such a system and then allow it to perform its function without supervision. The automated system consistently outperformed the hand sampling process it replaced and compared favorably with automated sampling systems described by other researchers with respect to both accuracy and precision of the datasets which were produced from the air samples collected. The automated system showed itself to be reliable throughout the sampling season, as evidenced by the positive results of the weekly system plumbing integrity tests. In conclusion, the system met all design objectives and functioned as intended, providing a collection of air samples which were

successfully analyzed for carbon isotopic composition and carbon dioxide concentration in the COAS Mass Spectrometer Laboratory.

Although the system met its design objectives and provided the basis for the primary dataset for this project, there is room for improvement in this system. The system would benefit from a further reduction in weight, making the system even more portable and thus easier to deploy in remote and rugged locations. This could be accomplished by reducing the size and changing the materials for both the sample valve enclosure and the control enclosure. The Valco sample collection valve is capable of producing air samples which accurately reflect the source, but this valve is not particularly well suited to the harsh environment where it was deployed. This situation could be remedied by either using a sample collection valve from a different manufacturer or by providing a more robust enclosure for the current valve. No suitable replacement valve candidates have been identified and providing a more robust enclosure is not a trivial undertaking due to the need for accessibility of the valve for laboratory analysis. Without design or hardware changes in this area, the weekly integrity testing should be continued to make certain the collected samples accurately reflect the source. With the addition of a wireless interface and transmitter, the system could be made to collect data independently and report back to the experimenter that the system is ready for retrieval. This feature could enhance the ability of a research team to take samples in multiple remote locations with minimal personnel requirements while ensuring data which meets pre-set requirements. A further system enhancement would be a lightweight, portable power source. This feature would decouple the sampling site from the requirement that a suitable source of power at

the site be available. Recent advances in hydrocarbon fuel cell technology may make this enhancement possible in the near future.

The weekly data stream from this automated system was analyzed for carbon isotope ratio and CO₂ concentration between May and November 2005. These data were used to determine the value of the carbon isotopic composition of ecosystem respiration ($\delta^{13}\text{C}_{\text{R-eco}}$) using a two end member mixing model known as a Keeling plot. These datasets were filtered to remove unsuitable data, resulting in a plot of seasonal $\delta^{13}\text{C}_{\text{R-eco}}$ from July 13, 2005 to October 5, 2005. Comparison of this plot with those of other studies show that the seasonal pattern of isotopic composition of ecosystem respiration observed in this study is consistent with the patterns observed in other studies for forests of similar vegetative compositions and climates. The isotopic pattern of ecosystem respiration shows a trend toward heavier isotopes as the summer progresses and the forest dries out. This pattern is consistent with a forest transitioning into a drought stressed condition. As the rains began to return in the fall, the isotopic trend returned to the lighter values, removing the drought stress condition.

The correlations between the seasonal pattern of $\delta^{13}\text{C}_{\text{R-eco}}$ and the environmental forcing variables of VPD and soil volumetric water content provided interesting results. VPD was not significantly correlated to seasonal Keeling intercepts, unlike results from previous studies. The correlations between soil volumetric water content and seasonal Keeling intercepts were strong for soil volumetric water content on transect plots facing south, but there were no significant correlations with soil volumetric water content on north-facing slopes. When soil volumetric water content was grouped by elevation, no

significant correlations were found. Past research has found lags of between 1 and 4 days between VPD and root respiration changes in $\delta^{13}\text{C}$ and 5-10 days between VPD and $\delta^{13}\text{C}_{\text{R-eco}}$, yet the data from this study do not show significant lags between either VPD or soil volumetric water content and seasonal Keeling plot intercepts. The lack of correlation between VPD and seasonal Keeling intercepts likely lies in the timing of the sampling period. In the early part of the growing season when soil moisture is plentiful, stomatal conductance is dominated by VPD. As the summer progresses and the soil moisture declines, stomatal conductance is dictated by soil moisture content. Unfortunately, the first Keeling intercepts of the 2005 season for this study did not occur until mid July, when the forest was already showing signs of drought stress due to soil moisture deficit. Had Keeling intercepts been available for earlier in the season, it is likely that the correlation between VPD and Keeling intercepts would have been much stronger. Data from the 2006 season reinforce this conjecture. The lack of correlation between soil volumetric water content grouped by elevation and seasonal Keeling plot intercepts is somewhat surprising and certainly non-intuitive. The patterns for soil volumetric content for upper and lower elevation plots were very well correlated, which implies that either both sets of data should be correlated to Keeling intercepts or neither set of data should be correlated. It turns out that the later is true. The correlation between south-facing plots soil volumetric water content and seasonal pattern of Keeling intercepts is likely due to the fact that the soil on the south slope is retains a large pool of water for vegetation to draw from. This larger pool allows the south-facing slope vegetation to continue to assimilate carbon well into the year, with an associated high rate

of respiration. The north-facing slope does not enjoy this luxury of a large pool of water and thus goes into drought stress earlier in the season, which reduces both its assimilation and respiration rates. The carbon isotope respiration signal collected at the base of the watershed was truncated in the early part of the 2005 season, when the respiration rate of the north slope was most probably equal to that of the south slope. It is likely that the north slope was simply not contributing much respiration to the isotope signal being gathered at the base of the watershed during the study period and thus was under-represented in the correlation. Anecdotal foliar and soil respiration information provided by other researchers lend support to this hypothesis.

Although the seasonal pattern of Keeling intercepts provided by this project was adequate in reaching the conclusions provided, there is room for improvement. A larger number of data points for the seasonal pattern would help to lower the threshold of correlation considered significant. For instance, increasing the number of Keeling intercepts for the seasonal pattern from nine to 27 would reduce the threshold of correlation significance from ± 0.58 to ± 0.32 . Increasing the number of weekly samples would also greatly improve the analysis if many of those samples could be produced early in the growing season. As seen from the analysis associated with this study, correlations of VPD and north-facing plots soil volumetric water content would likely have been significant had there been early season Keeling intercepts to compare with those variables.

Another aspect of the project worthy of consideration is the positions of the plots for gathering the environmental forcing variable data and the amount of data generated. It

would have been very informative to have plots in multiple locations around the watershed to verify the results obtained on the transect. Also, data on soil depth was minimal, as was information on soil and foliar respiration. Although it is clear that there are resource constraints which must be considered when addressing which data to pursue and at what volume, for the purposes of this investigation it would have been advisable to increase the efforts in these critical areas.

Aside from the insights gained from working on this project by the author, there have been other valuable lessons learned. I started this project with virtually no experience in ecosystem science. Although my background as an engineer provided me with some of the analytical tools necessary to contribute to such a study, I was not prepared for the level of complexity of these forest ecosystems. The process of producing this thesis has been both difficult and humbling, and I consistently underestimated the challenges of producing this document. All in all, I have gained new appreciation for the difficulties faced by the scientists attempting to advance this important field and I hope to be able to support them in their future endeavors.

Bibliography

Baldocchi, D. 1997. Measuring and modeling carbon dioxide and water vapour exchange over a temperate broad-leaved forest during the 1995 summer drought. *Plant, Cell and Environment* **20**:1108-1122.

Baldocchi, D. 2003, Assessing the eddy covariance technique for evaluating carbon dioxide exchange rates of ecosystems: past, present, and future, *Global Change Biology*, Vol 9, Page 479

Barbour, M.M., Hunt, J.E, Dungan, R. J., Turnbull, M.H., Brailsford, G.W., Farquhar, G.D., Whitehead, D., Variation in the degree of coupling between $\delta^{13}\text{C}$ of phloem sap and ecosystem respiration in two mature *Nothofagus* forests, *New Phytologist*, Volume 166, Page 497, May 2005.

Bhupinderpal-Singh, Nordgren A, Ottosson Lofvenius M, Hogberg MN, Mellander P-E, Hogberg P, 2003. Tree root and soil heterotrophic respiration as revealed by girdling of boreal Scots pine forest: extending observations beyond the first year, *Plant, Cell, Environ.*, 26, 1287-1296

Black, T. A., G. d. Hartog, H. H. Neumann, P. D. Blanken, P. C. Yang, C. Russell, Z. Nestic, X. Lee, S. G. Chen, R. Staebler, and M. D. Novak. 1996. Annual cycles of water vapour and carbon dioxide fluxes in and above a boreal aspen forest. *Global Change Biology* **2**:219-229.

Bowling, D.R., McDowell, N.G., Bond, B.J., Law, B.E., Ehleringer, J.R., 2002, ^{13}C content of ecosystem respiration is linked to precipitation and vapor pressure deficit, *Oecologia* (2002), 131:113-124

Bowling DR, Sargent SD, Tanner BD, and Ehleringer JR, 2003, Tunable diode laser absorption spectroscopy for stable isotope studies of ecosystem-atmosphere CO_2 exchange, *Agricultural and Forest Meteorology.*, 118, 1-19.

Bowling, D.R., Tans, P.P., Monson, R.K., 2001, Partitioning net ecosystem carbon exchange with isotopic fluxes of CO_2 , *Global Change Biology*, Vol 7, Page 127

Brooks, J.R., Meinzer, F.C., Warren, J.M., Domec, J.C., Coulombe, R., 2006, Hydraulic redistribution in a Douglas-fir forest: lessons from system manipulations, *Plant, Cell and Environment*, (2006) 138-150

Buchman, N., and J. A. Ehleringer. 1998, CO_2 concentration profiles, and carbon and oxygen isotopes in C_3 and C_4 crop canopies, *Agricultural and Forest Meteorology* 89:45-58.

- Buchman, N, Kao, W, Ehleringer, J, 1997, Influence of stand structure on Carbon 13 of vegetation, soils and canopy air within deciduous and evergreen forests in Utah, United States, *Oecologia*, Vol 10, number 1
- Buchman N, Guehl, JM, Barigah TS, Ehleringer JR, 1997, Interseasonal comparison of CO₂ concentrations, isotopic composition, and carbon dynamics in an Amazonian rainforest (French Giana), *Oecologia*, 110: 120-131
- Burton, A.J., Pregitzer, K.S., Zogg, G.P., Zak, D.R., 1998, Drought reduces root respiration in sugar maple forests, *Ecological Applications*, Vol. 8, No.3
- Condon, A.G., Richards, R.A., Farquhar, G.D., 1992, The effect of variation in soil water availability, vapour pressure deficit and nitrogen nutrition on carbon isotope discrimination in wheat, *Australian Journal of Agricultural Research* 43(5) 935 - 947
- Chambers, J.Q., Tribuzy, E.S., Toledo, L.C., Crisp, BF, Higuchi, N., Dos Santos, J., Araujo, A.C., Kruijt, B., Nobre, A.D., Trumbore, S.E., 2004, Respiration from a tropical forest ecosystem: Partitioning of sources and low carbon use efficiency, *Ecological Applications*, 14(4) Supplement 2004, 72-88
- Chernick, M.R., *Bootstrap methods: a practitioner's guide*, 1999, John Wiley & Sons.
- Ekblad A, Höglberg P. 2001. Natural abundance of ¹³C in CO₂ respired from forest soils reveals speed of link between tree photosynthesis and root respiration, *Oecologia* 127: 305-308
- Ehleringer, JR, 1993. Variation in leaf carbon isotope discrimination in *Encelia farinosa*: implications for growth, competition, and drought survival, *Oecologia*, Volume 95, number 3
- Ehleringer JR, Cook CS, 1997, Carbon and oxygen isotope ratios of ecosystem respiration along an Oregon conifer transect, preliminary observations based upon small-flask sampling, *Tree Physiology*, 18, 513-519
- Ehleringer, J. R., A. E. Hall, and G. D. Farquhar. 1992. *Stable isotopes and plant carbon-water relations*. Academic Press, San Diego, CA.
- Farquhar, G.D., Ehleringer, J.R., Hubick, K.T., 1989. Carbon isotope discrimination and photosynthesis. *Annu. Rev. Physiol. Plant Mol. Biol.*, 40, 503-537
- Fessenden, JE, Ehleringer, JR, 2003, Temporal variation in $\delta^{13}\text{C}$ of ecosystem respiration in the Pacific Northwest: links to moisture stress, *Oecologia* 136:129-136
- Flanagan, L. B., and J. A. Ehleringer. 1998. Ecosystem - atmosphere CO₂ exchange: interpreting signals of change using stable isotope ratios, *Trends in Ecology and Evolution* 13:10-14.

- Flanagan, L.B., J.R. Brooks, G.T. Varney, S.C. Berry, J.R. Ehleringer Carbon isotope discrimination during photosynthesis and the isotope ratio of respired CO₂ in boreal forest ecosystems, *Global Biogeochemical Cycles* 96GB02345
Vol. 10 , No. 4 , p. 629
- Flanagan, L.B., Kubien, D.S., Ehleringer, J.R., 1999, Spatial and temporal variation in the carbon and oxygen stable isotope ratio of respired CO₂ in a boreal forest ecosystem, *Tellus B*, Vol 51, April 1999
- Fleagle, R. G. 1950. A theory of air drainage. *Journal of Meteorology* 7:227-232.
- Gifford, R.M., 1994, The Global Carbon Cycle: a Viewpoint on the Missing Sink, *Australian Journal of Plant Physiology* 21(1) 1 – 15
- Harmon, M. E., K. Bible, M. G. Ryan, D. C. Shaw, H. Chen, J. Klopatek, and X. Li. 2004. Production, respiration, and overall carbon balance in an old-growth *Pseudotsuga-Tsuga* forest ecosystem. *Ecosystems* 7:498-512.
- Helsel, D.R., and R.M. Hirsch, *Statistical Methods in Water Resources*, Elsevier, Science Publishers, Amsterdam, The Netherlands, 1992.
- Jones, J.A. and Post, D.A., 2004, Seasonal and successional streamflow response to forest cutting and regrowth in the northwest and eastern United States, *Water Resources Research*, Vol. 40, W05203
- Keeling CD, 1958, The concentration and isotopic abundances of atmospheric carbon dioxide in rural areas, *Geochim Cosmochim Acta*. 13: 322-334.
- Keeling CD, 1961, The concentration and isotopic abundance of carbon dioxide in rural and marine air, *Geochim Cosmochim Acta*. 24: 277-298.
- Knohl, A., Werner, R.A., Brand, W.A., Buchman, N., (2005), Short-term variations in $\delta^{13}\text{C}$ of ecosystem respiration reveals link between assimilation and respiration in a deciduous forest, *Oecologia* ,142: 70–82
- Lavigne, M. B., D. E. Anderson, M. G. Ryan, D. E. Anderson, D. D. Baldocchi, P. M. Crill, D. R. Fitzjarrald, M. L. Goulden, S. T. Gower, J. M. Massheder, J. H. McCaughey, M. Rayment, and R. G. Striegl. 1997. Comparing nocturnal eddy covariance measurements to estimates of ecosystem respiration made by scaling chamber measurements at six coniferous boreal sites. *Journal of Geophysical Research* 105:28977-28985.
- Lin, G., Ehleringer, J.R., 1997, Carbon isotopic fractionation does not occur during dark respiration in C₃ and C₄ plants, *Plant Physiology*, Vol 114, Issue 1, 391-394

- Manins, P. C., and B. L. Sawford. 1979. Katabatic winds: A field study. *Quarterly Journal of the Meteorological Society* **105**:1011-1025.
- Martin, M.E., Aber, J.D., 1997, High Spectral Resolution Remote Sensing of Forest Canopy Lignin, Nitrogen, and Ecosystem Processes, *Ecological Applications*, Vol. 7, No. 2, pp. 431-443
- Meinzer, F.C., Z. Plaut, and N. Z. Saliendra, 1994, Carbon Isotope Discrimination, Gas Exchange, and Growth of Sugarcane Cultivars under Salinity, *Plant Physiol.* 1994 February; 104(2): 521–526
- Meinzer, F.C, Zhu, J., 1998. Nitrogen Stress reduces efficiency of the C_4CO_2 in the concentrating system, and therefore the quantum yield, in *Saccharum* (sugarcane) species, *Journal of Experimental Botany*, Vol. 49, Number 324
- Miller JB, Tans PP, White JWC, Conway TJ, Vaughn BW, 2003, The atmospheric signal of terrestrial carbon isotopic discrimination and its implication for partitioning carbon fluxes, *Tellus* 55B, 197-206
- Monteith, J. L., and M. H. Unsworth. 1990. *Principles of Environmental Physics*, 2nd edition. Edward Arnold, New York.
- Moore, G.W., Bond, B.J., Jones, J.A., Phillips, N., Meinzer, F.C., 2004, Structural and compositional controls on transpiration in 40 and 450-year-old riparian forests in western, Oregon, USA, *Tree Physiology* 24, 481-491
- Mortazavi, B, Chanton, JP, 2002, Carbon isotopic discrimination and control of nighttime canopy $\delta^{18}C-CO_2$ in a pine forest in the southeastern United States, *Global Biogeochemical Cycles*, Vol. 16, No. 1
- Ometto JHB, Flanagan LB, Martinelli LA, Moreira MZ, Higuuchi N, Ehleringer JR, 2002, Carbon isotope discrimination in forest and pasture ecosystems of the Amazon Basin, Brazil, *Global Biogeochemical Cycles*, Vol. 16, No.4
- Pataki DE, Ehleringer JR, Flanagan JB, Yakir D, Bowling DR, Still DJ, Buchman N, Kaplan JO, Berry JA, 2003, The application and interpretation of Keeling plots in terrestrial carbon cycle research, *Global Biogeochemical Cycles*, Vol 17, No.1
- Pypker, T.G., Unsworth, M.H., Mix, A.C., Rugh, W., Ocheltree, T., Alstad, K., Bond, B.J., 2006, Using nocturnal cold air drainage flow to monitor ecosystem processes in complex terrain: a pilot study on the carbon isotopic composition and advection of ecosystem respiration, in press (*Ecol. Apps*)

- Pypker, TG, Bond, BJ, Unsworth, MH, The role of epiphytes in the interception and evaporation of rainfall in old-growth Douglas-fir forests in the Pacific Northwest, *Agricultural and Forest Meteorology* 130 (2005) 113-129
- Rochette, P., L.B. Flanagan, E.G. Gregorich. 1999. Separating soil respiration into plant and soil components using analyses of the natural abundance of carbon-13. *Soil Science Society of America Journal* 63: 1207-1213
- Sanford A.P., Jarvis, P.F., 1986, Stomatal responses to humidity in selected conifers, *Tree Physiol* 2:89-103
- Schauer, AJ, Lai, CT, Bowling, DR, Ehleringer, JR, 2003, An automated sampler for collection of atmospheric trace gas samples for stable isotope analyses, *Ag. And For. Meteorology* (118), 113-124
- Schimel, D. S., T. G. F. Kittel, S. W. Running, R. K. Monson, A. A. Turnipseed, and D. E. Anderson. 2002. Carbon sequestration studies in Western US mountains. *EOS* **83**:445-448.
- Staebler, R. M., and D. R. Fitzjarrald. 2004. Observing subcanopy CO₂ advection. *Agricultural and Forest Meteorology* **122**:139-156.
- Swanson, F. J., and M. E. James. 1975. Geology and geomorphology of the H.J. Andrews Experimental Forest, Western cascades, Oregon. Pacific Northwest Forest and Range Experiment Station, Forest Service, U.S.D.A Research Paper PNW-188:1-14.
- Theis D, Saurer M, Blum H, Frossard E, Siegwolf TW, 2004, A portable automated system for trace gas sampling in the field and stable isotope analysis in the laboratory, *Rapid Communications in Mass Spectrometry*, Vol, 18, Issue 18, 2106-2112
- Trolier M, White JWC, Tans PP, Masarie KA, Gemery PA, 1996, Monitoring the isotopic composition of atmospheric CO₂: Measurements from the NOAA Global Air Sampling Network, *J. Geophys. Res.*, Vol. 101, No. D20
- Tu KP, Brooks PD, Dawson TE. 2001. Using septum-capped vials with continuous-flow isotope ratio mass spectrometric analysis of atmospheric CO₂ for Keeling plot applications. *Rapid Commun. Mass Spectrom.* 15:952-56
- Whiteman, C. D. 1990. Observations of thermally developed wind systems in mountainous terrain. Pages 5-42 in W. Blumen, editor. *Atmospheric processes over complex terrain*. American Meteorological Society, Boston.
- Yakir, D., Sternberg, L., 2000, The use of stable isotopes to study ecosystem gas exchange, *Oecologia*, Vol 123, No. 3

Zavitkowski, J., Ferrell, W.K, 1968, Effect of drought upon rates of photosynthesis, respiration, and transpiration of seedlings of two ecotype of Douglas-fir, Botanical Gazette, Vol. 129, No. 4 (Dec., 1968), pp. 346-350

Zobitz JM, Keener JP, Bowling DR, 2004, Sensitivity analysis and quantification of uncertainty for isotopic mixing relationships in carbon cycle research, American Geophysical Union, Fall Meeting 2004, in pres